

**King Solomon Property, Central Monitor Range,
Northern Nye County, Nevada:
Report on Geologic Mapping and Sampling
July to October 2025**

by

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SUMMARY

The King Solomon property in the central Monitor Range, northern Nye County, Nevada is being explored for both epithermal volcanic-hosted and Carlin-type Au-Ag-Sb mineralization. The property is located (Figure 1) in the central Great Basin's Oligocene caldera belt, which hosts a major epithermal Au-Ag deposit at Round Mountain (>14 Moz Au produced) and significant Carlin-type gold deposits at Eureka (>1.5 Moz Au produced) and Northumberland (>3 Moz Au resources).

King Solomon is underlain by extensive silicic volcanic rocks that are almost certainly associated with Oligocene caldera-related activity. The oldest volcanic rocks are a +1.5km-thick assemblage of successive, moderately welded rhyolitic to dacitic ash-flow tuff that is interpreted to derive from an intracaldera setting. These early tuffs are overlain by abundant, vent and near-vent facies, porphyritic quartz-biotite rhyolite lavas in three geographically distinct centers: Willow Creek in the northeast, House Canyon in the east-central, and Stone Cabin in the south. At least 16 vents for rhyolite lavas occur on or very near the King Solomon property. Rhyolite dikes, small plugs, and polymictic rhyolite breccia bodies related to these centers are abundant in deeper exposures cutting older rocks. These intrusions were a significant focus for late hydrothermal fluids, and it is likely that rhyolite lavas at vents cover similar intrusions. The rhyolite lavas and intrusions were likely localized by a caldera ring fracture and possibly the incipient King Solomon fault zone. A series of unaltered thin ash-flow tuffs overlie altered lavas in the hanging wall of the King Solomon normal fault and are regarded as post-mineral.

Curvilinear patterns of units and faults, younging of units southward, and general southward stratigraphic dips support the hypothesis that the property lies on the southern margin of a domal feature possibly related to caldera resurgence. The King Solomon Paleozoic window (KS window), which exposes probable Devonian through Ordovician lower-plate carbonates and quartzite, forms the core of the domal uplift. The lower-plate strata are structurally inverted in the window due to doming, whereas deformed upper-plate siliciclastic sedimentary rocks lay mostly at lower elevations surrounding the lower-plate core. A major thrust that separates the two domains is consistent with the Roberts Mountains thrust, which is a fundamental control for major Carlin-type deposits elsewhere in Nevada. The lower-plate strata are variable in character and several units present are interpreted as prospective hosts for disseminated Carlin-type gold mineralization including the thin-bedded carbonates of the Pogonip Limestone, Roberts Mountains Formation, and Devonian limestone possibly correlative with either the McColley Canyon formation or its downslope equivalent, the Rabbit Hill Limestone.

Extensive hydrothermal alteration and significant Au mineralization occur in four main areas of King Solomon: 1) the King Solomon mine area, 2) the KS Paleozoic window, 3) the Big Bend area of the King Solomon fault, and 4) West Ridge in the west-central part of the property. In addition, an emerging area of interest to the south (South Tail) contains strong alteration with high Hg along the King Solomon fault that may be a relatively higher-level expression. Only sparse drilling was done at Big Bend and West Ridge.

Exposures at the King Solomon mine are generally strongly altered with volcanics exhibiting moderate to strong sericitic alteration (feldspar- and biotite-destructive) and weak to moderate silicification whereas Paleozoic carbonate rocks are completely replaced by jasperoid quartz in exposures. Minor quartz veins and crackle breccia with open space contain stibnite, stibiconite, and barite. Big Bend is the southern extension of the King Solomon area and coincidentally occurs at the significant bend in the King Solomon fault.

The KS window area exhibits strong alteration, particularly around its fault-bound southern and eastern margins that include mineralized jasperoid and weak to moderate dolomitization and/or decarbonatization. Northerly striking splays of the King Solomon fault and northeast-striking dikes extend into the KS window and focused Carlin-type alteration and mineralization. Like at the King Solomon mine, a distinctive late vug-filling quartz-barite-stibnite stage is associated with jasperoid at the KS window. Steep terrain and overburden along the margins of the KS window obscure the volcanic-sedimentary contact in most places, although where exposed, adjacent volcanics are strongly altered.

The West Ridge area consists of numerous Stone Cabin rhyolite dikes that intrude the early ash-flow tuffs. Large composite rhyolite dikes are moderately to strongly altered to a quartz-sericite-iron oxide assemblage, with fine pyrite and other sulfides present in densely silicified rock. Polymictic intrusive breccia containing abundant juvenile rhyolite clasts and variably altered Paleozoic sedimentary clasts in an altered and commonly mineralized rock flour groundmass.

Mineralization at King Solomon is characterized by elevated Au, Ag, As, Sb, Hg, Tl, and Ba. Elemental concentrations in 213 rock chips from recent mapping have the following mean averages: 217 ppb Au, 12.2 ppm Ag, 345 ppm As, 428 ppm Sb, and 2.4 ppm Hg. Maximum values are: 5.1 ppm Au, 609 ppm Ag, 0.7% As, >1.0% Sb, 110 ppm Tl, and 1.76% Hg. There is a paucity of base metals except for Mo, which has a maximum value of 202 ppm. These values are consistent with both Carlin-type and epithermal mineralization in Nevada but many or most of these samples were from bulk-disseminated styles of mineralization, not veins. Paragenetically, gold tends to associate best with early silicification and/or sericitization whereas Ag and Sb are associated with the late-stage open-space infilling by colloform quartz-barite-stibnite stage.

Geophysical data include fairly detailed airborne magnetics, gravity, and limited CSAMT and IP. The magnetics and gravity data corroborate much of the detailed mapping, lending support to map interpretations but also providing some very interesting additional insights. For example, the KS window exposes lower Paleozoic carbonate rocks that coincide with a robust magnetic high that indicates a large intrusion at depth, which is supported by the mapping of numerous dikes in the area. The interpretation of a large intrusion under the KS window supports the interpretation of a broad, intrusion-sourced dome that also affected Cenozoic volcanic rocks.

At least 103 holes have been drilled by 9 companies at King Solomon for a total of over 41,340 feet (12,600m). The holes were drilled between 1981 and 1996, although all but 6 holes were drilled between 1984 and 1990. Drilling was concentrated in the pediment area from the King Solomon mine northeast to the House Canyon area. Six shallow holes were drilled by UKE in the West Ridge area. Mineralization in the King Solomon mine area remains open to the south and the entire King Solomon fault corridor is open at depth. Challenges will be in defining mineralization that has continuity and high enough grade to support sulfide processing as the weathered profile is thin. Regarding continuity, the potential exists for widespread disseminated Au-Ag mineralization in certain of the carbonate rocks like the Roberts and/or Devonian carbonates (e.g., Bartine member, McColley Canyon Fm.) under both Paleozoic and Cenozoic cover.

New recognition of the project as constituting part of a large silicic magmatic center with an early-stage intracaldera setting followed by voluminous, vent-facies rhyolite volcanism that includes abundant shallow intrusions and associated regional doming or resurgence provides better geological context for future exploration. It is the author's view that there is considerable exploration potential for additional mineralization at King Solomon. The project possesses robust and widespread alteration and at least four areas of surface Au-Ag mineralization are advanced enough for drilling. Two of these areas (King Solomon mine and KS window) have seen significant, albeit shallow, historical drilling whereas two other Au-mineralized areas (Big Bend and West Ridge) have seen only sparse drilling. Seven other prospective areas are less explored and require more detailed mapping and additional sampling to assess their potential for drilling (Appendix G: Target Map).

INTRODUCTION

The King Solomon project is operated by Phenom Resources Corp. (Phenom), a Canadian junior exploration company based in Vancouver, British Columbia with the ticker TSX.V:PHNM. Phenom has lease agreements with various individuals and companies for mineral rights at King Solomon and has expanded the project by staking additional unpatented mining claims. This mapping and sampling effort was contracted through Phenom in March 2025.

The purpose of this report is primarily to describe and interpret findings from the 2025 geologic mapping and sampling at King Solomon. A secondary purpose is to provide a preliminary assessment of exploration potential of the property based not only the current work but much of the archived historical data.

Location, Claims, and Access

Location – The King Solomon property is in the central Monitor Range of northern Nye County in central Nevada (Figure 1). The property lies entirely on land administered by the Humboldt-Toiyabe National Forest, which has an administrative office in Austin, Nevada. The Table Mountain Wilderness bounds the property on its northern and northwestern sides. The property lies entirely within the Elkhorn Canyon and Eagle Pass 7.5-minute U.S. Geological Survey quadrangles. Elevations on the property range from 2109m to 2696m (6920 ft to 8842 ft) much of which is mountainous terrain covered by pinyon, juniper, and mountain mahogany forest. A large burn area in the southern third of the property dates from around 2006-2007. A maintained road, NFD 324 bisects the property west-to-east from the Barley Ranch on the west to the abandoned Daugherty Ranch on the east and passes by the historical King Solomon antimony mine. Other roads include a few unmaintained ranch roads that access the range front area from the southeast, and an old drill road that extends northeast from the King Solomon mine to the informally named “Paleozoic Canyon” along the southern edge of the KS window; this road is the main access to the northern two-thirds of the property. In addition, the Willow Creek Road north of the Daugherty Ranch was used to access the extreme northeast corner of the property.

Claims - The property consists of 207 unpatented mining claims (Appendix A), some of which partly overlap. The claim block covers 16.13 km² (6.23 mi²). King Solomon is elongated northeast and is approximately centered on the historical King Solomon antimony mine and the King Solomon fault zone. A “tail” of the property is oriented due N-S and encompasses a narrower corridor centered on the southward strike extension of the King Solomon fault. Just south of the King Solomon mine is a bend in the strike of the King

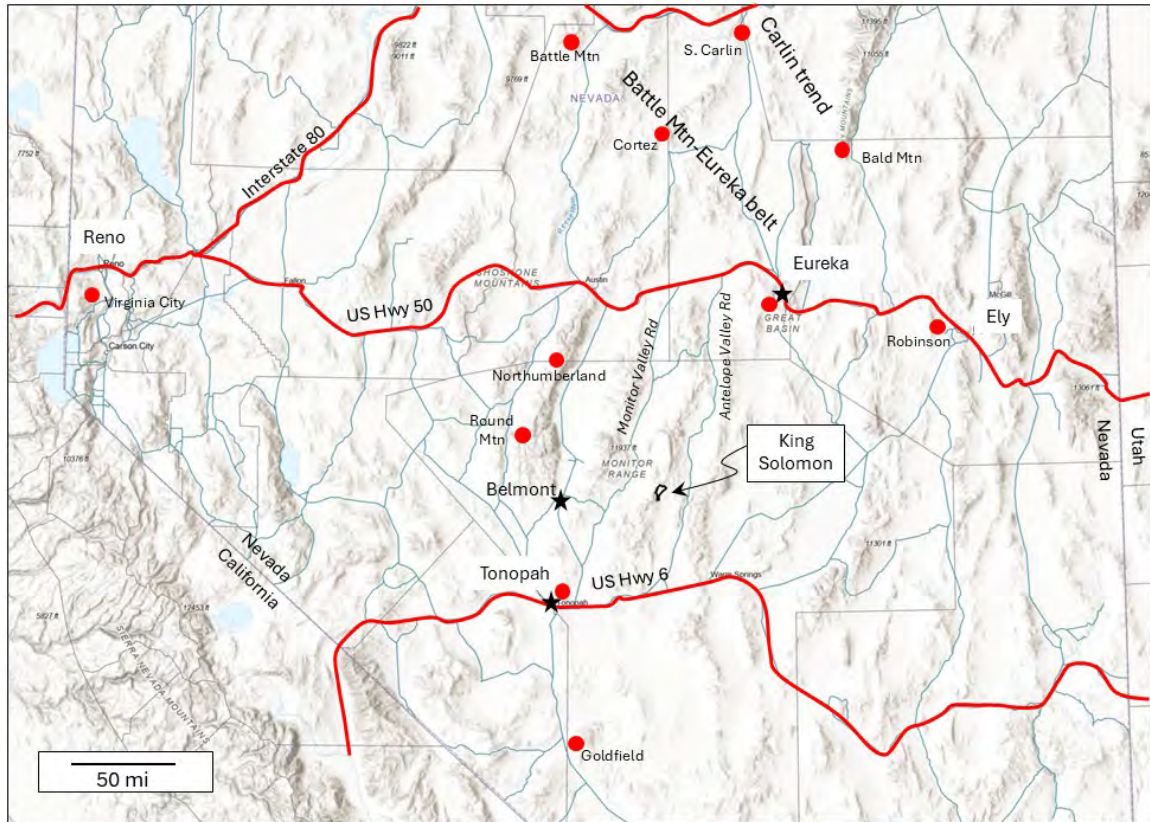


Figure 1. Location of the King Solomon project in central Nevada. Selected mines and districts are depicted with red dots.

Solomon fault from 050 degrees azimuth to 90 degrees azimuth, which accounts for the change in the property's orientation.

Access - The King Solomon project area is accessible by a combination of well-maintained, graded county roads and various dirt roads variably maintained by the U.S. Forest Service. The property is about 61 miles northeast of Tonopah by taking U.S. Highway 6 east of town, turning north onto State Highway 376, northeast onto the Belmont Road, through the town of Belmont, veering northeast onto the graded Monitor Valley Road before heading east on Forest Service roads NFD 005 and NFD 324 to the King Solomon mine.

An alternate route from Tonopah is to take U.S. Highway 6 for about 30 miles east of Tonopah then turning north at the graded Stone Cabin Valley-Willow Creek Road (NFD 404-405) about a mile east of Saulsbury Summit for another 35 miles, which tracks north along the west side of Stone Cabin Valley. At the ruins of the Daugherty Ranch, turn west onto the House Canyon Road that leads to the King Solomon mine.

Access from Eureka via Monitor Valley is about 108 miles. From Eureka, head west on U.S. Highway 50 for 25 miles, then take the graded Monitor Valley Road south for approximately 73 miles. Turn east at the junction with NFD 005, then near the entrance to the Barley Ranch, take NFD 324 east for about 10 miles up and over the crest of the Monitor Range to the King Solomon mine.

A more time-consuming route from Eureka is to head west on U.S. Highway 50 for 19 miles, then take the Antelope Valley Road south following the east side of the Monitor Range through Segura Pass and into Little Fish Lake Valley. From Fish Lake Valley, continue southward on NFD 139 (Hot Creek Road) to NFD 162. Head west several miles on NFD 162 until its junction with Willow Creek Road just south of the Daugherty Ranch.

Methodology

This report summarizes findings from mapping and sampling of the King Solomon exploration property during the period from June 19 through October 11, 2025. A total of 27 mapping days were spent at King Solomon to complete this property-wide geologic map. Mapping initially was done using 1:5000-scale mylar overlays on orthophoto bases and a handheld Garmin 60CX GPS to capture outcrops, lithologies, alteration assemblages, and structure but a decision was made early to improve and make more efficient the collection of field map data without compromising data quality by using a GPS-equipped Apple Ipad Pro tablet and Avenza Maps Pro mapping software. The digital format allowed for the upload of a wide range of user-defined detailed base maps (orthophotos, 7.5-minute topographic, digital elevation models) and supporting map information including geophysical and geochemical overlays, claims, and historical mapping and drilling information. Digital mapping also allowed for seamless zoom between high-resolution mapping (~100- to 1000-scale) where complexity demanded it to regional-scale surveying (>24,000-scale) and any scale in between. Detailed field notes were entered digitally into Avenza; data collection included lithologic descriptions, structural measurements, information on alteration and mineralization, sample cataloging and description, georeferenced photographs, the geometry and size information of outcrops or subcrops, and general notes and work-in-progress interpretations. Digital mapping with a water-resistant tablet proved effective during a particularly rainy field season. These data were downloaded into spreadsheets, organized, and then uploaded into ESRI ArcGIS Pro software each evening while in the field. Interpretations of geologic contacts (linework) were usually done every two weeks. A total of 1,515 map data points, 383 structural measurements, 143 geochemical samples, and 15 biostratigraphic age or geochronology samples were collected along with more than 100 library samples. The mapping resulted in



the identification of 22 distinct rock units and numerous faults, and it contributed to an emerging understanding of the area as a significant volcanic vent complex that imparted fundamental controls on the distribution of the observed Au-Ag-Sb mineralization. In addition to field activities, much of the legacy geological, geochemical, and geophysical databases were compiled, uploaded, and georeferenced in GIS to allow for rapid and efficient comparison and evaluation of spatial data during and after mapping was completed.

STRATIGRAPHY

The King Solomon property is underlain by extensive silicic caldera-related volcanic rocks of likely Oligocene age that were emplaced during the Great Basin's "Ignimbrite Flare-up." Although not isotopically dated at King Solomon, several isotopic ages on similar igneous rocks in the central Monitor Range indicate a span of silicic volcanic activity between ~29.8 Ma and 25.6 Ma. The Cenozoic volcanic rocks at King Solomon include what is interpreted to be a >1.5km-thick prism of moderately welded intracaldera ash-flow tuff that is overlain and crosscut by post-caldera rhyolite lavas and associated dikes and other intrusive bodies that were emplaced in and near the caldera's ring fracture faults. The combination of an intracaldera margin setting and abundant late rhyolite source vents and intrusions is the primary reason for the area's widespread hydrothermal alteration and Au-Ag-Sb mineralization. Furthermore, the early ash-flow tuffs rest on Paleozoic sedimentary rocks that include an upper-plate assemblage of highly deformed siliciclastic strata that was thrust over a lower-plate carbonate assemblage. This structural relationship of upper and lower thrust plates together with the age assignments and lithologies of Paleozoic strata are consistent with the Roberts Mountains thrust system in central Nevada, which is widely regarded as an important factor for the formation of many major Carlin-type gold deposits in north-central Nevada. King Solomon is somewhat unusual because alteration assemblages and alteration have characteristics of both volcanic-hosted epithermal Au-Ag and Carlin-type systems, which will be discussed further in the Alteration and Mineralization section.

Assemblages and individual units are grouped in the discussion below by relative age from youngest to oldest.

Quaternary Deposits

Alluvial fan, alluvial terrace, and over bank deposits (unit Qal) – The King Solomon property encompasses mountainous terrain that spans an elevation range between about 2743m to



2133m (9,000 ft and 7,000 ft), progressing from near the range's crest on the west to its eastern range front. Deeply incised canyons developed in the volcanic and sedimentary rocks drain to the southeast and east. In bedrock areas in canyons, small alluvial terraces and over bank deposits near springs developed but many stream bottoms expose bedrock; thus, small alluvial deposits in ephemeral streams cutting most bedrock areas of the property were not delineated in the mapping nor were areas with steep slopes containing thin talus or scree deposits. In contrast, extensive and coalesced boulder fan deposits (unit Qal) cover much of the range front to the east of stream mouths, coinciding with youthful fault scarps that mark the King Solomon fault zone. The property extends about 0.4km to 1.2km east-southeast of the range front. Within the range front zone, several tilted mesas and other areas of resistant bedrock project through alluvial fans. Near the heads of fans, levees up to 12m high channelize modern stream flows. Boulder size in fans is highly variable, with large boulders over 2m in diameter occurring near stream mouths.

Older alluvial fan deposits (unit Qoa) – Incised older alluvial fan deposits occur sporadically along the King Solomon fault zone. These older boulder fan deposits are thought to predate the latest significant movement on the King Solomon fault zone, thus preserving these deposits as perched zones in the immediate footwall of the fault.

Landslide deposits (unit Qls) – Hummocky, boulder strewn surfaces, commonly with upward flaring head scarps comprise landslide deposits. The deposits are relatively rare considering the steepness of slopes.

Cenozoic Rocks

The King Solomon property exposes more than 2km of tilted Cenozoic volcanic rocks. These volcanic rocks are a combination of intracaldera ash-flow tuff and younger rhyolite lavas and intrusions likely Oligocene in age based on isotopic ages on similar rocks in nearby locations of the Monitor Range. The hanging wall of the King Solomon fault contains unaltered outflow-facies glassy ash-flow tuff that may be substantially younger than the rocks of the King Solomon volcanic center.

Late Ash-flow Tuff and Bedded Tuffs (units Tbt2, Tbt, Trt, Tts2)

Three relatively thin but distinct cooling units of ash-flow tuff (map units Trt, Tbt, and Tbt2) overlie the House Canyon rhyolite (Thr) and Willow Creek rhyolite (Tpr) in the east-central and southern sectors of the King Solomon property. These well-exposed tuffs represent the youngest Neogene volcanic units on the property and are lithologically distinct from underlying rhyolite lava and intrusion-dominated units of the King Solomon center. In addition, the late tuffs, which rest upon mostly altered rocks of the King Solomon center,



lack hydrothermal alteration and are therefore considered to postdate mineralization. The ash-flow tuffs are separated from the underlying rhyolite lavas and subvolcanic intrusions by a thin, poorly exposed layer of fine-grained, white weathering tuff and tuffaceous siltstone (unit Tts2).

The lowermost of the three tuffs, called the red tuff (Trt) is a texturally variable, densely welded, dark reddish-brown weathering, locally vitrophyric ash-flow tuff with large (up to 10cm diameter) pale red pumice in its uppermost part. The red tuff is less evolved dacite and has characteristic dark green pyroxene phenocrysts in addition to hornblende and plagioclase phenocrysts. The tuff is sparsely phyric and pumiceous near its base. Thickness is approximately 25 to 35 meters. The red tuff forms a prominent tilted mesa about 1km due east of the King Solomon mine.

Unconformably overlying unit Trt is a tan weathering, fresh, welded, vitrophyric tuff (unit Tbt or lower biotite tuff) with abundant (>30%) compacted, black, glassy pumice to 0.5 cm diameter. This plagioclase-sanidine-biotite-quartz rhyolite ash-flow tuff is about 10m to 40 thick, its base generally porous, non-welded and poorly exposed. The unit breaks into sharp planar fragments.

The uppermost biotitic ash-flow tuff (unit Tbt2) is another fresh, glassy ash-flow tuff similar to Tbt but separated from it by a cooling break in outcrops east of the King Solomon mine, where the Tbt2 unit is approximately 30 m thick. The Tbt2 was not recognized as a separate cooling unit in exposures south of the King Solomon mine, and biotite rhyolite ash-flow tuff there is mapped only as unit Tbt.

Within the project area, these late ash-flow tuffs are exposed only in two small, fault-bound, tilted mesas that lay east of and in the hanging wall of the King Solomon fault zone.

Interpretation of late ash-flow tuffs

Ash-flow tuffs are not a significant component of the underlying rhyolite lavas of the King Solomon volcanic center. These younger and distinct ash-flow tuffs with sparse intervening bedded tuff (planar surge beds, tephra, and/or tuffaceous siltstone) are therefore considered separate from the King Solomon rhyolite center although they could be distal ash flow tuffs derived from other nearby caldera centers (e.g., southern Monitor and Toquima ranges). A further line of evidence differentiating these late tuffs from lavas of the King Solomon center is their lack of hydrothermal alteration whereas underlying rhyolitic rocks of the King Solomon center have been moderately to strongly affected by hydrothermal alteration.

King Solomon Volcanic Center (units Tfpr, Tts, Tpr, Thr, Tqpr, Tqprl, Tpbx)

A large part of the King Solomon property is underlain by locally derived silicic volcanic and subvolcanic rocks of probable Oligocene age that are considered part of the King Solomon volcanic center. In general, rhyolite lavas and their subvolcanic dikes, plugs, and stock-like bodies crop out moderately to very well. Outcrops tend to be large and weather prominently relative to underlying and generally more high altered early ash-flow tuffs. Lavas form extensive tabular outcrops whereas dikes and other intrusions have rib-like or plug-like exposures.

Rocks of the King Solomon center consist largely of quartz-bearing, flow-foliated to massive rhyolite lavas, and underlying, widespread and abundant silicic dikes, other subvolcanic intrusions, and intrusive breccia bodies with minor small-volume ash-flow tuff and related tephra and surge deposits (Figures 2, 3, and 4). Most of the lavas (units Thr, Tfpr, Tpr) are moderately porphyritic (15-25% phenocrysts) plagioclase > biotite > quartz > sanidine ± hornblende rhyolites; phenocryst size is mostly consistent with most rhyolites having phenocrysts between 1mm and 3mm diameter. A distinct sparsely and finely porphyritic, likely high-silica rhyolite (Tqpr) is less abundant and occurs mostly as dikes and stock-like bodies. Some quartz phenocrysts in the sparsely porphyritic variety are smoky.

An estimated 16 volcanic and subvolcanic rhyolitic source vents were mapped (Appendix E, Geologic Map). The vents were identified by tracing dikes to lava flows, by prominent, spine-like topographic highs of contorted flow-foliated rhyolite, and by following flow lobes upgradient. Most mapped vents occur in the south to southwest and northeast sectors of the property and coincide with the Stone Cabin and Willow Creek rhyolite lavas and domes (Figures 2, 3, 4). No vents were identified for the House Canyon rhyolite but it is likely it also vented in areas near the King Solomon mine. The vents coalesce with thinner and more tabular rhyolite lavas exhibiting many irregular pod-like to dike-like bodies of lithic-rich polymictic breccia containing abundant juvenile rhyolite clasts together with ubiquitous but varying amounts of other igneous and sedimentary clasts in a comminuted ashy matrix that most commonly was intensely hydrothermally altered. Polymictic breccia bodies are considered to derive from phreatomagmatic activity associated with shallow silicic vents commonly along the range front normal fault in the southern sector of the property and in a cluster cutting older ash-flow tuff in the west-central part of the property.

The total thickness of rhyolite lavas varies substantially owing to the viscous nature of the lavas during emplacement. Maximum thicknesses proximal to rhyolite domes exceed 240 meters (>800 feet). In the vicinity of the King Solomon mine, the average thickness of the House Canyon rhyolite (Figure 5), a subunit within the King Solomon volcanic center, is probably about 50 to 100 meters.



Figure 2. Stone Cabin rhyolite lavas and domes (Tfpr). Flow-layered, finely porphyritic plagioclase-biotite-quartz rhyolite. The yellow color is from modest jarosite and clay alteration that is characteristic of the unit.

Rhyolite lavas of the King Solomon center generally rest unconformably on a >1.5km-thick sequence of successive, relatively thin, compositionally similar ash-flow tuffs. At the King Solomon mine, two outcroppings of jasperoid derived from Paleozoic thin- to medium-bedded limestone underlie the House Canyon rhyolite with little or no intervening early ash-flow tuff. The jasperoids occur along the King Solomon fault zone and are interpreted to reflect a NE-trending horst block also defined by a narrow gravity high.

Subvolcanic intrusions of the King Solomon center are widespread and abundant (units Tqpr, Tpbx; Figures 3 and 4). The intrusions occur as small stock-like bodies and abundant dikes. Several dikes over 700 meters long occur with the longest at least 1.1 km. The dikes are invariably quartz-bearing rhyolite and possibly high-silica dacite composition based on their abundant phenocrystic quartz. Dikes commonly form resistant, rib-like outcrops and some are more than 100 m wide and have wall-like outcrops exceeding 20 meters height. These larger dikes are commonly composite, consisting of a complex mix of flow-foliated to

massive porphyritic rhyolite and polymictic breccia containing abundant juvenile rhyolite clasts as well as smaller and less abundant Paleozoic sedimentary clasts.



Figure 3. Variably altered Sone Cabin composite rhyolite dikes and stock-like intrusions (units Tqpr and Tpbx) cutting early ash-flow tuff. Some dikes are as much as 100m-thick.

Rocks of the King Solomon center are variably altered with intrusive units significantly more altered than volcanic flow units, an exception being the lavas and subvolcanic intrusions of the House rhyolite exposed along the central King Solomon fault zone from the historical King Solomon antimony mine northeast to the southern edge of the KS window, which are intensely altered. Tuffs that unconformably overlie the House rhyolite (Thr) in the hanging wall of the King Solomon fault zone are not altered and regarded as post-mineral (units Tbt2, Tbt, Trt, and Tts2). Large areas of rhyolite lavas in the footwall of the King Solomon fault zone, particularly south of the King Solomon mine, have been variably bleached and altered to an assemblage dominated by sericite (likely illite) after feldspar and biotite and secondary quartz.



Figure 4. Stone Cabin polymictic rhyolite intrusive breccia (unit Tpbx) associated with dikes. Clasts are highly variable in size and shape but invariably contain abundant juvenile rhyolite clasts. The matrix is comminuted rock flour, which is commonly altered to quartz, sericite, and sulfides where unoxidized. Polymictic intrusive breccia is considered potentially important for exploration because of its disseminated mineralization.



Figure 5. Strongly altered and mineralized, finely porphyritic House Canyon rhyolite from the King Solomon mine area. Dense, finely porphyritic plagioclase-biotite-quartz rhyolite is pervasively sericitized and moderately silicified. Jarosite (upper left), barite (upper right), stibnite (lower photographs) are common features of the Au-Ag-Sb mineralization.

Geochronology of volcanic rocks at King Solomon

Isotopic ages from ash-flow tuff in the vicinity of the central Monitor Range are sparse. The most precise ages are from $^{40}\text{Ar}/^{39}\text{Ar}$ dating of sanidine (29.87 ± 0.07 Ma, 1σ) and of hornblende (29.70 ± 0.12 Ma, 1σ) from the tuff of Haystack Canyon from Table Mountain about 12km north of King Solomon (Shawe et al., 2014). These ages are thought to best represent the timing of caldera collapse associated with the Table Mountain caldera. A less precise K-Ar biotite determination from welded silicic ash-flow tuff taken from the east flank of Table Mountain in the Danville mining district 14km north of King Solomon gave an age of 29.12 ± 0.7 Ma (Sloan et al., 2003; Marvin and Cole, 1978). A rhyolitic intrusion near the headwaters of Willow Creek ~4 km northeast of King Solomon gave a K-Ar biotite age of

29.3±0.9 Ma (McKee and John, 1987); this rhyolite is equivalent to the rhyolite lavas and subvolcanic intrusions in the northeast corner of the King Solomon property (unit Tfprf). Two K-Ar biotite ages were determined for samples collected west and southwest of King Solomon, one gave an age 25.57 ±1.50 Ma from a sample of rhyolite vitrophyre and was taken 5.5km southwest of the southern boundary of King Solomon (Marvin et al., 1973) and a second sample from a rhyolite intrusion near the head of Wattles Creek, ~2.5km west of King Solomon gave an age of 26.2±0.8 Ma; McKee et al., 1987). The ages suggest that tuff of the Table Mountain caldera (i.e., the tuff of Haystack Canyon) erupted at ~29.8 Ma and that considerably younger (~26-25 Ma) post-caldera rhyolites lavas and intrusions that are considered equivalent to the King Solomon center were emplaced around 25.6 Ma. For comparison, collapse of the nearby Round Mountain caldera occurred at ~26.5 Ma and major gold mineralization occurred at ~26.0 Ma (Henry et al., 1997).

Interpretation of the King Solomon rhyolite center

Rhyolitic rocks of the King Solomon volcanic center are interpreted to be the latest eruptive stage of caldera-related volcanism related to resurgence. The rhyolite lavas and their subvolcanic intrusions overlie a very thick (>1.5km) sequence of stacked rhyolitic to dacitic ash-flow tuff (see below: *Early Ash-flow Tuffs*) that are considered to represent the pyroclastic products of either multiple, smaller eruptions and protracted caldera collapse, or successive small-volume eruptions that followed caldera collapse. Many intracaldera tuff sequences consist of monotonous, successive, moderately welded ash-flow tuff (e.g., Best et al., 2013), and this interpretation is preferred for King Solomon based on the compositional similarities between the King Solomon rhyolites and underlying ash-flow tuffs. The King Solomon rhyolites were also part of a regional doming event that resulted in curvilinear faults and map patterns for units around a central Paleozoic dome-shaped “window” that is interpreted to reflect central resurgence in a caldera setting.

The central Great Basin in Nevada and Utah was a locus for intense caldera volcanism in the Eocene through Oligocene, part of the regional “ignimbrite flare-up” that also affected much of the southern U.S. Rockies and the Sierra Madre Occidental in Mexico. In the general area of King Solomon, several large, partly overlapping silicic calderas are inferred or known including the Round Mountain (26.5 Ma), Mount Jefferson (~27.3 Ma), and Lunar Lake (~25 Ma) calderas in the adjacent Toquima Range and the Table Mountain (~27.8 Ma) and Big Ten Peak (~27 Ma) calderas in the Monitor Range. One of the largest of these, the Table Mountain caldera, is more than 31km in diameter and its southeastern boundary coincides with the King Solomon property (Ludington et al., 1996); much of the Table Mountain caldera is inferred to underlie Monitor Valley. Findings from the present mapping

at King Solomon support the interpretation that the property lies along the southeastern ring fracture of the Table Mountain caldera.

Early Ash-flow Tuffs (unit Tcpt)

Silicic volcanic rocks of the King Solomon center rest unconformably on a thick sequence of hydrothermally altered rhyolitic to dacitic ash-flow tuff (Figure 6) Alteration has bleached the tuffs nearly everywhere to mottled white, yellow brown, and reddish brown. The unit is characterized by successive, relatively thin (~5m to 40m thick) pumice-rich, moderately to densely welded silicic tuff layers or individual cooling units, each with moderately abundant phenocrysts (12-20%) of plagioclase > biotite > quartz ± hornblende ± pyroxene phenocrysts and sparse (up to 5%) lithic fragments. There are likely scores of individual ash flows and cooling units represented in unit Tcpt. Phenocrysts vary in size but generally are between 1mm and 3 mm diameter although some quartz and sanidine can be larger. Large, compressed pumice up to or occasionally exceeding 10cm diameter in some ash-flow tuff suggests proximal sources for the sequence. The tuffs possess a characteristic anastomosing compaction or eutaxitic foliation defined largely by pumice. Welding in tuffs is usually moderate with aspect ratios of compressed pumice averaging between 1:3 and 1:5. Lithic fragments are ubiquitous to early ash-flow tuffs but they rarely exceed more than about 5% of the rock volume. Lithic fragments are subround to angular, commonly siliceous and fine grained and are thought to derive mostly from Paleozoic siliciclastic sedimentary rocks although some may be altered from limestone. Intercalated with ash-flow tuffs are minor thin-bedded tuffs that mark subtle unconformities between ash flows.

In mapping at King Solomon, efforts to subdivide the sequence were unsuccessful because of the similarity in composition and texture between ash-flow tuffs and the pervasive hydrothermal alteration affecting them. Thick, monotonous sequences of moderately welded ash-flow tuff characterize many intracaldera settings in the Great Basin (Best et al., 2013).

Interpretation of early ash-flow tuff

Stacked, relatively thin, mostly moderately welded ash-flow tuffs that collectively are more than 1.5km thick underlie compositionally similar quartz-biotite rhyolite lavas and domes at King Solomon. Commonly, thick sequences of compositionally similar ash-flow tuff are associated with major caldera collapse events. In some cases like the Caetano caldera in north-central Nevada, much of the intracaldera tuff is contained in two thick, densely welded tuffs with lesser moderately welded successive sheets on top (e.g., John et al., 2008). In other calderas, like those of the Indian Peak caldera complex, intracaldera tuffs are mostly successive, thin layers that accumulated to great thickness in the intracaldera

setting (Best et al., 2013). At King Solomon, the older succession of ash-flow tuffs is underlain by Paleozoic strata, not massive welded tuffs like the Caetano caldera. This implies there was no singular, catastrophic caldera collapse, but rather many smaller eruptions of similar composition spaced temporally apart to allow for individual cooling of many successive tuffs. Examples of calderas that consist of successive, moderately welded ash-flow tuff are the Big Ten caldera just south of King Solomon in the southern Monitor Range, and the Monotony Valley caldera in the Pancake Range (Best et al., 2013).

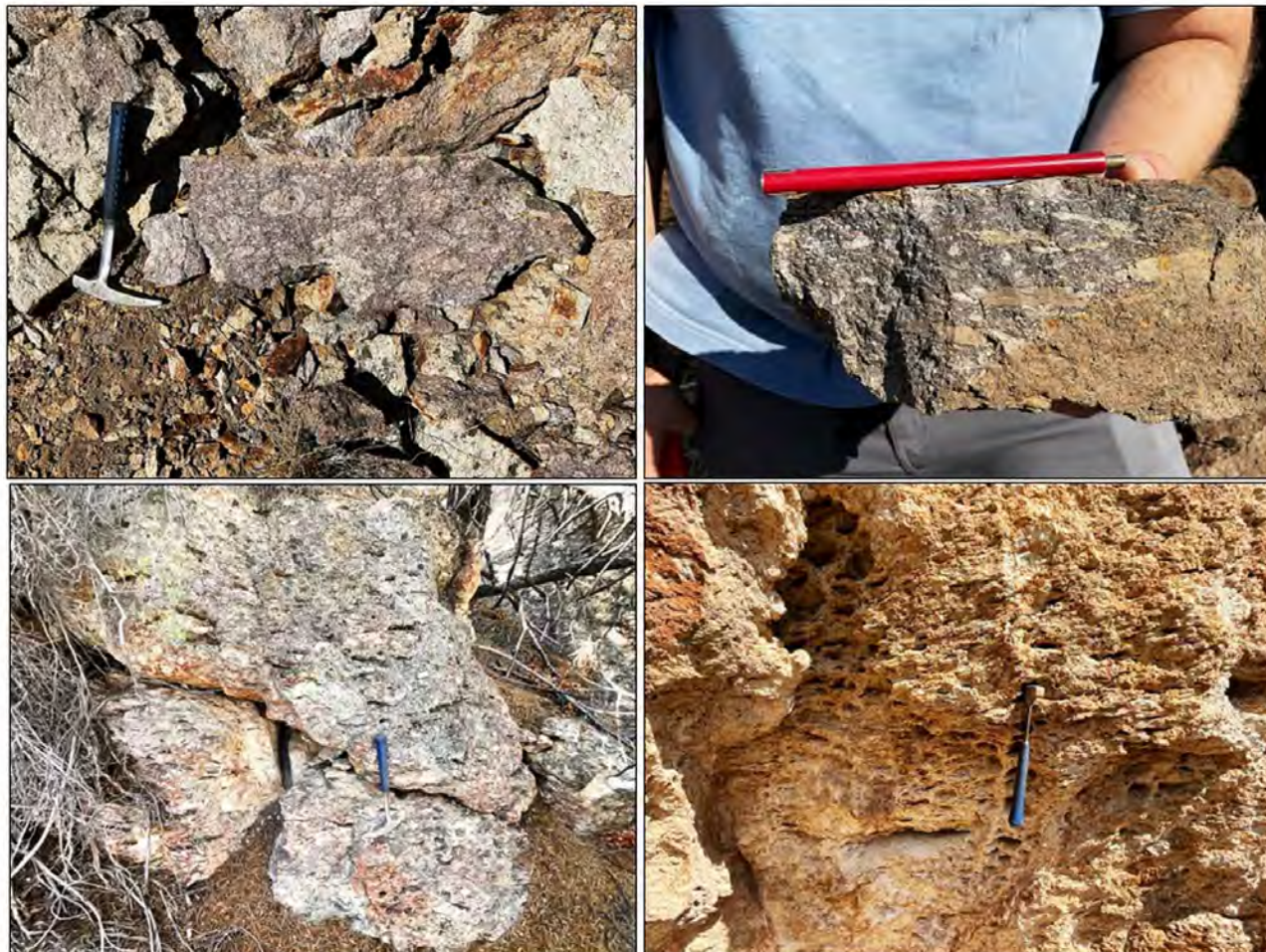


Figure 6. Images of moderately welded early ash-flow tuff (unit Tcpt). Tuff is typically coarsely and abundantly pumiceous, with eutaxitic pumice variable in composition and size. Largest pumice are >10cm diameter. Tuff contains 1-3mm dia. phenocrysts of plagioclase, biotite, quartz, sanidine, and hornblende.

The abundance of late vent-facies rhyolite lavas and intrusions that are compositionally similar to early ash-flow tuff is best explained by post-collapse effusive magmatism, which commonly occupies ring fractures and/or a central resurgent dome in a caldera setting. Similar occurrences of successive, thin ash-flow tuffs with compositionally similar late

lavas and domes are described in post-caldera collapse settings where doming or resurgence occurred (e.g., Lipman, 2006; San Juan volcanic field, Colorado).

Which caldera is King Solomon associated with?

The King Solomon property is inferred to lie astride the southeastern ring fracture of the Table Mountain caldera, a large (31km-diameter) silicic caldera that includes the large, tilted plateau comprising Table Mountain in the central Monitor Range. The intracaldera tuff of Haystack Canyon (~29.8 Ma) is considered to be at least one of the major intracaldera tuffs for it (Shawe et al., 2014). The character of the tuff of Haystack Canyon is poorly described other than being a silicic ash-flow tuff containing sanidine and hornblende. No information on the style of intracaldera tuffs (e.g., successive, moderately welded vs. massive, densely welded) is provided.

Based on previous work (e.g., Ludington et al., 1996; Best et al., 2013; Shawe et al., 2014) for the Table Mountain caldera, the interpretation that the Table Mountain caldera ring fracture traverses through or very near the King Solomon property is reasonable. This interpretation is further supported by new mapping that indicates the emplacement of voluminous rhyolite lavas and associated intrusions after the deposition of +1.5km of successive, moderately welded ash-flow tuff. The rhyolite lavas and related intrusions are considered to have been sourced along ring fracture faults and along younger normal faults like the King Solomon fault. As a note, however, the southeastern boundary of the Table Mountain caldera of Ludington et al. (1996) needs to be extended ~3.4km to the east to accommodate intracaldera tuff present at King Solomon.

A more speculative alternative is that King Solomon represents its own caldera, which is nested with the Table Mountain caldera farther to the west. This alternative is based on a few criteria: 1) the intracaldera tuff at King Solomon lies outside of the caldera boundary of Ludington et al. (1996), 2) the resurgent dome (i.e., KS Paleozoic window) is not centered on the Table Mountain caldera, and 3) curvilinear faults and unit map patterns at King Solomon are centered on the KS window, which defines a separate 14.5km x 13.2km caldera-like feature containing abundant early ash-flow tuff and late rhyolitic lavas and domes. At this stage of study, a simpler interpretation is the former, that King Solomon lies along the southeastern ring fracture of the Table Mountain caldera.

Paleozoic Sedimentary Rocks

The King Solomon property exposes a small (2.0 km²), circular structural window of “lower-plate” Ordovician through Devonian carbonate rocks (Appendix E: Geologic Map). The carbonates occur structurally beneath an outer rind of allochthonous, upper-plate strata comprised of chert, siliceous mudstone, carbonaceous mudstone, siltstone, and



sandstone. The two diverse rock assemblages are separated by a major thrust fault, which has intensely deformed upper-plate rocks and immediately subjacent lower-plate carbonates. The stratigraphic and structural relationships between lower and upper plates at King Solomon are consistent with relationships observed regionally for the Roberts Mountains thrust and its associated thrust faults. Subsequent Cenozoic domal uplift to expose the Paleozoic strata resulted in structural inversion such that the lower-plate carbonate rocks are exposed at elevations mostly surpassing the flanking upper-plate siliciclastic rocks. Subsequent normal faulting, principally along the King Solomon fault zone, exposes the deepest part of lower-plate Paleozoic section along the eastern margin of the window. Additional descriptions and interpretations related to doming are provided in the Structure section.

Upper-plate Sedimentary Rocks (undifferentiated; unit D-Osc)

A diverse group of clastic sedimentary rocks comprises the upper plate of a significant thrust exposed in the northern sector of the King Solomon property. Rocks include brick red sandstone and siltstone > black carbonaceous mudstone > siliceous or cherty mudstone and chert, with sparse tectonized lenses of lime mudstone. These siliciclastic rocks occur marginal to the carbonate rocks and invariably are separated from them by significant faults (Appendix E: Geologic Map; Stewart, 1980; Mathewson, 2002). None of the siliciclastic rocks crops out well; limestone and chert crop out best and their discontinuous and lensoidal shapes are evidence for the considerable contractional deformation that affected them. Bedding is rarely preserved in chert, but cleavage representing the intersection of bedding with a shear fabric is pervasive. The effect of this cleavage is a wavy or anastomosing penetrative foliation in most rocks.

Upper-plate strata were not subdivided based on lithologies because of considerable internal deformation, the rarity of outcrops, and mixed lithologies in areas of steep slopes and scree.

Red sandstone-siltstone with carbonate lenses

On the east side of the window, the structurally lowest unit that is in fault contact with lower-plate carbonate rocks is a red weathering, fine sandstone and siltstone with tectonized lenses of medium gray lime mudstone. Red sandstone-siltstone rarely crops out but is characterized by small blocky chips (up to 1cm dia.) of arkosic sandstone and fine tabular chips (up to 0.5cm dia.) of laminated siltstone. Where slopes are less steep, subcrop of red sandstone-siltstone is homogeneous and the soils are dark reddish brown to brick red.



Red sandstone-siltstone also comprises a sliver of float on steep slopes along the southwest shoulder of *Peak 8842* where it is structurally underlain by highly contorted and poorly exposed chert, siliceous mudstone, and black mudstone.

Black mudstone, siliceous mudstone, and chert

Mottled medium gray and brownish gray chert and siliceous mudstone due south of *Peak 8842* are in thrust contact with underlying lower-plate carbonate. The contact between chert and sandstone-siltstone is occupied by a large, northeast-striking composite Tqpr rhyolite dike, which effectively buttresses the peak. In contrast, a saddle just east of *Peak 8842* and areas to the northeast are underlain by black, carbonaceous mudstone that occurs only in subcrop float. It seems likely that this mudstone underlies the chert and siliceous mudstone present farther south although with structural disruption, this apparent stacking order may not persist laterally.

Lower-plate Sedimentary Rocks (lower-plate units, Dmls, Dm stms, Srm, Ohc, Oe, Op)

Several distinct lower-plate carbonate units and a quartzite were mapped. The strata are thought to range from Ordovician through early Devonian. Both the Silurian and early Devonian sections in the KS window are considered slope facies (Appendix E: Geologic Map). The strata from youngest to oldest include:

Massive, fossiliferous limestone (Early Devonian limestone Ddls) – Coils Creek member, McColley Canyon formation, or Rabbit Hill Limestone (?)

Small fault-bound slivers of massive ledge-forming limestone and dolomitic limestone occur on the eastern flank of *Peak 8842*. The bluish-gray weathering cliff bands are as much as 30 m high and consist of bioclast-rich grainstones that are commonly flat-pebble conglomerates. Fossil debris of crinoids, brachiopods, and bryozoans are abundant. Subordinate and thinner beds (<0.5m) of packstone and wackestone of nearly black lime muds are rarely interbedded with the massive grainstones (Figure 7).

Interpretation – Based on documented exposures at Lone Mountain, the Roberts Mountains, and in the Monitor Range at Table Mountain and Dobbin Summit (Murphy and Gronberg, 1970; Wise, 1977; Johnson et al., 1978), the massive grainstone exposed along the eastern flanks of *Peak 8842* may correlate with the Coils Creek member of the McColley Canyon formation or its down-slope equivalent Rabbit Hill Limestone. The abundance of grainstone and ledgy nature of this unit in the KS window suggests a more proximal slope depositional facies. Johnson et al. (1978) place the central Monitor Range in the transition between McColley and Rabbit Hill facies based on brachiopod fauna.

Thin- to medium-bedded limestone (Devonian limestone, unit Dd stms)

A monotonous and thick section (possibly structurally overthickened) of thin- to medium-bedded, dark gray, dark greenish gray, and yellowish gray weathering silty to muddy limestone characterizes much of the surface exposures on *Peak 8842* (Figure 7). The limestone is commonly partly or weakly dolomitic. Planar to wavy beds average between 2cm and 20cm thick and consist dominantly of wackestone > packstone. Fine interbeds of fine-grained sand and/or silt are present and are suggestive of allodapic deposition, perhaps associated with turbidites.

The thickness of unit Dd stms is difficult to ascertain because of significant deformation that affects it. Both tops and bottoms of the unit are faulted. The upper contact lies beneath the major thrust that separates upper and lower plates in the western margin of the KS window, whereas a high-angle fault juxtaposes the Roberts Mountains, Hanson Creek, and Eureka formations against it and unit Ddls along the easternmost margin of the window. In addition, a nearly pervasive sub-bedding parallel foliation affects this unit. The generally small angle (<30 degrees) between bedding and tectonic foliation imparts a distinctive hackly or anastomosing foliation in the limestone. The intensity of foliation varies but is most pronounced in structurally higher areas near the upper thrust that is interpreted to be the Roberts Mountains thrust. Small- to medium-scale folding and contractional deformation suggest that the observed thickness of unit Dd stms may be significantly more than its depositional thickness.

Planar thin-bedded limestone (Roberts Mountains Formation, unit Srm)

Planar thin bedded to laminated silty limestone (wackestone > packstone) forms a distinctive subcropping map unit (Figure 8) with an areal extent of 0.5km² in fault contact with the Hanson Creek unit in the far eastern margin of the KS window along the King Solomon fault zone. The unit occurs as angular, platy subcrop fragments and rare small outcrops that weather tan to pale red but are dark gray on fresh cut surfaces. The platy limestone contains occasional thin (~1mm) interbeds of clear quartz sand. Where more hematitic alteration occurs, thin, partly discontinuous, “pinstripes” of more intense hematite after pyrite is common. With greater alteration, the angularity of plates decreases and the edges become rounded. This platy limestone, which occurs in association with the Hanson Creek dolomite and Eureka quartzite, is consistent with an assignment to the Roberts Mountains Formation. Although exposure of the Roberts Mountains Formation is



Figure 7 (previous page). Devonian carbonate rocks of the KS Paleozoic window. Upper row: Cliff-forming coarse carbonate pebble conglomerate unit (McColley Canyon Fm., Coils Creek member(?); Second row left: Close-up of Devonian carbonate flat-pebble conglomerate. Second row right: White, planar thin-bedded limestone (wackestone). Third row: Fossil fragments in a thin bedded packstone; wavy interbeds in carbonate wackestone (silty limestone). Bottom row: Tight folds and shear fabric in Devonian thin-bedded limestone, possibly of the Bartine member of the McColley Canyon Formation.



Figure 8. Silurian Roberts Mountains Formation limestone. Top images show the typical exposure of platy silty limestone (wackestone) in subcrop with occasional medium beds. Middle left: medium bedded limestone with fine planar laminations with carbonate silt and fine sand. Middle right: packstone with ~10% fossil fragments. Bottom left: “Wispy” layering in silty limestone (wackestone). Bottom right: partly decarbonated, weakly carbonaceous and dolomitic silty limestone.

modest along the King Solomon fault, the unit is considered to be widespread beneath Devonian carbonates that comprise the bulk of exposures in the KS window.

Dolomite (Hanson Creek Dolomite; unit Ohc)

A massive, granular-textured, medium gray dolomite that weathers distinctively brown is present along the southeast sector of the King Solomon lower-plate window (Figure 9). Cliffy dolomite outcrops are as much as 20m high. Exposed structural thickness of the Hanson Creek Dolomite at King Solomon is about 50m. The dolomite locally preserves low-angle cross bedding but most bedding is massive and primary textures have been destroyed. Dolomite outcrops occur in a fault-bound wedge along the southern and eastern boundaries of the structural window, its eastern boundary being the NE-striking King Solomon fault zone and its southern boundary defined by the NW-striking Paleozoic Canyon fault. Considerable fault disruption occurs in the intersection zone and hydrothermal alteration is widespread. A few thin, mottled gray to white tectonized quartzite lenses occur in the upper part of the unit close to its contact with more massive quartzite of the Eureka Quartzite.

Quartzite (Eureka Quartzite ; unit Oe)

Two faulted ledges of dense orthoquartzite and quartzite breccia characterize the Eureka Quartzite at King Solomon. The quartzite ledges overlie dolomite of the Hanson Creek unit (Figure 9). The basal 10-15m thick ledge consists of mottled gray and white quartzite with considerable fracturing and small-scale faulting. The upper quartzite layer is heavily brecciated and consists of mottled white and gray quartzite with a pinkish brown tint. The upper quartzite may be as much as 25m thick, but like the dolomite of the Hanson Creek, has been extensively faulted. The map pattern of the quartzite is as an isolated circular body surrounded on three sides by the Hanson Creek dolomite with its northern boundary a fault separating quartzite from thin- to medium-bedded limestone and dolomitic limestone.

Thin- to medium-bedded limestone (Pogonip Group; unit Op)

Within strands of the King Solomon fault zone are slivers of poorly exposed thin-bedded limestone that is interpreted to underlie dolomite of the Hanson Creek. The float exposures on steep slopes are admixtures of platy gray limestone with considerable coarse blocky dolomite from above. No outcrops of the limestone were encountered. The Pogonip limestone is considered to be widespread beneath the Devonian carbonates and Ordovician Eureka Quartzite and Hanson Creek Dolomite that comprise most surface exposures in the KS window.



Figure 9. Top row : Typical outcrops of Eureka Quartzite. Two fault-bound ledges are present along the eastern margin of the KS window. Center left: Less altered Hanson Creek Dolomite. Center right and bottom row: Hydrothermally altered, red-brown weathering, vuggy Hanson Creek Dolomite along splays of the King Solomon fault zone in the eastern limit of the KS Window.

STRUCTURE

Late Normal Faults

The youngest faults in the King Solomon area form youthful scarps along the range front on the east side of West Stone Cabin Valley. These faults comprise the King Solomon fault zone, which in the south “tail” of the property strikes N-S and dips east in two subparallel strands separated by ~300-400m, but which, at the King Solomon mine, curves abruptly northeastward and splays into three distinct strands that appear to have lower total displacement northeastward. The King Solomon fault and its equivalent faults to the south of the King Solomon property bound the Monitor Range on its eastern side and have a cumulative strike length of more than 50km. Mineralization at King Solomon and south at Longstreet occurs along the fault zone but it is unclear if mineralization is directly related to early movement on the fault. A possibility is that the fault’s footwall exposes deeper and thereby more intensely altered parts of the King Solomon center potentially giving the false impression that the fault was a conduit for hydrothermal fluids. If the King Solomon fault was active during development of the King Solomon rhyolite center, then its activity was episodic over at least 20 million years. However, faults subparallel to the King Solomon fault control stibnite veins at the King Solomon mine, indicating that similar faults were open during mineralization (Lawrence, 1963).

Faults and Folds in the King Solomon Window

Two main fault sets were observed cutting Paleozoic strata in the King Solomon window. The youngest set are high-angle normal faults that are northward projections of the King Solomon fault zone where it splays into at least three major strands that cut both Cenozoic volcanic rocks and Paleozoic upper- and lower-plate rocks. Bedding replacement jasperoids are present in and near these normal fault strands of the King Solomon fault with jasperoid bodies commonly possessing elongation subparallel to faults (Appendix E. Structure Map and Geologic Map).

An earlier set of northwest-striking faults is cross cut by northerly striking normal faults associated with the King Solomon fault. On average, these high-angle faults have azimuth strikes of between 315 and 325. A prominent fault that coincides with the informally named “Paleozoic Canyon” localized jasperoid, particular near intersections with crosscutting northerly striking normal faults. In addition, an inferred northwest-striking fault bounds Paleozoic strata from early ash-flow tuff on the southern margin of the KS window.

There are abundant small-scale (1-5m-scale $\frac{1}{2}$ wavelength) folds in lower-plate carbonates of the KS window. The folds are best observed in the Dd stms unit because it is the most extensive Paleozoic unit and contains moderately good outcrop exposures. The folds

typically trend similarly to northwest-striking faults. Larger folds are generally difficult to trace laterally in part because of masking by the penetrative foliation. One large anticline traces north-northwest for more than 1km in the northeast part of the window.

Roberts Mountains Thrust

A major thrust separates siliciclastic sedimentary rocks from underlying carbonate-dominant sedimentary rocks in the KS window. The mapped thrust wraps partially around the carbonate window with inferred outward dips that support a dome geometry for the window (Appendix E: Structure Map and Geologic Map; Mathewson, 2001). The fault rarely crops out but is evident from the contrasting sedimentary lithologies of upper and lower plates. In addition, a considerable amount of carbonate veining (>30% or more) occurs in subjacent carbonate rocks together with more intense shearing evident by increased foliation development. The shear foliation affects a more than 200m thickness of the Dd stms unit becoming less pervasive at depth beneath the thrust.

Similarly, upper-plate siliciclastic rocks are intensely deformed, often with outcrop-scale kink folding and fault dislocations. Silicification in the form of quartz druses, local lenses of intense organic carbon alteration, and fracture surfaces contains yellow-green iron and arsenic oxides are common in proximal upper-plate lithologies. The upper-plate rocks are unconformably overlain by the early ash-flow tuffs, indicating considerable erosion prior to deposition of the volcanic rocks. Therefore, little information on the cumulative thickness of upper-plate rocks in the area is available.

Interpretation – The major thrust fault exposed in the KS window is interpreted to be part of the Roberts Mountains thrust system. This interpretation is consistent with regional stratigraphic and structural relationships for the thrust including distinct upper- and lower-plate stratigraphic units (e.g., Wise, 2010). The Roberts Mountains allochthon is mapped in the Simpson Park Mountains, Roberts Mountains, and only in small areas of the Monitor Range including western Table Mountain (Johnson and Gronberg, 1978) and Dobbin Summit-Clear Creek (Wise, 1977). This mapping at King Solomon documents the distribution and lithologies associated with the Roberts Mountains allochthon considerably farther south of previous locations. It is consistent with the inferred location of the Roberts Mountains thrust by Stewart and Poole (1974) and Stewart (1980).

Tilts on Volcanic Units

Volcanic rocks at King Solomon have variable foliation dip directions. Despite the variability, some generalities emerge from analyzing the strikes and dips of foliation in various Cenozoic units. Footwall to the King Solomon fault zone, the majority of volcanic strata have foliations that dip southward (both southeast and southwest). This overall trend



is supported stratigraphically by the southward younging observed in volcanic rocks from the thick succession of early ash-flow tuff (unit Tcpt) on the north to the rhyolite lavas of the King Solomon center (i.e., Stone Cabin lavas and domes) to the south. In contrast, volcanic rocks in the hanging wall of the King Solomon fault zone dip mostly eastward except for a tilted mesa in the bend of the King Solomon fault south of the King Solomon mine, which dips westerly.

Dome Structure

In the KS window, Paleozoic strata mostly have bedding dips that are east-northeast in the northeastern sector and east-southeast in the southwestern sector (Figure 10). This is consistent with bedding generally dipping outward from the center of the dome. Fewer bedding measurements on the western side of the window due to poor exposure make it more difficult to access bedding geometry there.

Doming of the entire area in and surrounding King Solomon is considered the principal reason for exposing both lower- and upper-plate Paleozoic sedimentary rocks through a thick cover of early Cenozoic volcanics (Mathewson, 2002). In addition to the indication of domed Paleozoic strata based on topographic expression and bedding orientations, older ash-flow tuff inside the King Solomon property occurs in a semi-circular pattern immediately south of the Paleozoic window, and the relatively young King Solomon rhyolite flows lay farther outboard still. This pattern of southward younging is considered to reflect regional resurgence centered near the KS window. Later truncation of the southeast sector of the dome by the King Solomon normal fault occurred, down-dropping younger volcanic rocks of the King Solomon center against Paleozoic strata and older ash-flow tuff. Several vents of the House rhyolite are localized near the King Solomon fault zone, suggesting that the fault may have been active in the Oligocene.

Based on knowledge gained from mapping, the area north of the window and property probably has opposite regional dips (i.e., northward dips) that reflect the north flank of resurgence.

The cause of the doming is not certain but a few important lines of evidence points to injection of magmas during caldera resurgence. The older ash-flow tuffs and Paleozoic strata of the KS window contain a swarm of porphyritic rhyolite dikes, plugs, and polymictic subvolcanic breccia bodies that indicate larger rhyolitic source plutons at depth. High viscosities in near-surface settings limit the lateral travel of most rhyolitic melts to very short distances (≤ 3 km) although vertical distances can be greater due to buoyancy forces. Thus, the pattern of diking at King Solomon outlines the breadth of underlying intrusive

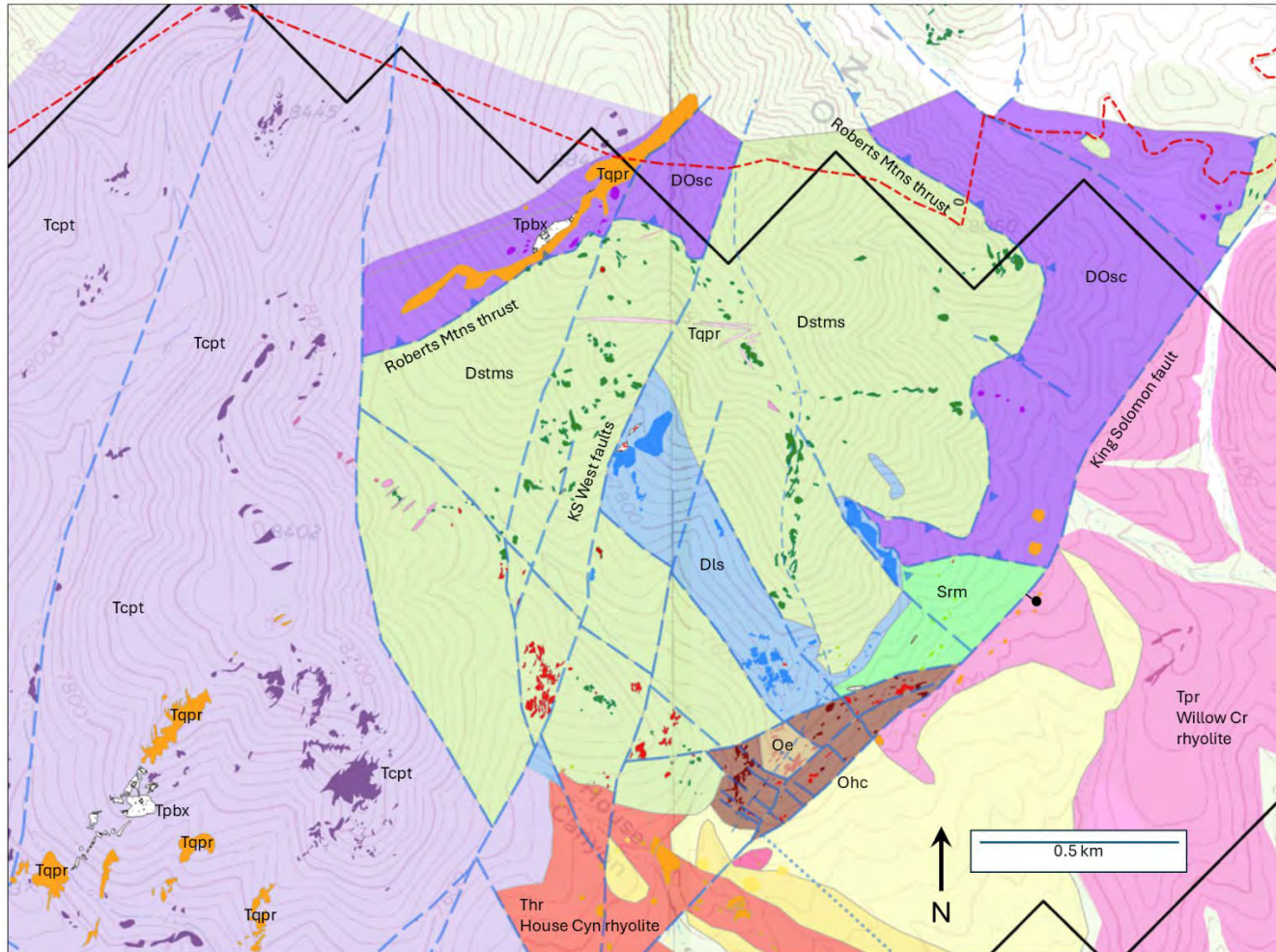


Figure 10. KS Paleozoic window. The window is the core of a broader domal uplift, the east part of which is cut by and down-dropped along the King Solomon fault. Units are: Tpr (Willow Creek rhyolite; fresh), Thr (House Canyon rhyolite), Tpbx (polymictic intrusive breccia), Tqpr (Stone Cabin porphyritic rhyolite dikes), DOsc (Devonian-Ordovician upper-plate siliciclastic rocks), Dstms (Early Devonian McColley Canyon Fm – Bartine member), Dls (Early Devonian McColley Cyn Fm – Coils Creek member), Srm (Roberts Mountains Fm.), Ohc (Hanson Creek Dolomite), and Oe (Eureka Quartzite). Black line is the King Solomon property line. Dashed red line is the approximate boundary of the Table Mountain Wilderness.

bodies although it is clear that the inferred bodies extend well beyond the property boundary. Other supporting data for doming by intrusion is from geophysical data; the gravity data show a pronounced gravity high centered over the KS window but extending into areas underlain by tuff and rhyolite lava in the footwall of faults like the King Solomon fault. The gravity high over the KS window is characterized by a corresponding and very prominent magnetic high as well. A magnetic high in areas of extensive Paleozoic carbonate rocks with low magnetic susceptibility infers a large and relatively magnetic intrusive body at depth, thus supporting the mapping that shows abundant diking.

Structure of Dikes, Plugs, and Breccia Bodies

Dikes and other small intrusions are abundant at King Solomon. The intrusions are correlated with the King Solomon volcanic center and cut caldera-related early ash-flow tuff. In some cases, intrusive bodies are interpreted in the lowest exposed parts of the King Solomon lavas, particular near the base of flows and exposed by faulting. Dikes and other bodies also occur widely in the KS window.

Most dikes at King Solomon strike northeast with azimuths ranging from 020 to 055 degrees. A few dikes strike northwest. The dikes are subparallel with the King Solomon extensional fault zone, which implies that the fault may have a long history of activity including early movement that partly controlled the distribution of Oligocene dikes. In addition, recent Quaternary displacement is evident in the southern arm of the fault. A dominant northeast orientation for Oligocene dikes in the King Solomon volcanic center implies a NW-SE-directed extension direction (σ_3).

ALTERATION AND MINERALIZATION

Extensive hydrothermal alteration and significant mineralization affects both Cenozoic volcanic and Paleozoic sedimentary rocks of the King Solomon property. The property is a rare example of co-spatial and coeval Carlin-type and volcanic-hosted epithermal Au-Ag-Sb mineralization, which therefore provides important constraints on the processes that contributed to both styles of mineralization. Higher silver contents in the Carlin-type mineralization is characteristic of shallower settings as described by Ressel et al. (2015).

Hydrothermal Alteration

Hydrothermal alteration affects rocks of the King Solomon rhyolite center, which are among the youngest rocks found on the property (Appendix E: Geology Maps, see Alteration Maps). The alteration in the King Solomon rhyolites mostly affects intrusive rocks, although the House rhyolite (unit Thr) is also moderately to strongly altered. Rhyolite lavas and domes of the Stone Cabin and Willow Creek vent areas are generally less altered, although more intense alteration occurs in associated dikes and breccia bodies in these areas. Late, small-volume glassy ash-flow tuffs in the hanging wall of the King Solomon fault are not hydrothermally altered, indicating that these units postdate alteration and mineralization. Early ash-flow tuff and Paleozoic carbonates in the KS window are variably hydrothermally altered. The early ash-flow tuffs are almost entirely moderately altered throughout their extent inside the property, a function perhaps of their intracaldera depositional setting.

Willow Creek domes and flows – This vent area in the northeast corner of the property contains very little alteration. The unit lies entirely on the hanging wall side of the King Solomon fault, including several vents for rhyolite domes (unit Tpr). Paleozoic strata in the footwall of the fault adjacent to the rhyolite are variously altered to jasperoid and locally affected by decarbonatization of carbonate, including sanding of dolomite. The relationships of alteration suggest but are not conclusive evidence that the Willow Creek lavas postdate alteration and mineralization.

The Stone Cabin rhyolite domes and lavas – This significant vent area in the southern part of the property bears more extensive hydrothermal alteration than those of the Willow Creek area. Alteration is common but generally weak in the lavas including where lavas vented in domes (units Tfpr, Tqprl), consisting of partial conversion of feldspar to clays (likely illite > smectite) and development of weak jarosite staining after partial alteration and oxidation of biotite. The weak alteration resulted in a mild bleaching of the rocks to a mottled white and yellow. Rarely in the lavas are zones of more intense alteration that include secondary quartz, pervasive sericitization (illitization), and moderate iron oxides after pyrite.

In contrast to the Stone Cabin lavas, underlying intrusive rocks related to Stone Cabin magmatism, including dikes, small plugs, and intrusive breccia bodies (units Tqpr, Tpbx), are commonly moderately to strongly altered. These altered and often mineralized intrusions occur abundantly along the N-S range front in the property's south end and in the southwestern sector of the main body of claims west of the King Solomon mine where intrusions cut the early ash-flow tuffs. The alteration in intrusive units is variable but commonly consists of a white bleaching of the rocks, and pervasive sericite replacement of feldspar and biotite. Silicification ranges from mild to pervasive, and where more

intense, fine chalcedonic veinlets appear often containing fine sulfides. Iron oxide contents are commonly weak to moderate, with relict pyrite and/or marcasite present as fine disseminations in the most intensely silicified rocks.

A particular kind of rock type in the Stone Cabin center associated with dikes and plugs is a polymictic breccia (unit Tpbx) containing abundant juvenile rhyolite clasts together with varying amounts of angular to round Paleozoic sedimentary fragments in a fine-grained, crystal-rich comminuted igneous groundmass; these polymictic intrusive breccias are considered phreatomagmatic, and their groundmasses are typically moderately to strongly silicified with abundant very fine-grained sulfides where less oxidized.

House Canyon rhyolite lavas and intrusions – The House rhyolite (unit Thr) underlies the King Solomon mine area mostly between strands of the King Solomon fault zone but also present in the hanging wall of the fault. The unit extends northward along the fault zone to the KS window, but its southern extent is limited to the King Solomon mine. Most of the unit consists of massive to flow-foliated porphyritic rhyolite lava, but dikes and small intrusions occur in the area of the King Solomon mine where they crosscut older units. Moderate to strong sericitization (illitization) from nearly pervasive feldspar and biotite destruction occurs in rhyolite over a broad area around the mine. The sericite is accompanied by quartz addition through moderate silicification of the groundmass and sparse quartz-chalcedony veining containing coarse stibnite and tabular barite. Locally, at the King Solomon mine proper, nearly pure blue-gray veins of stibnite to a few centimeters thickness are present cutting the House rhyolite. The prominent vein set is sub-vertical and strike ~350° azimuth. Polymictic intrusive breccia, similar but more silicified and pyritic than observed breccia from the Stone Cabin dike area, occur scattered throughout the area of the King Solomon mine mainly as float boulders to 0.5m diameter.

Au-Ag-Sb Mineralization

Mineralization exposed on surface at King Solomon appears relatively simple and expressed similarly in both volcanic and sedimentary rocks. As expected, most mineralization occurs in zones with the most intense hydrothermal alteration, which in mapping included areas having complete feldspar and biotite destruction in rhyolitic rocks and moderate to pervasive silicification in both carbonate and volcanic rocks. A caveat to the latter in carbonates is a difficulty in mapping exposures of decalcified and/or partially dolomitized rocks that weather recessively peripheral to areas of jasperoid; these rocks are underestimated in their extent. In both quartz-sericite altered volcanic rocks and in jasperoid after carbonate rocks, local zones with late-stage irregular veins and crackle

breccia contain drusy to comb quartz, and coarse stibnite, barite, ± calcite, commonly in association with open space. High silver values in vein-rich samples suggest that silver is contained in stibnite. The stibnite-bearing veins and veinlets are paragenetically late whereas zones of fine-grained chalcedonic quartz in both jasperoid and silicified rhyolite contain relict, extremely fine-grained iron sulfides thought to predate stibnite-bearing veins and open space development. The poor correlation between gold and late stibnite mineralization supports the hypothesis that gold was introduced early, probably in these fine iron sulfides. Little or no base metal sulfides were observed.

Post-mineral oxidation of sulfides at King Solomon is generally considered weak. Sulfides, although scarce, are present in numerous surface exposures in both volcanic and carbonate rocks, particularly those in the footwall of the King Salomon fault zone. It is possible that deep oxidation that could affect precious metal recovery has occurred in the fault zone and perhaps its hanging wall. An evaluation of drilling that may address oxidation at depth has not been done as part of this study.

Styles of Mineralization

King Solomon is unusual in possessing intermingled epithermal and Carlin-type styles of mineralization (Figure 11). In this case, the styles of mineralization are a function of differing host rock rheology and reactivity to what was almost certainly the same ore fluids. As such, the two types of mineralization share geochemical expressions and sulfide/sulfate mineralogy, but they differ widely in the nature and style of hydrothermal alteration, the volcanic-hosted alteration resulting in variable degrees of sericitization due to the relative impermeability of the feldspathic host. Partial silicification occurs through replacement of volcanic groundmasses, particularly in areas with quartz veining. In contrast, moderately acidic ore fluids reacted more broadly with and partly dissolved calcite in limestone, producing very fine-grained bedding replacement type jasperoid in some areas, and in other areas either leaving residual dolomite at the expense of calcite dissolution or depositing hydrothermal dolomite. The reactivity of the rock effectively creates its own porosity and permeability, given enough ore fluid to combat the buffering effects of the carbonate. The product of carbonate dissolution is a lower-density, micro-porous and relatively soft rock that typically has enriched organic carbon due to volume loss and weathers recessively. Such decalcified rocks crop out poorly if at all, necessitating detailed mapping and sampling in areas of carbonate float in future exploration endeavors. The contrast in alteration and mineralization processes among volcanic versus carbonate rocks results in a dominance of vein or veinlet-hosted styles for volcanic hosts whereas the dissolution process in carbonates results in a bulk-disseminated type of mineralization

that is a hallmark of large Carlin-type systems. One exception to the dominant vein control on mineralization in volcanic rocks is the pervasive style that accompanies polymictic intrusive breccia bodies. High porosities and permeabilities in these breccias led to pervasive alteration in groundmasses.



Figure 11. Representative types of mineralization observed at King Solomon. Top left: Decarbonized, porous, carbonaceous silty limestone in the Roberts Mountains Formation (Zn = 1690 ppm); Top right: jasperoid with relict bedding textures in Devonian McColley Canyon Limestone(?) (3.9 g/t Au, 63 g/t Ag); Middle left: jasperoid with late stibnite-stibiconite in voids and fractures, Devonian McColley Canyon (26 ppb Au, 39 g/t Ag, 38 ppm As, >1% Sb, 28 ppm Se); Middle right: Jasperoid of the Roberts Mountains Formation with late-stage white, drusy quartz infilling irregular vugs (165 ppb Au, 266 g/t Ag). Bottom left: Silicified and sericitized House porphyritic rhyolite with late-stage stibnite-quartz-barite in irregular vug infilling (2.82 g/t Au, 609 g/t Ag, >1% Sb); Bottom right: Polymictic rhyolite intrusive breccia from King Solomon antimony mine with fine-grained stibnite (389 ppb Au, 278 g/t Ag, >1% Sb).

Timing of Alteration and Mineralization

Hydrothermal alteration and mineralization at King Solomon affect nearly all Cenozoic and older rocks except for the youngest ash-flow tuffs (units Trt, Tbt, Tbt2) present in the hanging wall of the King Solomon fault zone, which are considered post-mineral. In addition, some rhyolitic lavas (e.g., Willow Creek and Stone Cabin) range from essentially fresh to only weakly altered. The Willow Creek and Stone Cabin lavas may reflect a post- or late-mineral timing, respectively.

GEOCHEMISTRY

Smaller Au, Ag, As, and Sb anomalies in soils are associated with shallow trenches in the King Solomon mine area and in an area to the south coincident with the major bend in the King Solomon fault, called the Big Bend area.

Rock-chip samples – Data from 213 rock-chip samples that have reliable tabulated locations were compiled, of which 143 samples were from sampling during this mapping (Appendix C: Rock-chip and Soil Geochemistry). Rock chips data corroborates the soils data and mapping for identifying surface expressions of mineralization, namely in the KS window, the King Solomon mine, Big Bend, and an area called West Ridge on the west-central side of the property. The KS window and King Solomon mine have received the bulk of exploration and drilling, and fewer than 6 holes were drilled in each of West Ridge and Big Bend. All four areas contain significant gold in rock chips >1 ppm along with elevated Ag, As, Sb, Hg, and Tl. In addition, numerous other areas associated with a swarm of dikes and other intrusions of the Stone Cabin center are associated with moderate to strong hydrothermal alteration share coincident elevated Au, Ag, As, Sb, Hg, and/or Tl. These areas are identified as priority 2 targets on.

GEOPHYSICS

Magnetics

Several types of geophysical surveys have been conducted over the King Solomon property. Kennecott contracted with High-Sense Geophysics Limited of Toronto, Canada for an airborne magnetics survey in 1995 that had 200m spacing on flight lines oriented northwest-southeast (High-Sense, 1995). The magnetometer position was 50m above the ground surface. The survey is extremely interesting in that it delineates a prominent magnetic high (e.g., reduced-to-pole data; Appendix F: Geophysics, see Magnetics RTP Map) centered over the KS window, which mapping indicates is underlain by essentially non-magnetic Paleozoic carbonate rocks. However, several mapped dikes including one major northeast-striking dike on Peak 8842 indicate the likelihood of a significant rhyolitic intrusion underpinning the window and helping to explain the regional doming evident from

the geology. Furthermore, the magnetics data from farther south show a pronounced magnetic low in an area that coincides well with the mapped Stone Cabin rhyolite center. Although the rhyolite lavas at Stone Cabin are weakly altered, the rock is magnetic and the magnetic low is almost certain related to reversed polarity (e.g., remanent magnetism) of this volcanic unit relative to the Willow Creek rhyolite lavas, which are strongly magnetic with normal polarity. The magnetics clearly map the King Solomon fault zone and less altered volcanics in the hanging wall of the fault have the characteristic alternation between normally and remanently magnetized units.

Gravity

Gravity data have been integrated for four separate surveys by Wright Geophysics of Spring Creek, Nevada in 2017, and again in 2018 (Magee Geophysics, 2017; Wright Geophysics, 2017). The data cover an area larger than the size of the property (Appendix F: Geophysics, see Gravity Maps). All of the data were collected by Magee Geophysics of Reno, Nevada except for some control points obtained from U.S. Geological Survey data. The King Solomon surveys were from Golden Chalice's work in 2006 (595 points), a survey by Orla Mining in 2017 (767 points), and another add-on survey in 2018 (440 points). In all, a total of 1,802 gravity data points were used to generate a series of maps depicting gravity models.

Gravity data for King Solomon delineate equally remarkable features as do the magnetics data. The data show a robust gravity high centered over the KS carbonate window as would be expected for relatively denser carbonates, and the gravity high corresponds very well with the magnetics high over the carbonate window, which as discussed in the previous subsection, is not expected for carbonate rocks and implies the KS window is underlain by a significant rhyolite intrusion at depth. The data also show a strip of rocks with higher gravity response along the immediate footwall of the King Solomon fault in areas underlain by relatively less dense volcanic rocks. The data therefore suggest that a zone along the fault's footwall contains carbonate rocks at relatively shallow depths. Other areas, particularly to the southwest of the KS window and underlain by the early ash-flow tuff also delineate gravity highs in topographically low areas along incised canyon. While no carbonate rocks were mapped in these areas, the data suggest that carbonates may lie at shallow depths.

The gravity data together with geological mapping are the best tools available for delineating faults at King Solomon. The King Solomon fault is well depicted in the gravity data (e.g., horizontal gradient map) and supports map data for more than one fault strand.



The horizontal gravity data also corroborates the extent of the Roberts Mountains thrust from mapping.

CSAMT and IP Resistivity and IP Chargeability

Cougar Gold LLC commissioned 6 northwest-to-southeast lines of controlled-source audio-magnetotelluric data (CSAMT) to be collected over the main area of King Solomon including the southern half of the KS window to just south of the King Solomon mine in 2006 (Zonge Geosciences Inc, 2006). An inversion of the data was later made by Jim Wright of Wright Geophysics in Spring Creek, Nevada, and depth slices of resistivity were generated from surface to 400m depth (Appendix F: Geophysics, see CSAMT Depth Slice Maps). The 200m to 400m depth slice, in particular, depicts higher resistivity (and potentially shallow carbonate) in an area underlain by early ash-flow tuff a short distance from the southwest margin of the exposed KS window. The same depth slice shows an elongate “tail” of high resistivity rock trending southward from the KS window along the King Solomon fault, suggesting fault-bound carbonates within relatively shallow depths.

Three northwest-to-southeast lines of IP chargeability and resistivity were completed by Cougar Gold LLC in 2006 for the area near the King Solomon mine (Appendix F: Geophysics, see IP Chargeability and Resistivity Maps). The data broadly show a sizeable tabular chargeable volume in the footwall of the King Solomon fault zone. The tabular zone coincides with a topographic bench that is interpreted to lie between two strands of the fault. The chargeability is inferred to reflect disseminated sulfides at fairly shallow depth of less than 50m but extending to considerable depth. Both IP chargeability and resistivity depict a sharp gradient in both chargeability and resistivity gradients that is interpreted to be the easternmost strand of the King Solomon fault.

Aster Hyperspectral Data

Only a single map was located depicting processed Aster short-wave infrared hyperspectral imagery (SWIR) over the King Solomon property (Appendix F: Geophysics, see Hyperspectral Map). The data were processed for Cougar Gold LLC in 2007. The composite log residual data show a nearly pervasive area of pink hues over the volcanic rocks at King Solomon that were broadly interpreted to reflect clay alteration. The pink hues are most intense in two concentric bands that lay outboard of the KS window. The King Solomon mine area, the West Ridge target area, and the broad area underlain by altered early ash-flow tuffs are interpreted as the strongest areas of clay alteration. Weaker clay alteration occurs in the area of the Stone Cabin rhyolite domes and lavas, although



intensity increases nearer the King Solomon fault. The Willow Creek lavas exhibit a green color that apparently indicates unaltered volcanic rocks. Carbonate rocks of the KS window also lack significant pinkish hues indicative of clay. Hyperspectral data correspond well with different intensities of clay alteration identified from mapping.

DRILLING

Data for 103 holes are recorded from the King Solomon project (Appendix B. Drill Hole Map). The holes were drilled by 8 companies spanning the period from 1981 through 1996. A majority of holes were drilled from 1984 through 1991, with only 6 holes drilled before or after. The total drilled footage is 41,340 ft (12,600m). All but 6 holes were drilled in the King Solomon fault corridor between the KS window and King Solomon, including a small area just south of the King Solomon mine. The 6 outlier holes were drilled in the West Ridge target just north of NFD 324 access road at the western boundary of the property.

This report does not include a thorough assessment of historical drilling at King Solomon. However, it is important to note that several holes in the King Solomon mine area and the southernmost margin of the KS window encountered interesting gold grades. Most of the holes were relatively shallow (i.e., <500 ft). Finally, no drilling has taken place at King Solomon in 30 years, which is significant because of new understanding of Carlin-type systems.

RECOMMENDATIONS

2026 Exploration

1. *Field work*: Detailed mapping and sampling in Priority 1 and Priority 2 areas (see Target Map in Appendix). 20 days.
2. *Geochemical survey*: consider tightening the existing 152m soil grid in the KS window and expanding the grid farther west to include the entire window and the adjacent fault-bound volcanics. Consider a small soils grid over the West Ridge area and tighten the existing grid over the Big Bend target.
3. *Work-up of existing data*: Compilation and evaluation of drill data; construction of cross sections and/or a 3D model of the King Solomon-south KS window using available datasets. An effort to locate and log drill cuttings from some historical drilling would potentially be useful. 10-15 days.

4. *Geophysical surveys*: consider extending IP surveys over West Ridge and the space between King Solomon mine and Big Bend.
5. *Drill hole planning*: King Solomon mine area, KS window, Big Bend, and West Ridge.
6. *Continuation of drill permitting process with the USFS*.
7. *Petrography, geochronology, and whole-rock geochemistry*: for the purpose of improving the geological understanding of the property and providing a strong framework.

Notes

1. The King Solomon mine area is quite prospective despite numerous historical holes. I recommend testing both strands of the King Solomon fault including drilling more of the Paleozoic carbonate rocks at depth. The southward extension of King Solomon is open with only sparse drilling.

2. The KS window is also highly prospective and the two areas share continuity in structure (King Solomon fault) and geology (e.g., Paleozoic lower-plate strata, vent-facies volcanic rocks). It possesses many of the components of many major Carlin-type systems: intrusion-cored, arched Paleozoic uplift, Roberts Mountains thrust with prospective lower-plate carbonates and a less favorable upper-plate siliciclastic section, a major early Cenozoic igneous center, complex structure in the window and major bounding normal faults, and abundance of dikes and other intrusions. The intersections between the Roberts Mountains thrust, high-angle faults, and prospective units at depth should be a high-priority target.

3. The caldera-related resurgence at King Solomon is much larger than the property. King Solomon covers the southwestern quarter, whereas there is likely a northwestern segment to the north (in the Table Mountain Wilderness). The resurgence isn't limited to the KS window, but rather it involves the early ash-flow tuffs and the King Solomon rhyolites, too. At King Solomon, both Paleozoic and Cenozoic rocks are very broadly tilted southward (south-younging). The eastern half of the resurgence is cut off by the King Solomon fault and down dropped to the east. The resurgence is important to understand because it represents the area underlain by intrusions at shallow depth in the footwall of the King Solomon fault; many of the smaller intrusions are exposed and mapped as dikes. There is a good correlation between the distribution of rhyolite lavas (i.e., near-vent volcanics), dikes and other small intrusions, doming, and hydrothermal alteration. It is reasonable to assume that ore fluids would have used the same conduits as roughly coeval magmas.

4. Polymictic intrusive breccia bodies are common and interesting feature of the volcanism and hydrothermal activity at King Solomon and point to a very shallow level of exposure

(mostly ≤ 1 km paleodepth). These breccias were mapped at the King Solomon mine, in the West Ridge area, in the broader area of dikes and other shallow intrusions in the northern Stone Cabin center, and along the immediate footwall of the King Solomon fault south of the mine. The breccias are interpreted to be phreatomagmatic rocks, i.e., shallow rhyolitic magmas that erupted due to the interaction between magma and groundwater.

Phreatomagmatic bodies directly tap the tops of intrusions and are therefore, commonly mineralized. Higher porosity and permeability of breccias make them more favorable to mineralize as well.

BRIEF DESCRIPTIONS OF SPECIFIC EXPLORATION TARGET AREAS

KS Window

This large target area covers ~ 1.7 km² in the northern third of the King Solomon property (Appendix G). The geological setting of this area is a Paleozoic lower-plate carbonate window that hosts numerous prospective strata for Carlin-type gold mineralization. The window occurs as a domal structure interpreted to be underpinned by a large silicic pluton. Geophysical surveys including gravity and magnetics support domal uplift by intrusion. The Roberts Mountains thrust occurs on the flanks of the window where it bounds less favorable upper-plate siliciclastic strata against lower-plate limestone in a structural setting similar to that hosting major Carlin-type gold deposits of the Carlin and Cortez trends in north-central Nevada. Major north-northeast striking normal faults like the King Solomon fault further provide structural preparation and helped to focus hydrothermal fluids. The KS Window contains abundant evidence for a robust Carlin-type system including: extensive alteration in the form of jasperoid, decalcified and dolomitic limestone and strong, broadly disseminated Au, As, Sb, and Hg geochemistry in soils and rock chips. This combination of a large target size, regional doming, a strong structural framework, prospective host rocks, widespread and robust surface geochemistry, and only sparse historical drilling make the KS Window target the highest ranking target at King Solomon. Only the extreme southern margin of the KS Window was drilled, in part by very shallow airtrack holes (30-165 ft) and also somewhat deeper conventional drill holes (~ 300 -700 ft).

West Ridge

This sizable target occurs along the west-central boundary of King Solomon (Appendix G). The area was explored by Newmont in 1989-1990 but no further work was done. Extensive rock-chip and soil sampling show robust Au, As, Sb, and Hg. West Ridge is underlain by abundant, highly altered rhyolitic dikes that locally formed polymictic breccias containing juvenile rhyolite clasts along with altered Paleozoic limestone (polymictic subvolcanic

breccia). The breccia bodies and margins of dikes are consistently altered and variably mineralized. Six widespread RC holes were drilled at West Ridge in 1990 ranging from 355 ft to 800 ft depth. Detailed mapping coupled with better definition of breccia pipes and bodies will be undertaken to improve targeting at depth.

Chalice

The Chalice target area is underlain by strongly bleached and quartz-sericite-altered, older ash-flow tuff immediately west of the KS Window (Appendix G). The area is bound by two significant north-northeast normal faults that are interpreted as splays of the King Solomon fault. The faults have resulted in what is interpreted as an area of shallow limestone based on the gravity data. The area contains an extensive soil geochemical expression in Au and As. Silicic dikes associated with the King Solomon center cut the older tuffs in this area. The dikes together with subparallel normal faults are considered to be important for localizing alteration and mineralization. Relatively shallow lower-plate carbonates may yield bulk-disseminated styles of Carlin-type mineralization. No drilling has occurred in this target.

King Solomon Mine

The King Solomon Mine area includes the area of historical, small-scale antimony mining (Appendix G). It is extensively drilled but only to shallow depths generally less than 500 ft. The area contains both highly mineralized jasperoid bodies derived from lower-plate limestone and widespread and robust quartz-sericite alteration in the volcanic and subvolcanic rhyolites. A small, non-43-101-compliant resource exists in the south end of the target area and consists of numerous holes that bottomed in significant gold mineralization. Other areas beyond the small resource are much less evaluated. The King Solomon mine area should be evaluated with deeper drill holes across the King Solomon fault zone, which here consists of two major subparallel strands. The latest drilling by Kennecott in 1996 included on hole directed northwest away from the resource (KKS-3) that encountered robust mineralization in what is interpreted as the footwall of the westernmost King Solomon fault strand.

Big Bend

The Big Bend target is centered on the prominent flexure of the King Solomon fault zone from northeast-striking on the north to north-south-striking on the south (Appendix G). The bend localized strong alteration including silicification in volcanic rocks. Several shallow holes drilled near the range front encountered thin intervals of gold but the area remains poorly tested. It's possible that strain developed at the bend in the King Solomon fault may have contributed to more extensive breakage of rocks there. A gravity high suggests that the

footwall of the King Solomon fault is prospective for relatively shallow carbonate. Dikes and polymictic breccia bodies exposed along the KS fault's footwall comprise the altered and locally mineralized roots to lava flows.

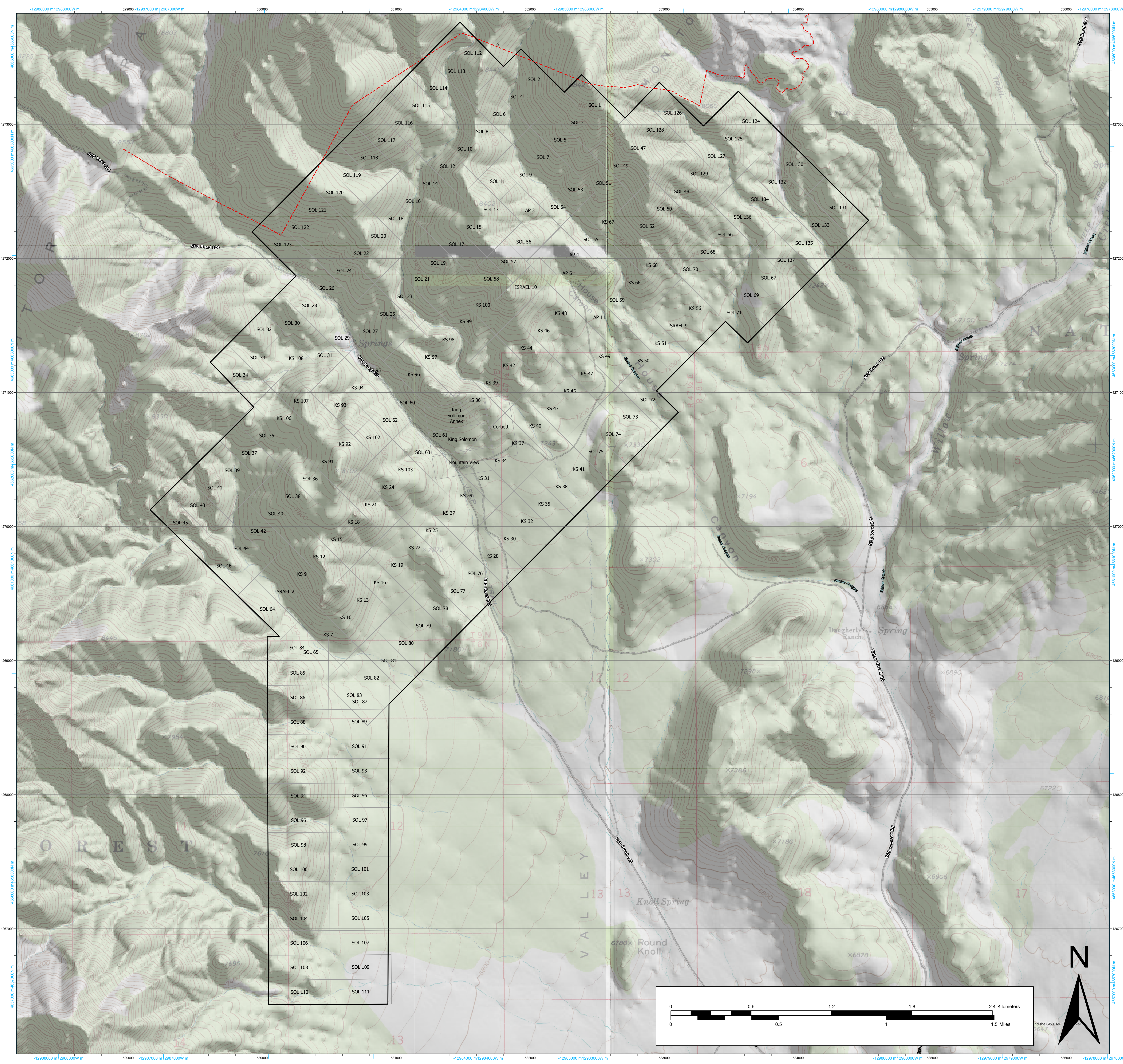
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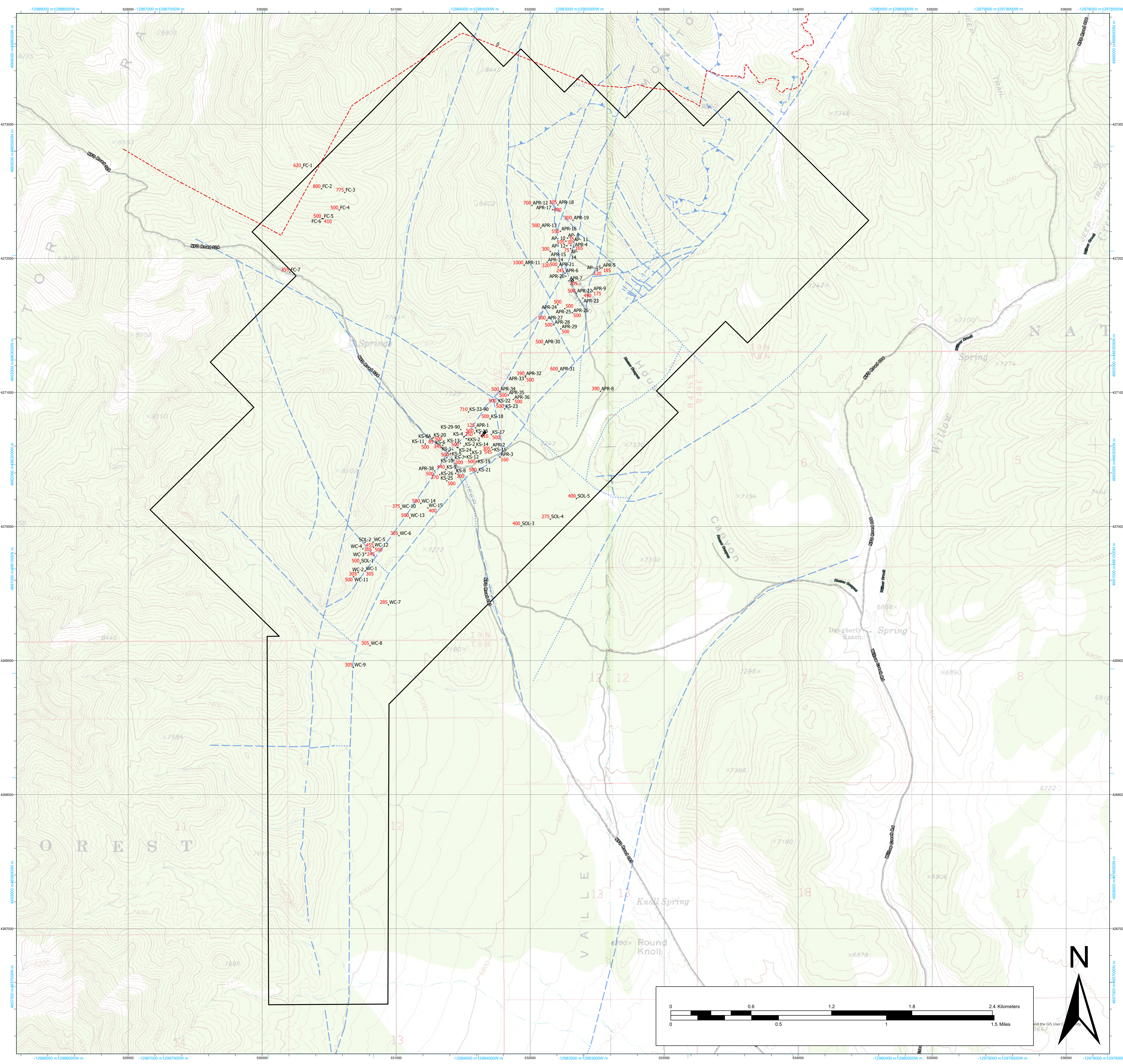
Appendix A.
Claims Map
King Solomon Exploration Project,
Northern Nye County, Nevada
March 2026



Claim map of the King Solomon project, Monitor Range,
 northern Nye County, Nevada

by
 Michael W. Ressel
 for
 Phenom Resources Corp.
 October 2025

Appendix B.
Drill Holes Map
King Solomon Exploration Project,
Northern Nye County, Nevada
March 2026

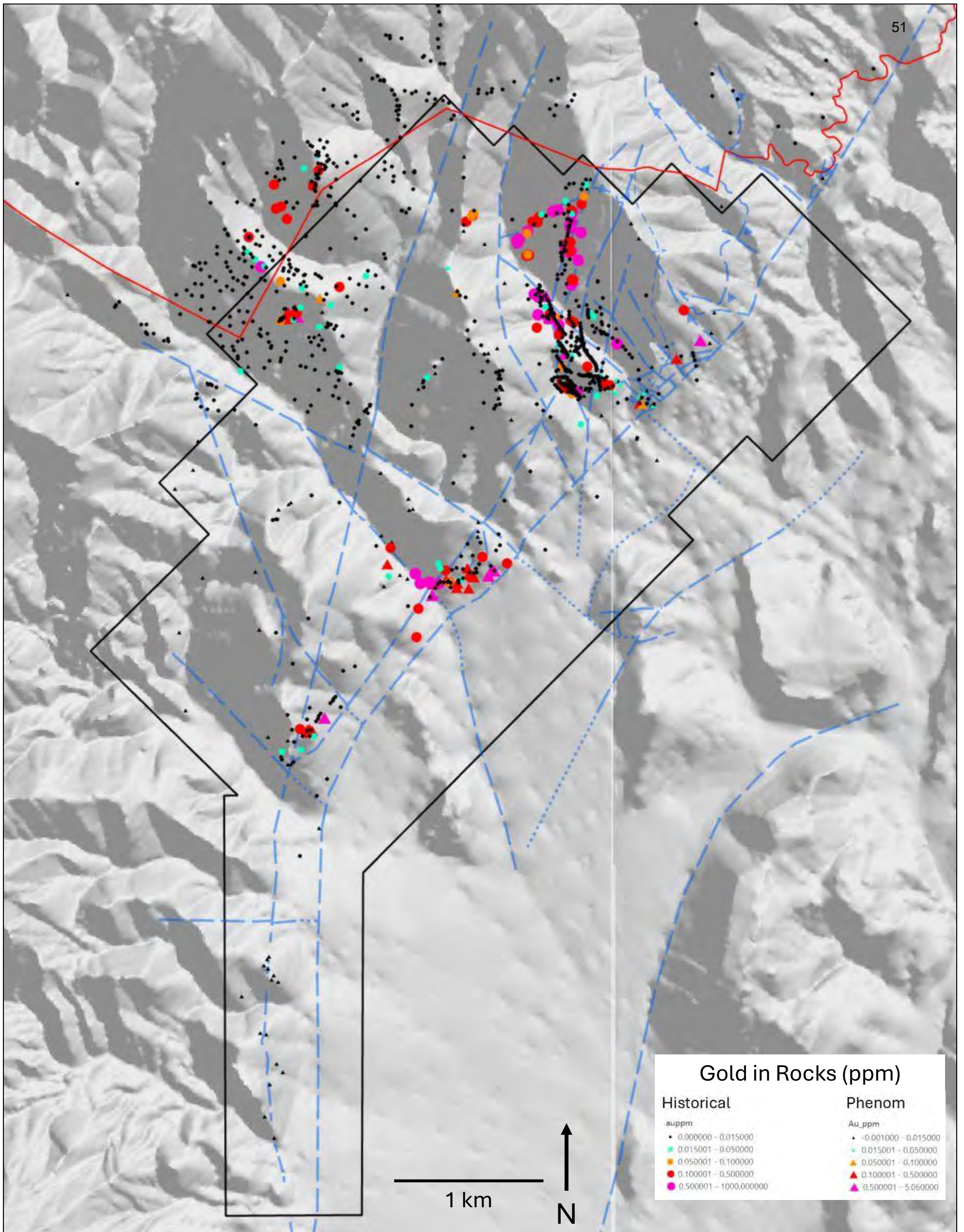


● Drill hole and total depth (ft)

Drill hole map of the King Solomon project, Monitor Range, northern Nye County, Nevada

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

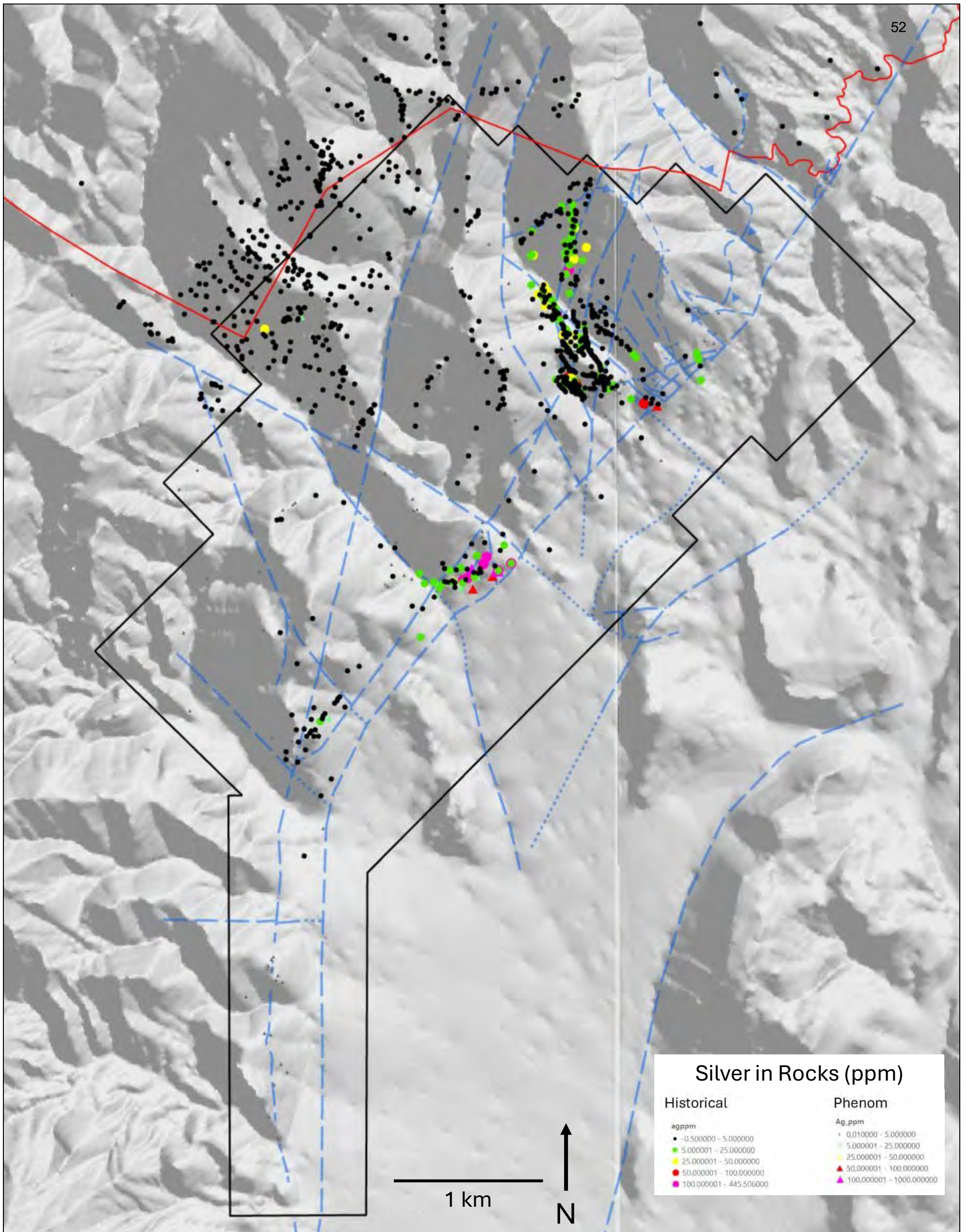
Appendix C.
Rock Chip Samples
2025 Samples + Historical Samples
(Au, Ag, As, Sb, Hg)
King Solomon Exploration Project,
Northern Nye County, Nevada
March 2026



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Gold in Rocks
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada

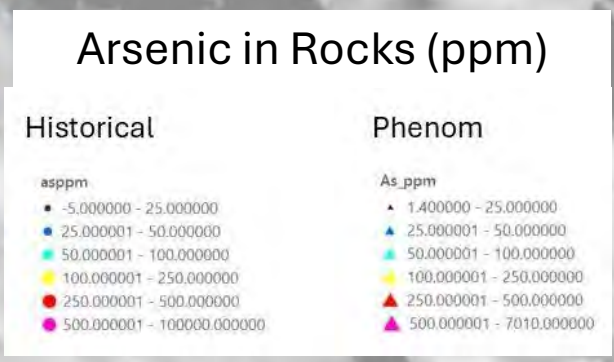
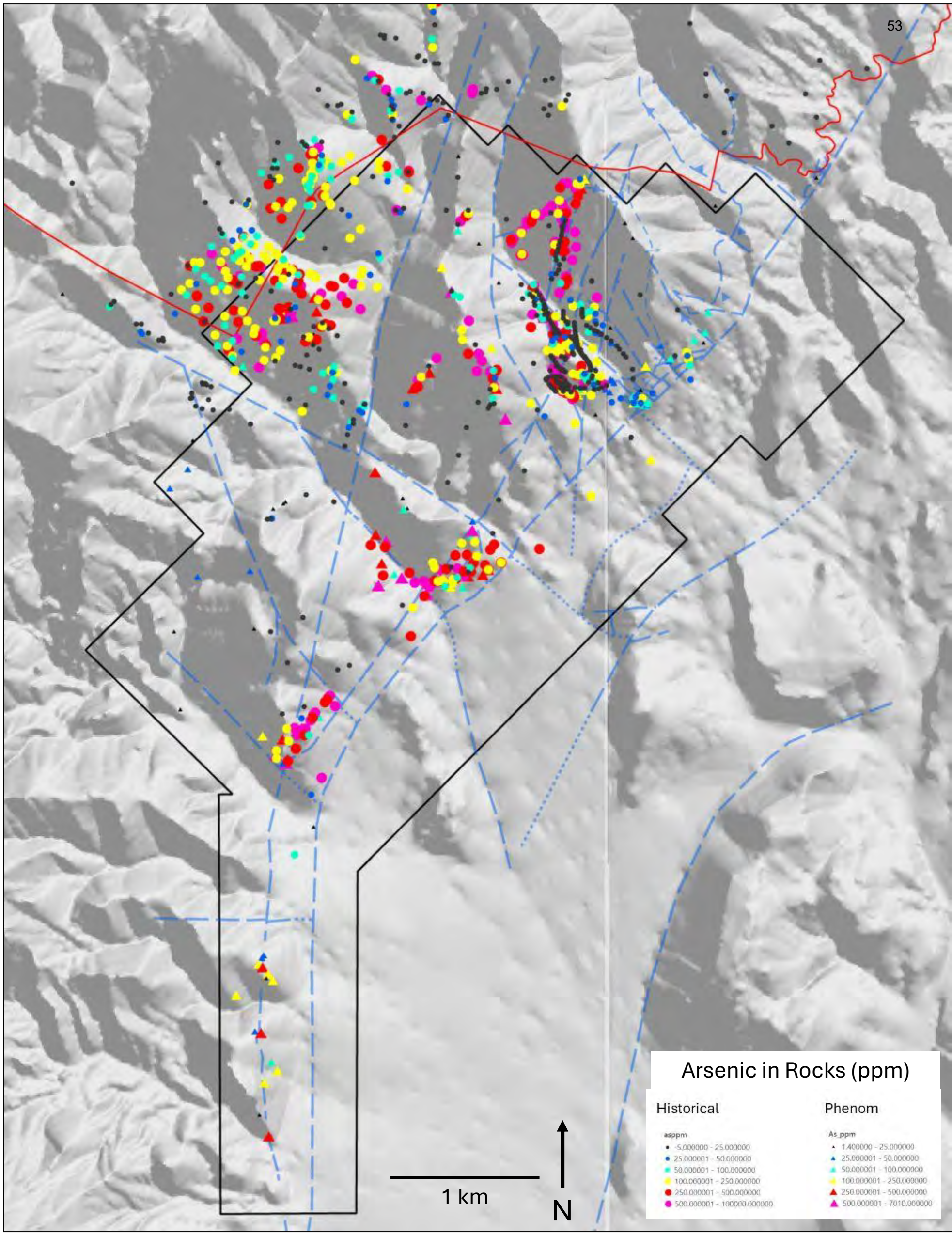
 Phenom Resources Corp.
 April 2026



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Silver in Rocks
King Solomon Project, Monitor Range,
Northern Nye County, Nevada

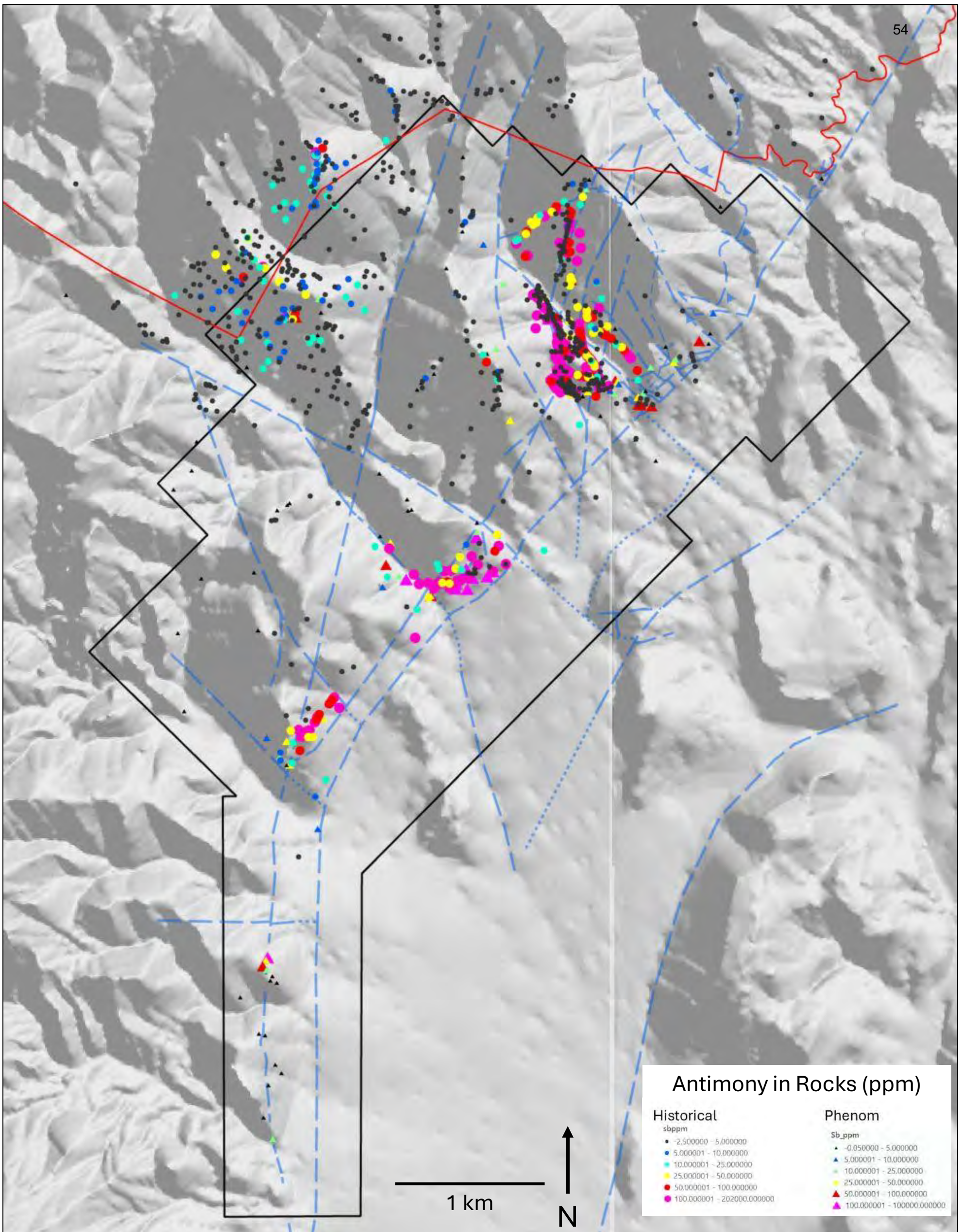
Phenom Resources Corp.
April 2026



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Arsenic in Rocks
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada

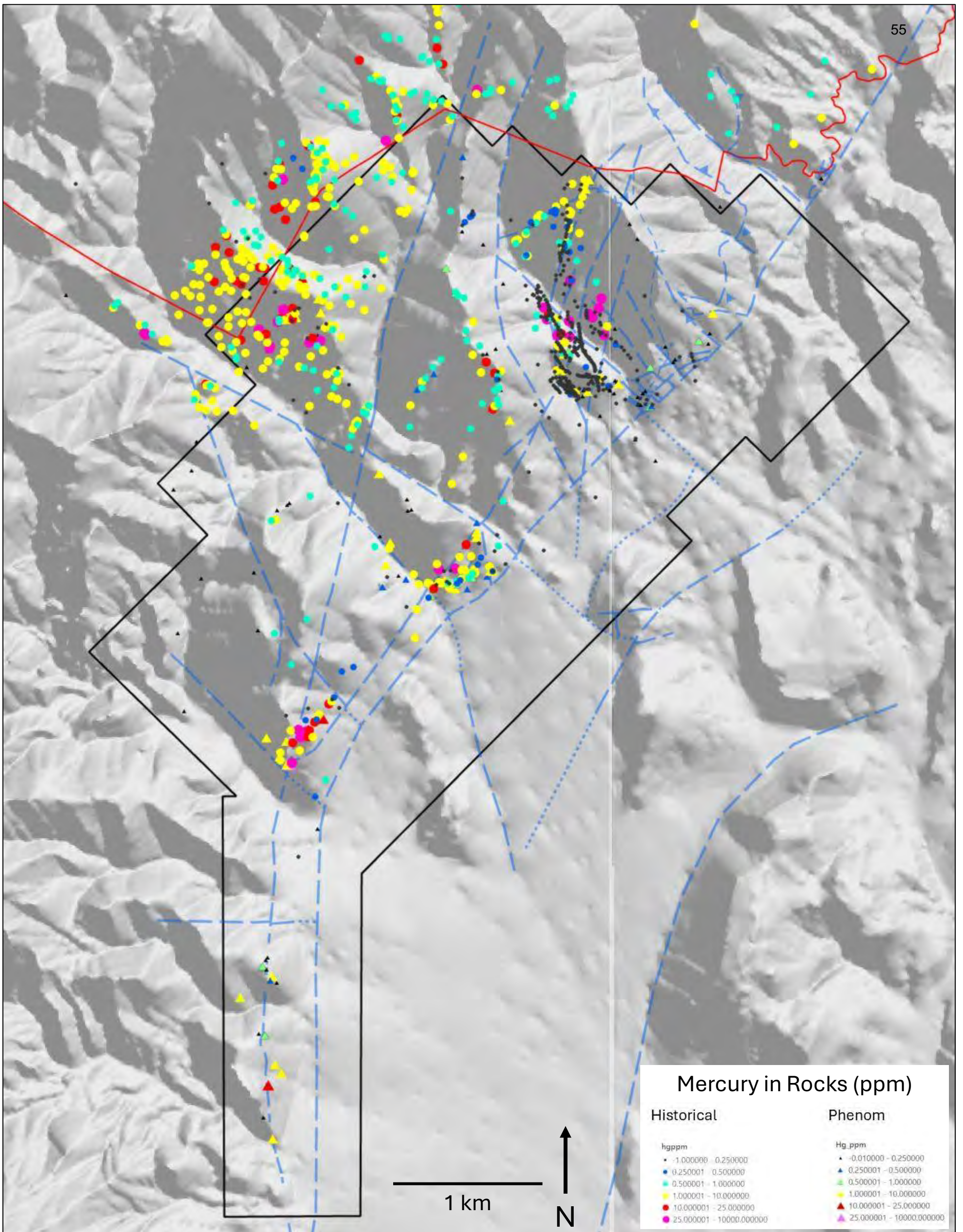
 Phenom Resources Corp.
 April 2026



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Antimony in Rocks
King Solomon Project, Monitor Range,
Northern Nye County, Nevada

Phenom Resources Corp.
April 2026



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Mercury in Rocks
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada

Phenom Resources Corp.
 April 2026

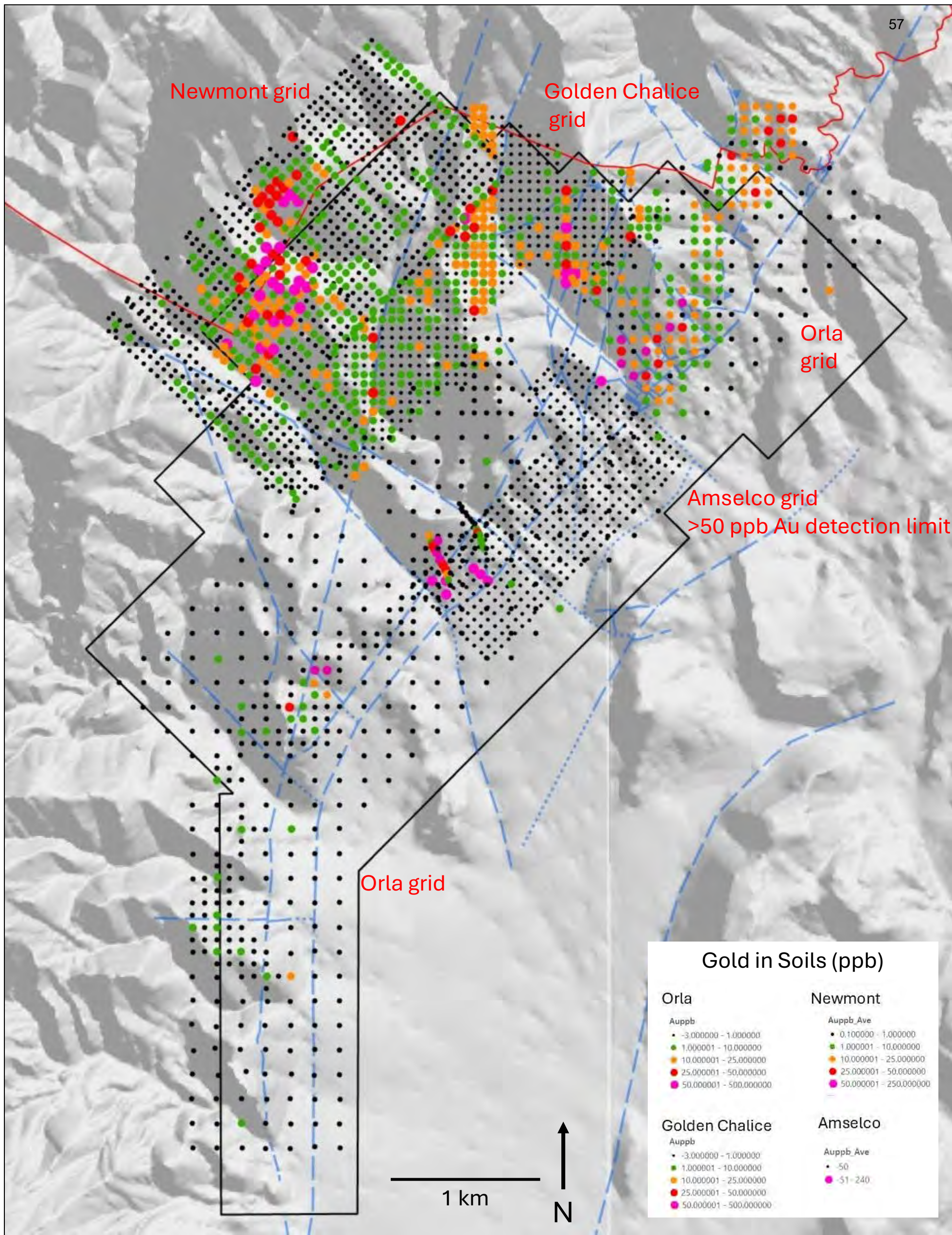
Appendix D.
Soil Samples
Historical Data
(Au, Ag, As, Sb, Hg)
King Solomon Exploration Project, Northern Nye
County, Nevada March 2026

Amselco survey (1981; >50 ppb Au detection limit)

Newmont survey (1990)

Golden Chalice survey (2006-7)

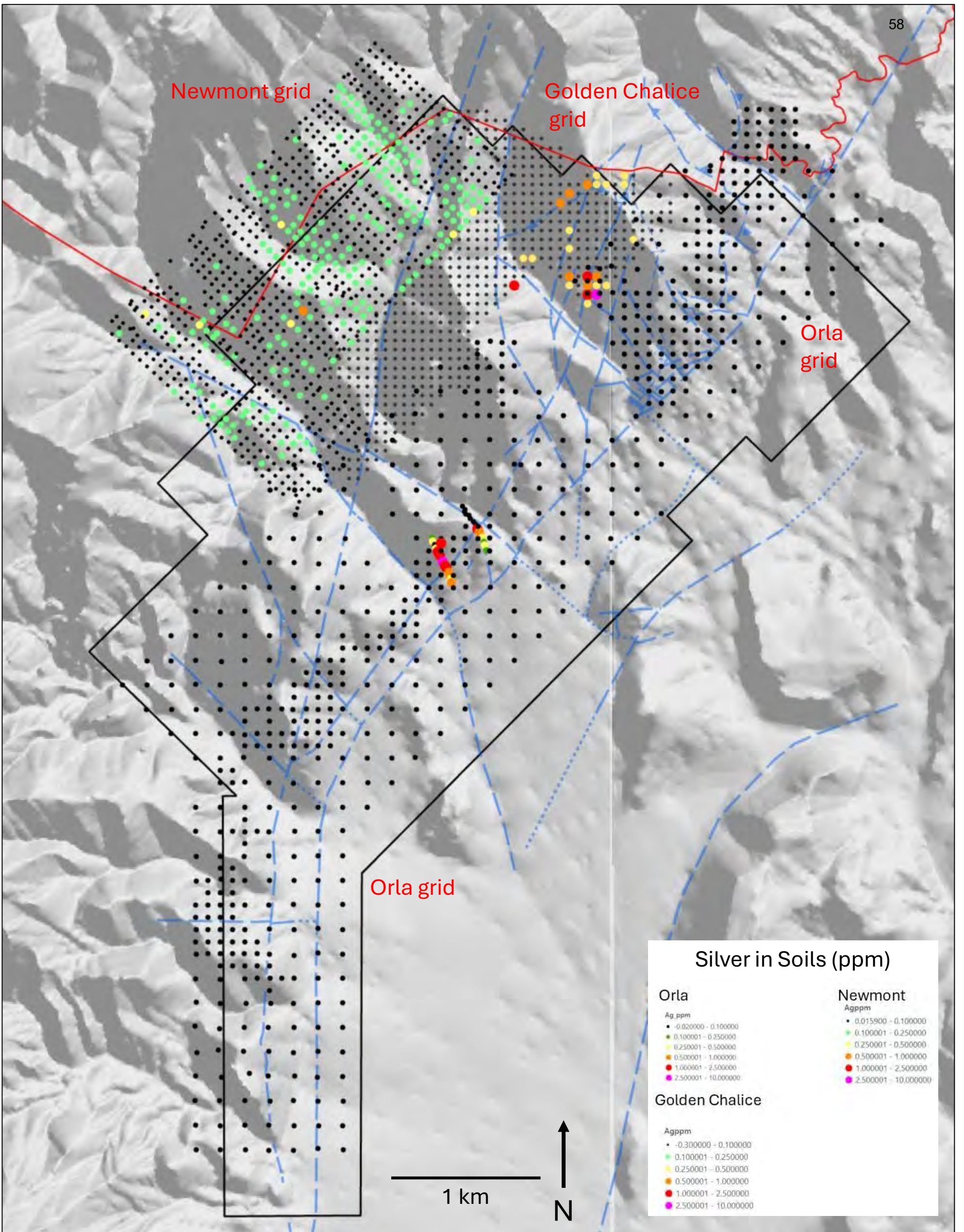
Orla survey (2017)



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Gold in Soils
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada

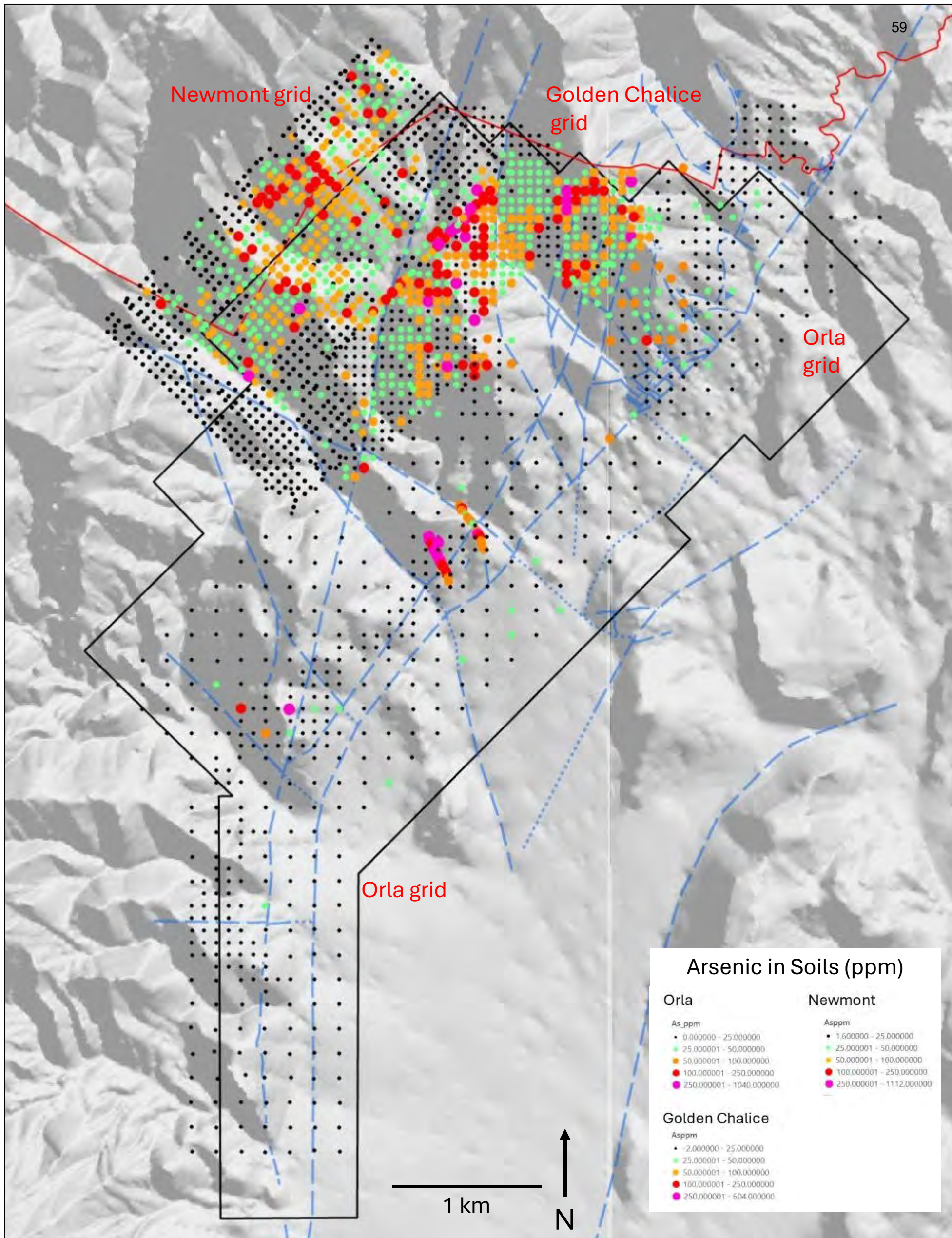
Phenom Resources Corp.
 April 2026



- King Solomon claim outline
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Silver in Soils
King Solomon Project, Monitor Range,
Northern Nye County, Nevada

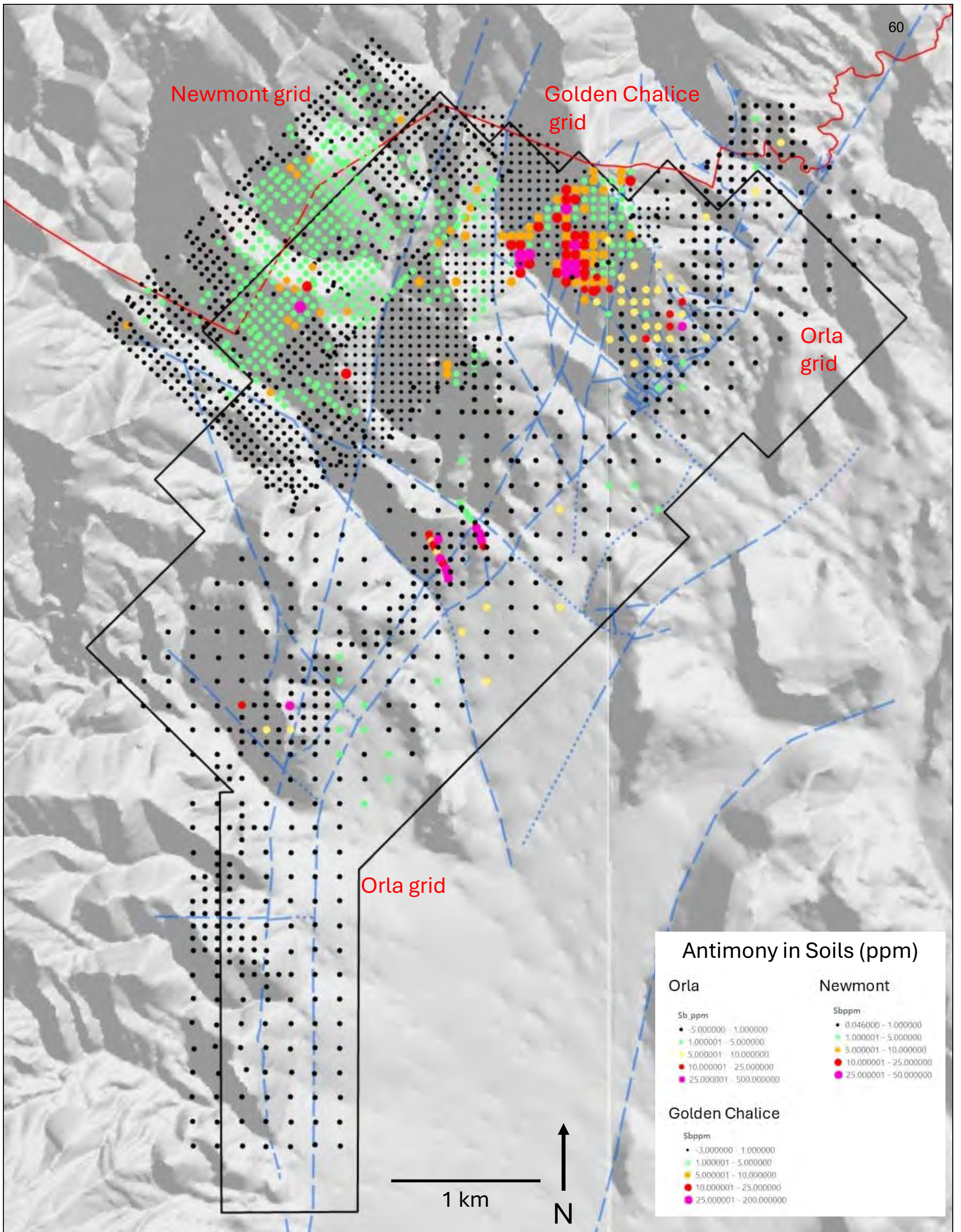
Phenom Resources Corp.
April 2026



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Arsenic in Soils
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada

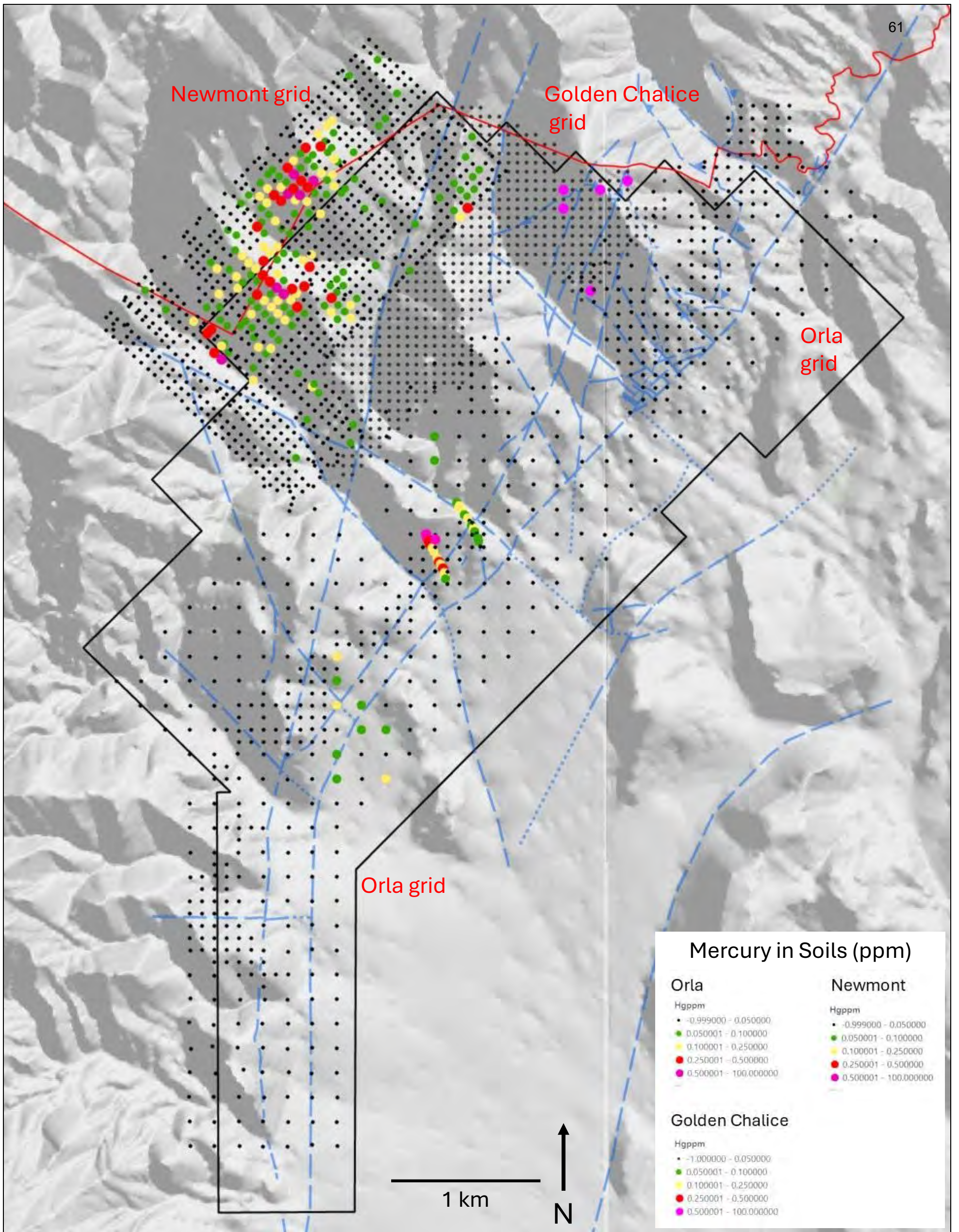
 Phenom Resources Corp.
 April 2026



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Antimony in Soils
King Solomon Project, Monitor Range,
Northern Nye County, Nevada

Phenom Resources Corp.
April 2026



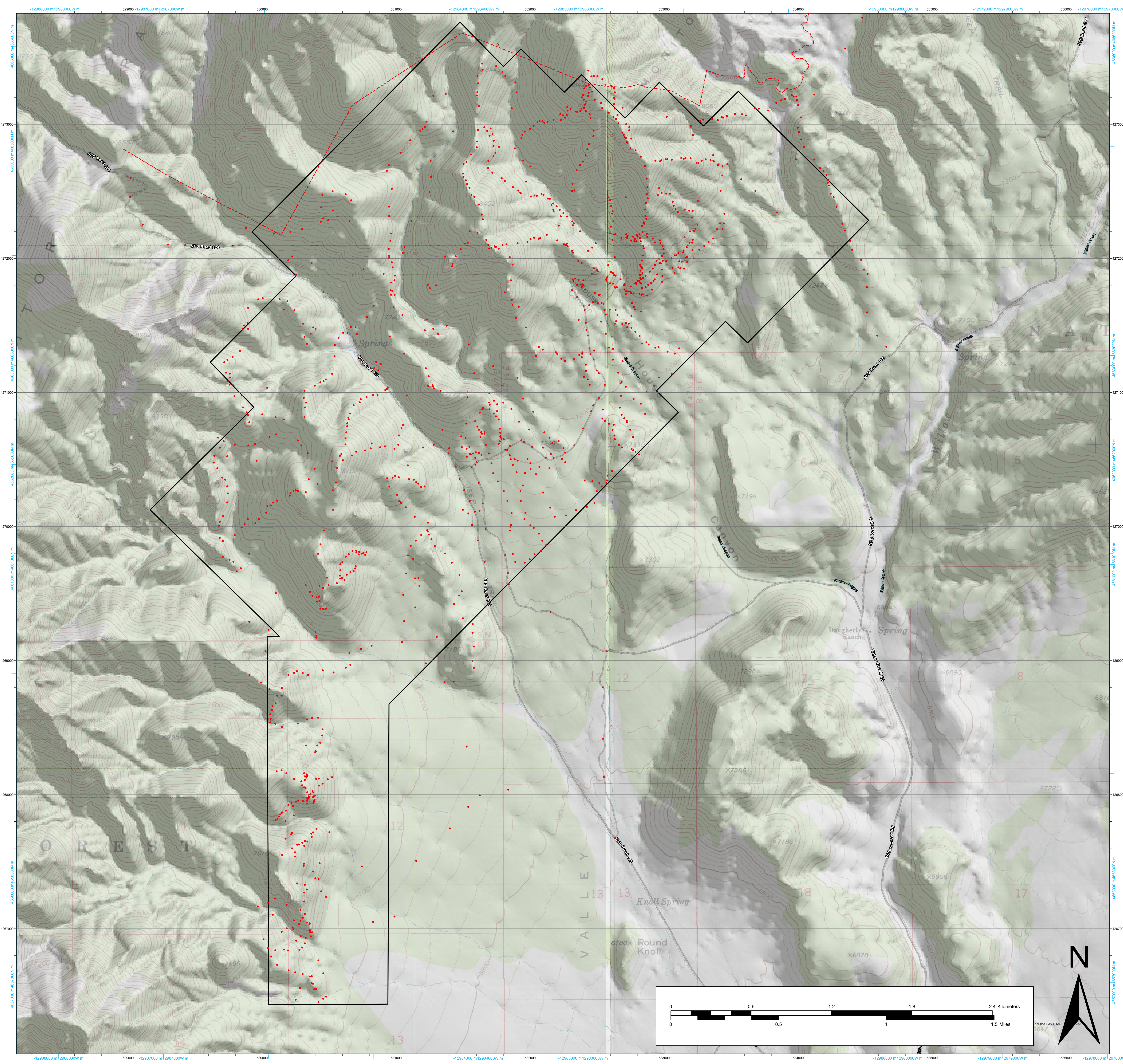
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Mercury in Soils
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada
 Phenom Resources Corp.
 April 2026

Appendix E.
2025 Map Data:

1. Map Data Points
2. Outcrop Maps
3. Structural Measurements
4. Geologic Maps
5. Alteration Maps

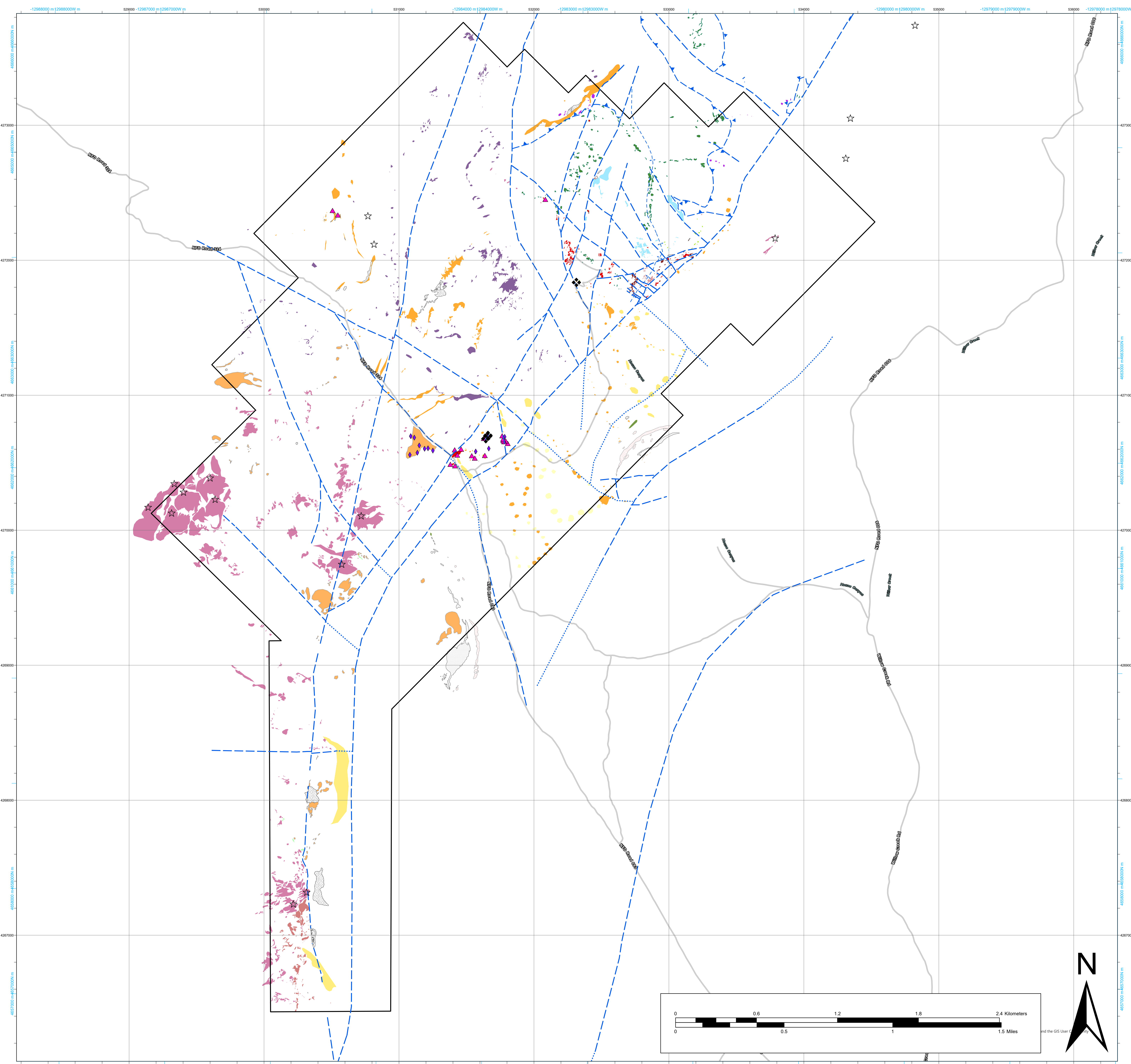
King Solomon Exploration Project,
Northern Nye County, Nevada
March 2026



Claim map of the King Solomon project, Monitor Range,
northern Nye County, Nevada

Map Data Points (red dots)

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025



**Outcrop map of the King Solomon project, Monitor Range,
northern Nye County, Nevada**
by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

- Quaternary**
 - Qf_Qal_Qt Alluvial (Qal) and alluvial fan (Qaf) deposits; lesser talus apron deposits (Qt) on steeper slopes
 - Qoa Older alluvial and alluvial fan deposits present as perched terrace remnants at canyon mouths
 - Qls Landslide deposits
 - Tbt2 Upper beige biotite-phyric, fine pumice-bearing, vitrophyric rhyolite ash-flow tuff
 - Tbt Lower beige biotite-phyric, fine pumice-bearing vitrophyric ash-flow tuff
 - Trt Red tuff, welded pyroxene-bearing, coarse pumice-rich dacite ash-flow tuff
 - Ts2 White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks
 - Tfpr Stone Cabin rhyolite; yellow weathering, strongly flow foliated, sparsely porphyritic rhyolite lavas and domes
 - Tts White, planar-laminated to thin-bedded tuffaceous sedimentary rocks, generally altered and with Liesegang bands
 - Tpr Willow rhyolite. Spheroidally weathering, boldly outcropping, flow-foliated, maroon gray, plg>bio>qtz rhyolite lavas
 - Thr House rhyolite. Intrusive, moderately porphyritic (plg>bio>qtz>san). May be intrusive equivalent to Tpr Willow lavas.
 - Tqprl Strongly flow foliated, white, altered, quartz-phyric rhyolite lavas associated with massive Tqpr dikes
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 - Tpbx Massive, speckled dark gray and white, polymictic lithic breccia with abundant rounded to subangular, coarse to fine, juvenile rhyolite and fine black sedimentary clasts. Groundmass consists of rock flour. Ubiquitous hydrothermal alteration. Associated with Tqpr rhyolite dikes
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 - Tcpt Older, maroon, hackly fractured, coarse to medium pumice-rich rhyolitic ash-flow tuff; multiple thin cooling units
- Oligocene(?) caldera-related ash-flow tuffs**
 - D-Osc Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. *Roberts Mountains or similar thrust*
 - Dd-stms Mostly tectonized, medium to light gray, hackly fractured, thin- to medium-bedded lime mudstone/wackestone, weathers to mottled gray and yellow. Possibly the Denay or Rabbit Hill formation.
 - Dd-ls Cliff-forming, massive, medium to pale gray carbonate packstone to grainstone; clasts rounded to 1cm, debris-flow carbonate; locally dolomitic. Possibly the lower Denay or Rabbit Hill formation.
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 - Oe Massive, cliff-forming, spheroidally weathering, mottled pale gray with gray streaks, medium-grained quartzite, has distinct clear, rounded quartz grains. Eureka Quartzite.
 - Op Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

map-scale alteration

- jasperoid breccia
- jasperoid
- wilderness boundary
- mine
- property outline

measurements

- volcanic vent
- anticline_approx
- thrust_inferred
- fault_inferred
- bedding
- volcanic foliation
- tectonic foliation
- veins
- folds
- faults
- dikes
- Barite occurrence
- Silbite occurrence

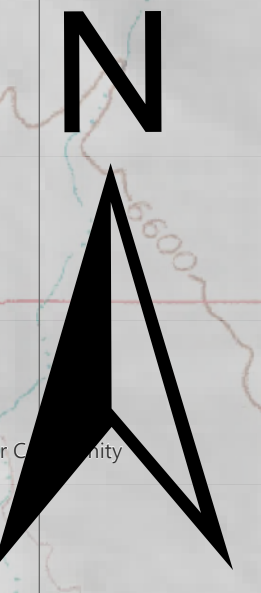
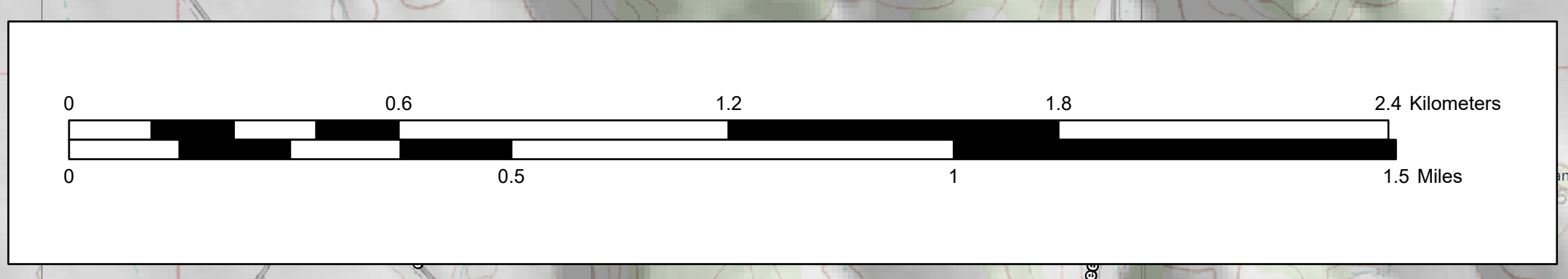
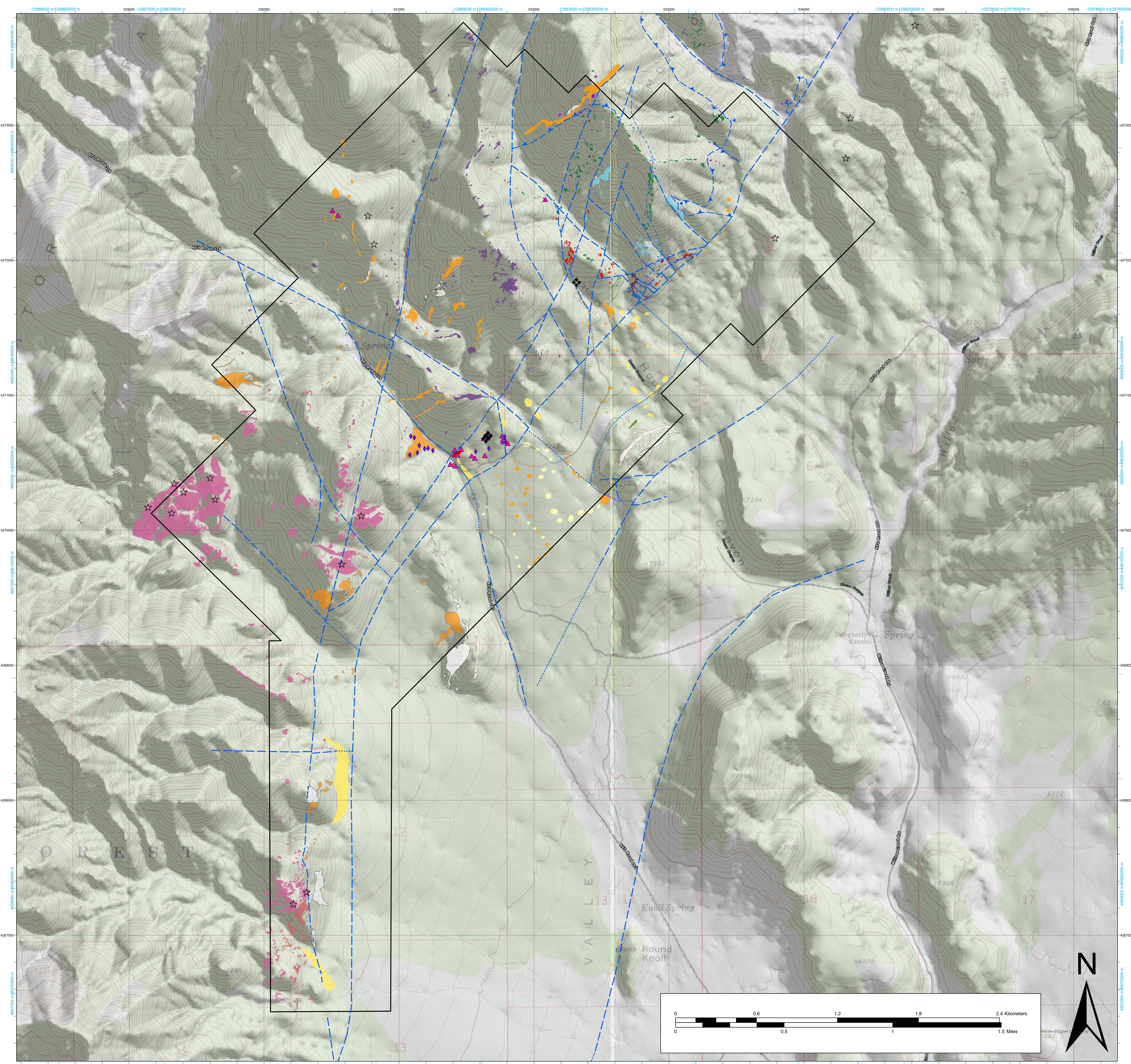
*dark polygons within map units are outcrops

UTM NAD83 (2011) datum

Magnetic declination: 12°E

Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM

Grid: UTM 1983, zone 11, 500 meters

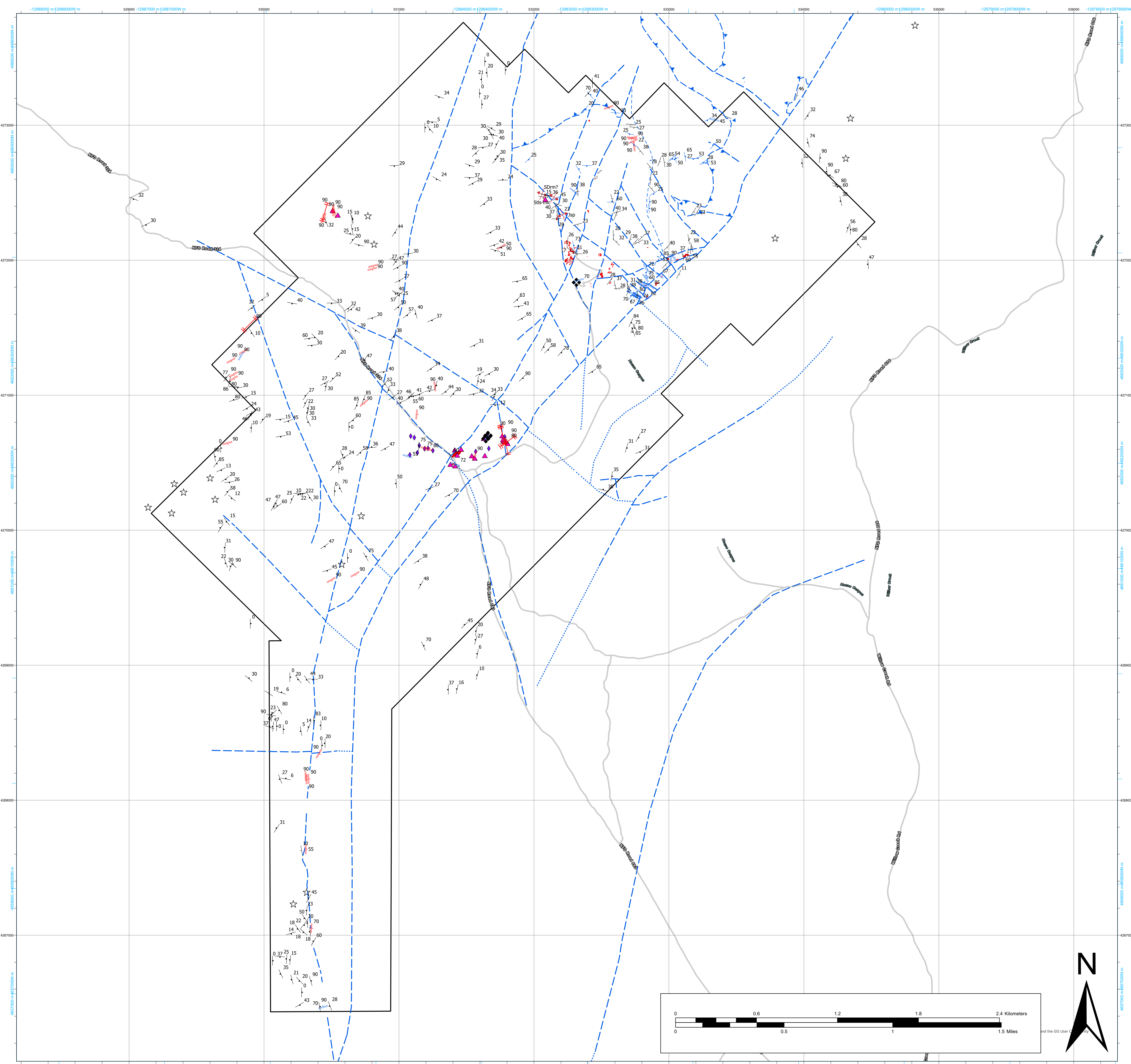


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**Outcrop map of the King Solomon project, Monitor Range,
northern Nye County, Nevada**

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

- map-scale alteration**
- jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
 - property outline
- measurements**
- volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - folds
 - faults
 - dikes
 - Biotite occurrence
 - Silbite occurrence
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E
- Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM
- Grid: UTM 1983, zone 11, 500 meters

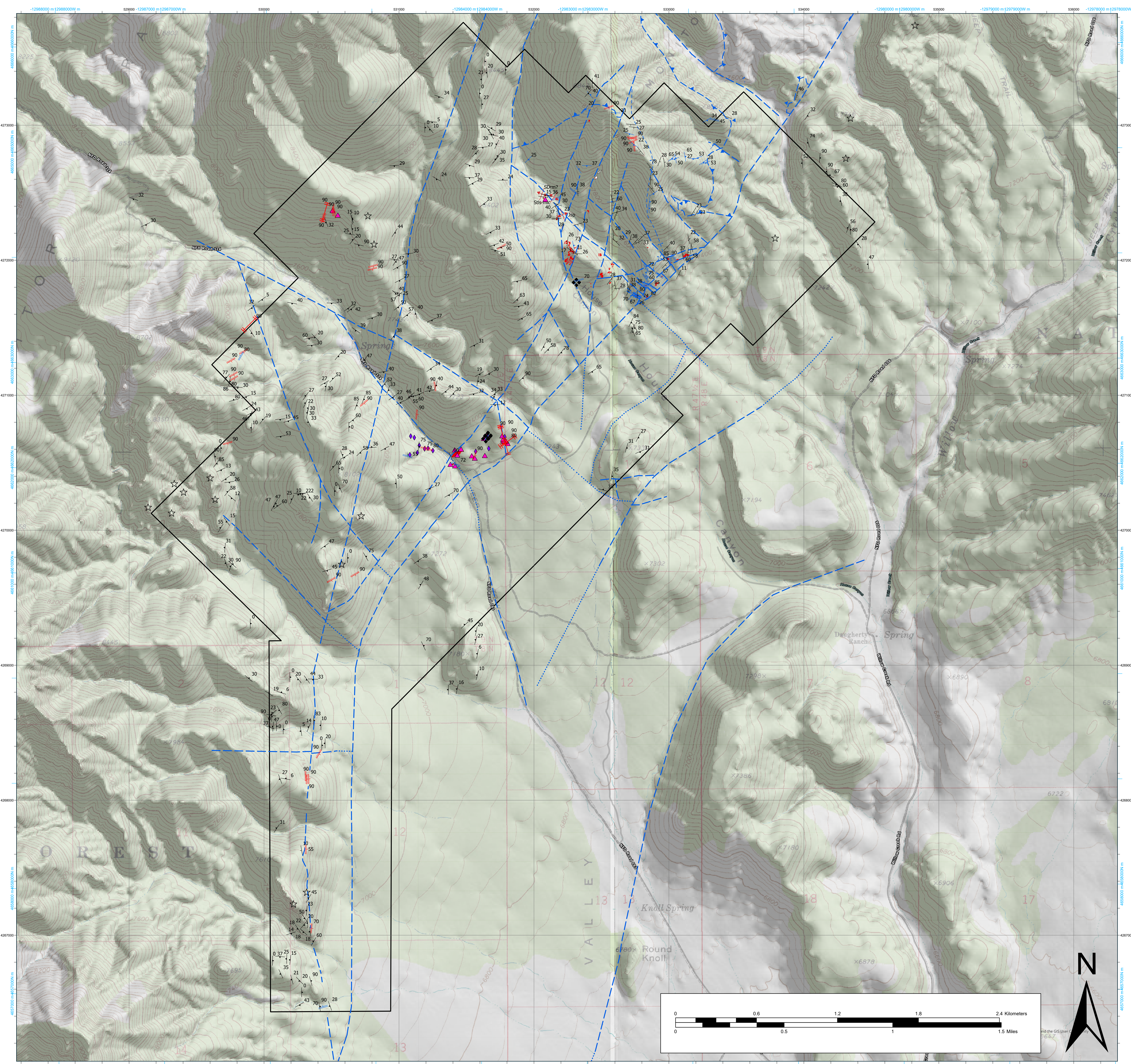


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Structural features map of the King Solomon project, Monitor Range, northern Nye County, Nevada

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

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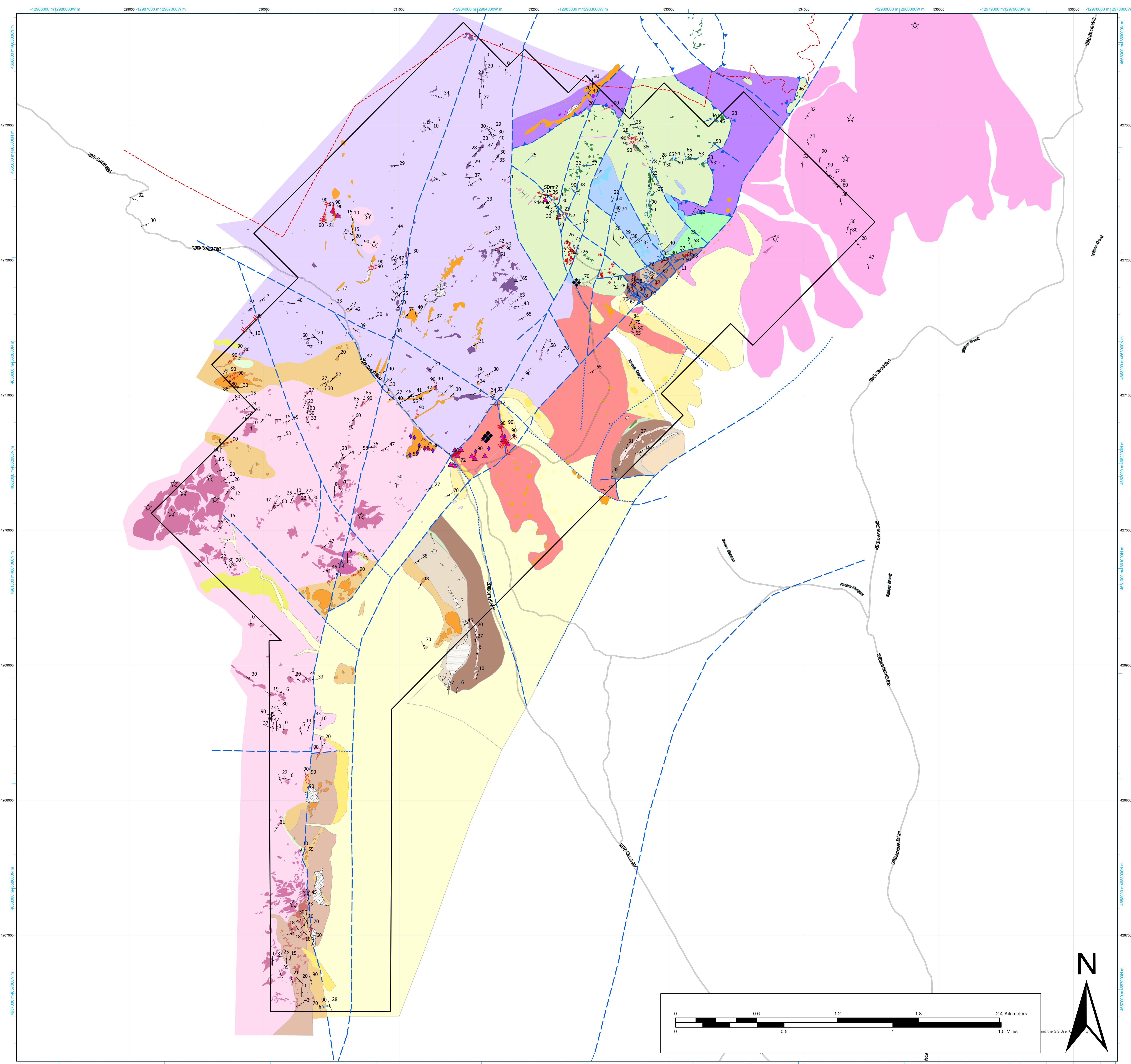


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**Geologic map of the King Solomon project, Monitor Range,
northern Nye County, Nevada**
by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

- | | | |
|--|-----------|---|
| Quaternary | Qf_Qal_Qt | Alluvial (Qal) and alluvial fan (Qaf) deposits; lesser talus apron deposits (Qt) on steeper slopes |
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| | Qls | Landslide deposits |
| | Tbt2 | Upper beige biotite-phyric, fine pumice-bearing, vitrophyric rhyolite ash-flow tuff |
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King Solomon
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| Oligocene(?)
caldera-
related ash-
flow tuffs | Tcpt | Older, maroon, hackly fractured, coarse to medium pumice-rich rhyolitic ash-flow tuff; multiple thin cooling units |
| Paleozoic | D-Osc | Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. <i>Roberts Mountains or similar thrust</i> |
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- dikes

mineral occurrences

- barite
- silbrite

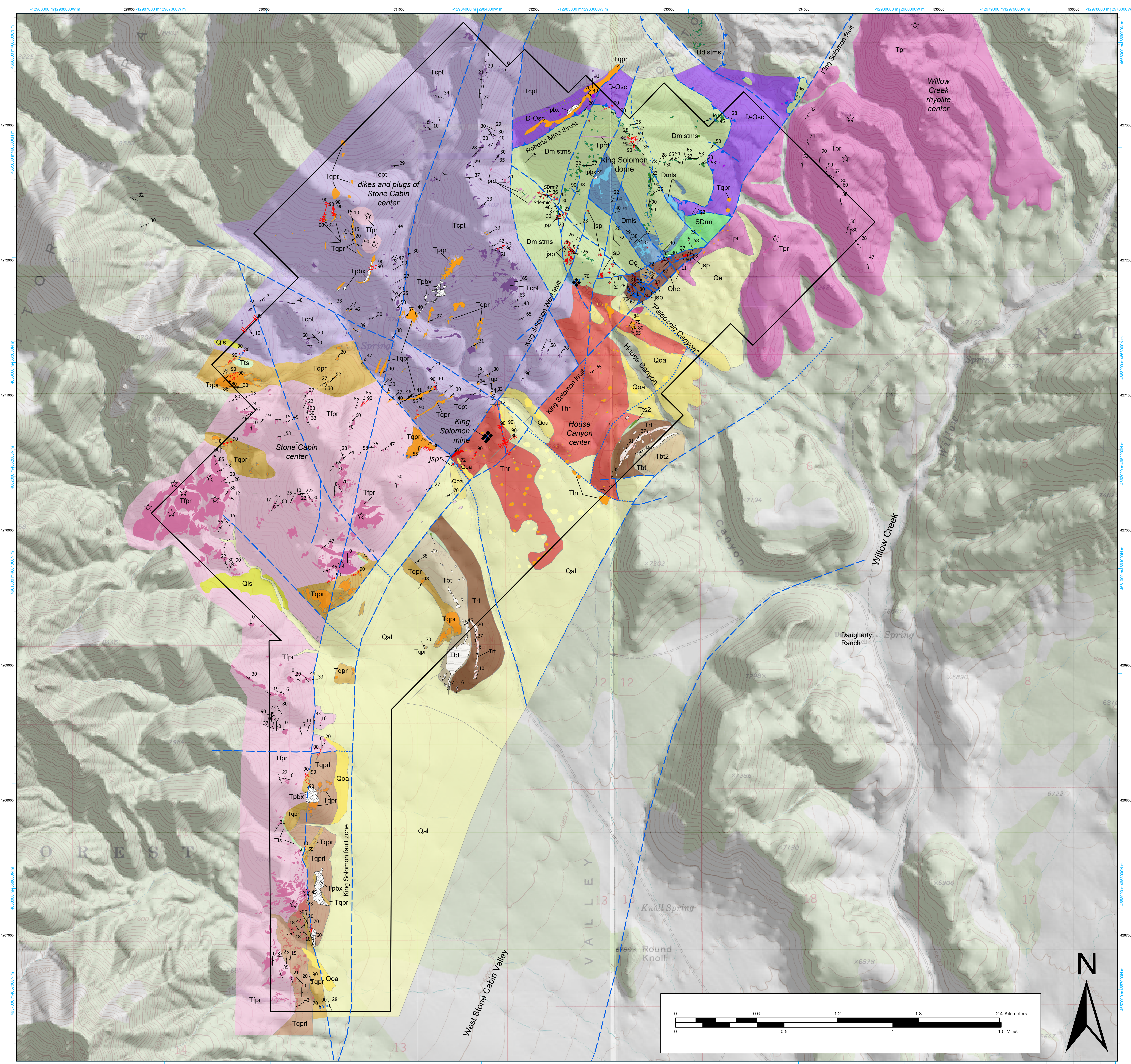
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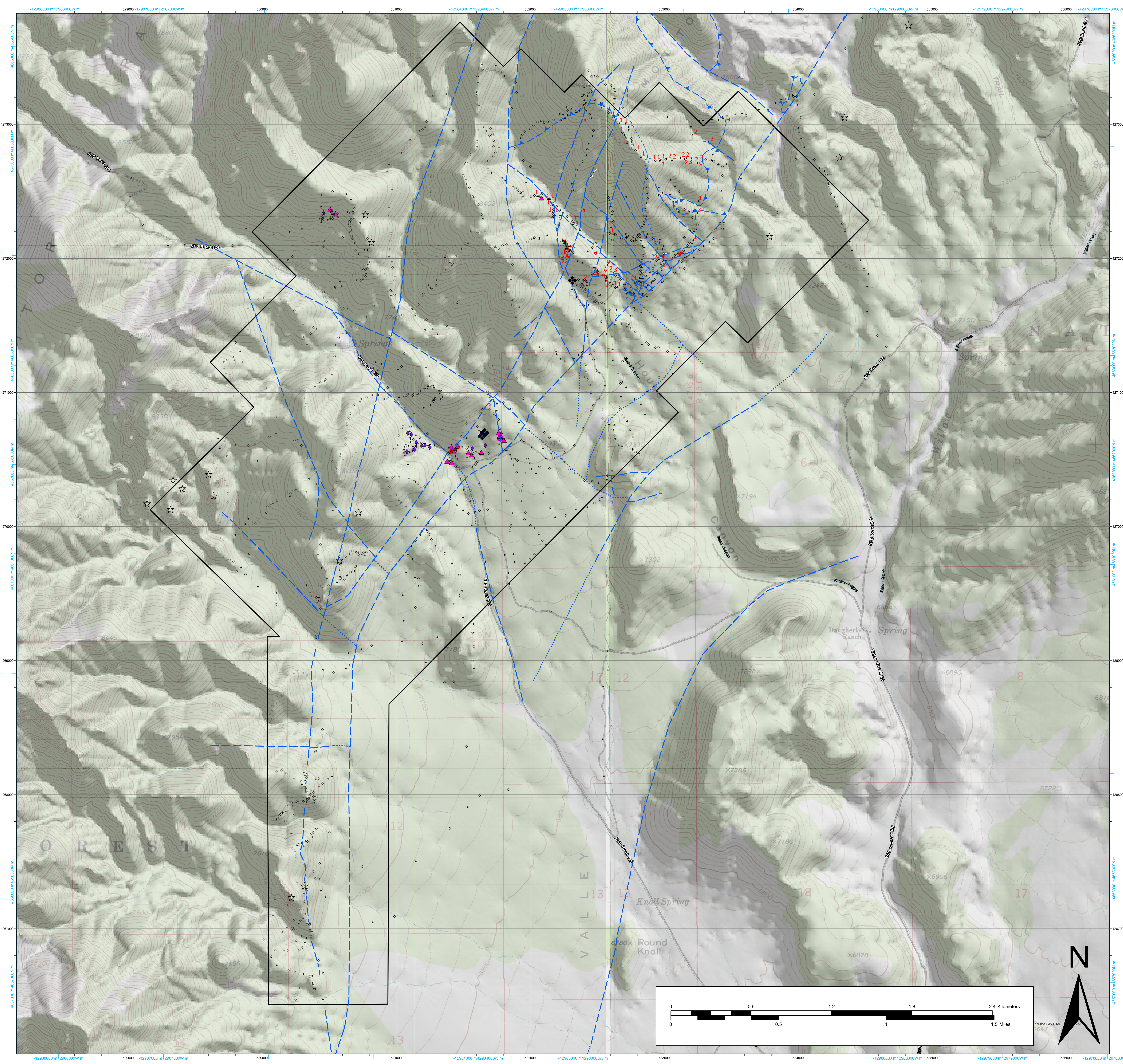
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volcanic
center | Tts2 | White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks |
| | Tfpr | Stone Cabin rhyolite; yellow weathering, strongly flow foliated, sparsely porphyritic rhyolite lavas and domes |
| | Tts | White, planar-laminated to thin-bedded tuffaceous sedimentary rocks, generally altered and with Liesegang bands |
| | Tpr | Willow rhyolite. Spheroidally weathering, boldly outcropping, flow-foliated, maroon gray, plg>bio>qtz rhyolite lavas |
| | Thr | House rhyolite. Intrusive moderately porphyritic (plg>bio>qtz>san). May be intrusive equivalent to Tpr Willow lavas. |
| | Tqprl | Strongly flow foliated, white, altered, quartz-phyric rhyolite lavas associated with massive Tqpr dikes |
| | Tqpr | Massive, commonly wall-like exposures of spheroidally weathered, sparsely porphyritic qtz-rich rhyolite dikes |
| | Tpbx | Massive, speckled dark gray and white, polymictic lithic breccia with abundant rounded to subangular, coarse to fine, juvenile rhyolite and fine black sedimentary clasts. Groundmass consists of rock flour. Ubiquitous hydrothermal alteration. Associated with Tqpr rhyolite dikes |
| Oligocene(?)
caldera-
related ash-
flow tuffs | Tcpt | Older, maroon, hackly fractured, coarse to medium pumice-rich rhyolitic ash-flow tuff; multiple thin cooling units |
| Paleozoic | D-Osc | Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. <i>Roberts Mountains or similar thrust</i> |
| | Dd-stms | Mostly tectonized, medium to light gray, hackly fractured, thin- to medium-bedded lime mudstone/wackestone, weathers to mottled gray and yellow. Possibly the Denay or Rabbit Hill formation. |
| | Dd-ls | Cliff-forming, massive, medium to pale gray carbonate packstone to grainstone; clasts rounded to 1cm, debris-flow carbonate; locally dolomitic. Possibly the lower Denay or Rabbit Hill formation. |
| | SDrm | Beige weathering, planar-laminated to thin-bedded, dark gray limestone and dolomitic limestone; has thin quartz-rich interlayers. Slope-former. Roberts Mountains Formation. |
| | Ohc | Massive, cliff-forming, yellowish-brown weathering, recrystallized, granular dolomite with considerable karst dissolution features. Hanson Creek Formation. |
| | Oe | Massive, cliff-forming, spheroidally weathering, mottled pale gray with gray streaks, medium-grained quartzite, has distinct clear, rounded quartz grains. Eureka Quartzite. |
| | Op | Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone. |

Geologic map of the King Solomon project, Monitor Range, northern Nye County, Nevada
 by
 Michael W. Ressel
 for
 Phenom Resources Corp.
 October 2025

- map-scale alteration**
- jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
 - property outline
- measurements**
- volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - folds
 - faults
 - dikes
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E
- Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM
- Grid: UTM 1983, zone 11, 500 meters



- 2 ○ Dolomitization intensity (1 through 3)
- ◆ Barite occurrence
- ▲ Stibnite occurrence

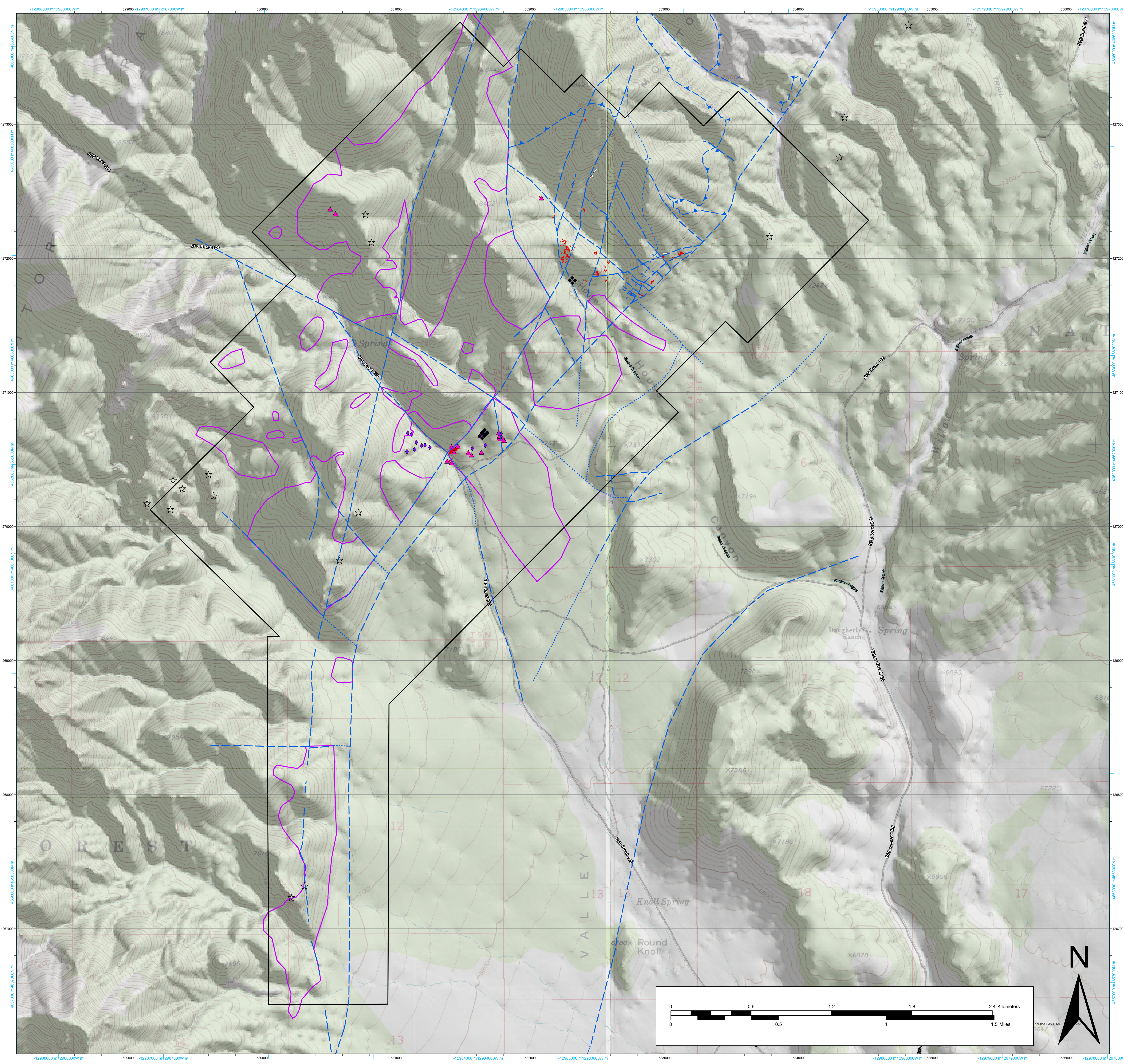
- map-scale alteration**
- ◻ jasperoid breccia
 - jasperoid
 - - wilderness boundary
 - ◆ mine
 - ◻ property outline
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E




- measurements**
- ☆ volcanic vent
 - + anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - ↔ folds
 - faults
 - dikes

Alteration map of the King Solomon project, Monitor Range, northern Nye County, Nevada

Dolomitization

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

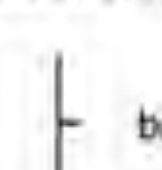
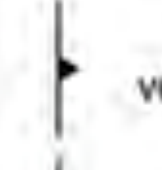
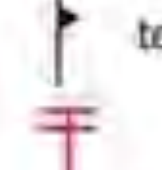










-  Moderate to strong sericitization
-  Barite occurrence
-  Stibnite occurrence

map-scale alteration

-  jasperoid breccia
 -  jasperoid
 -  wilderness boundary
 -  mine
 -  property outline
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E

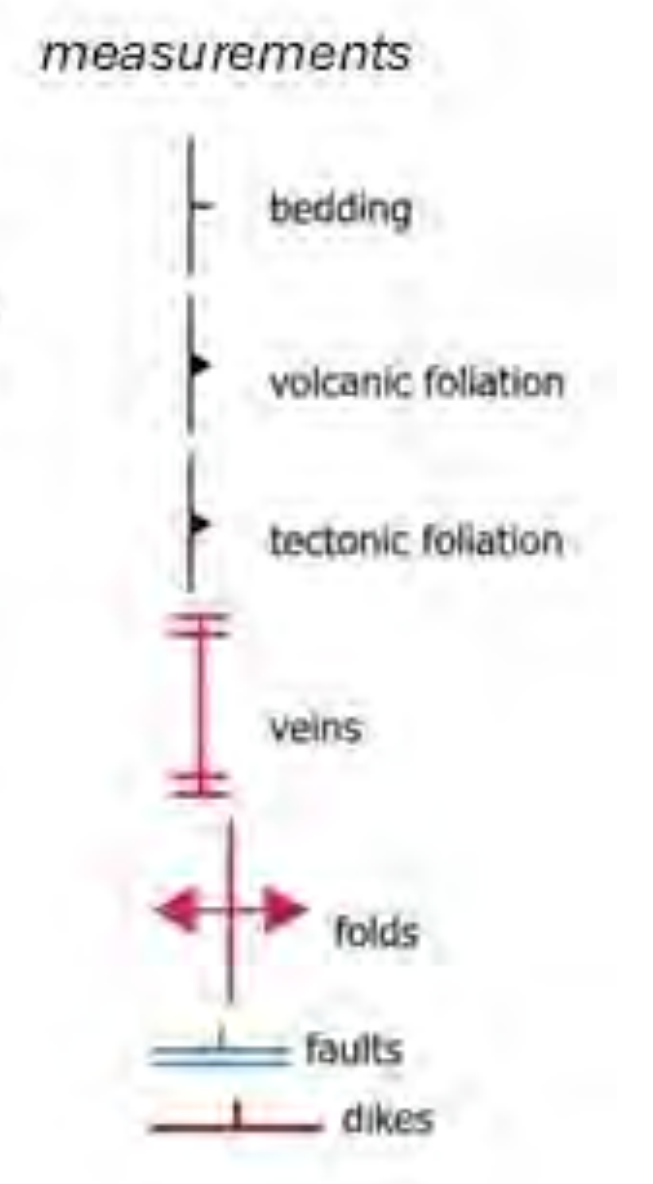
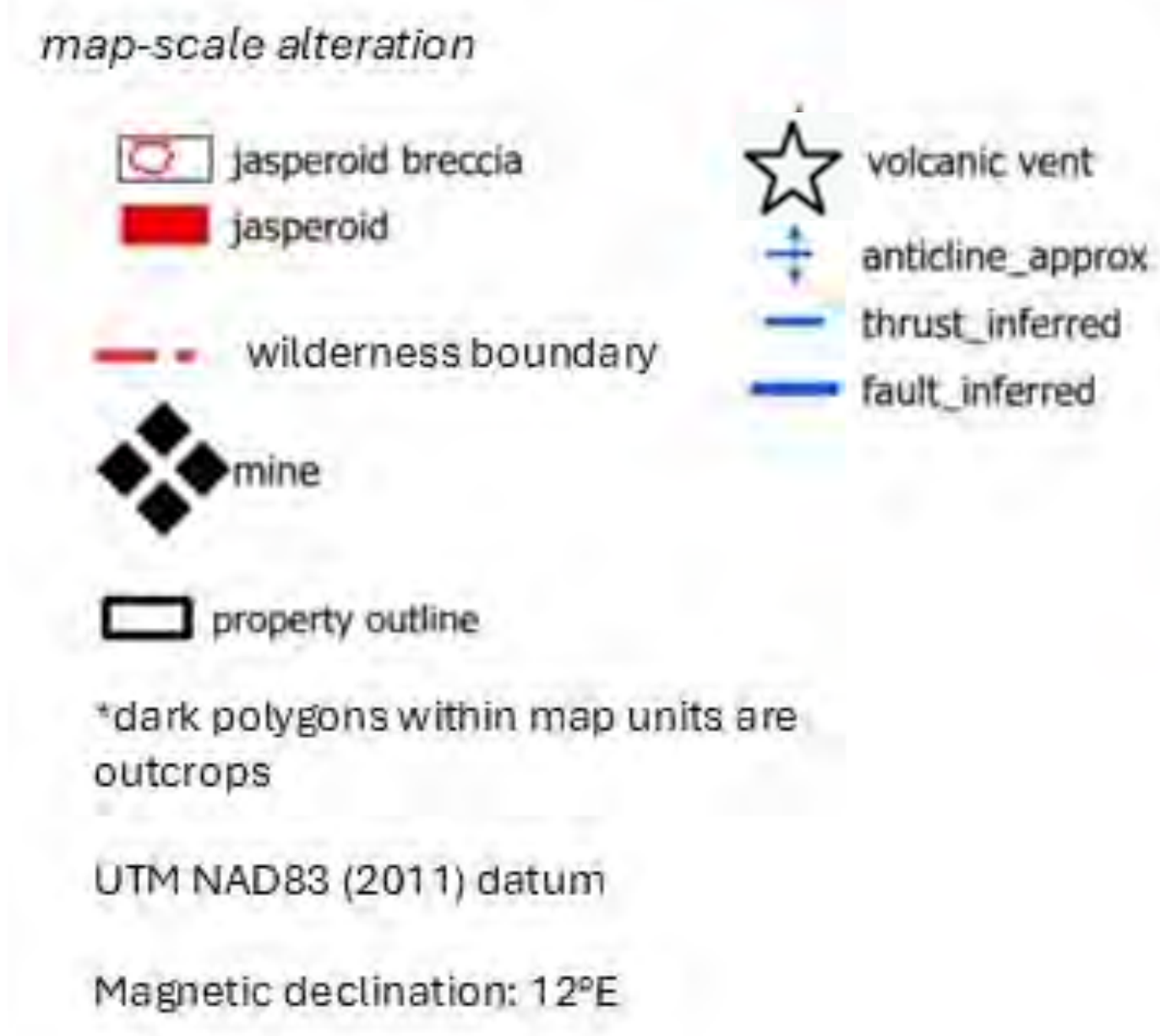
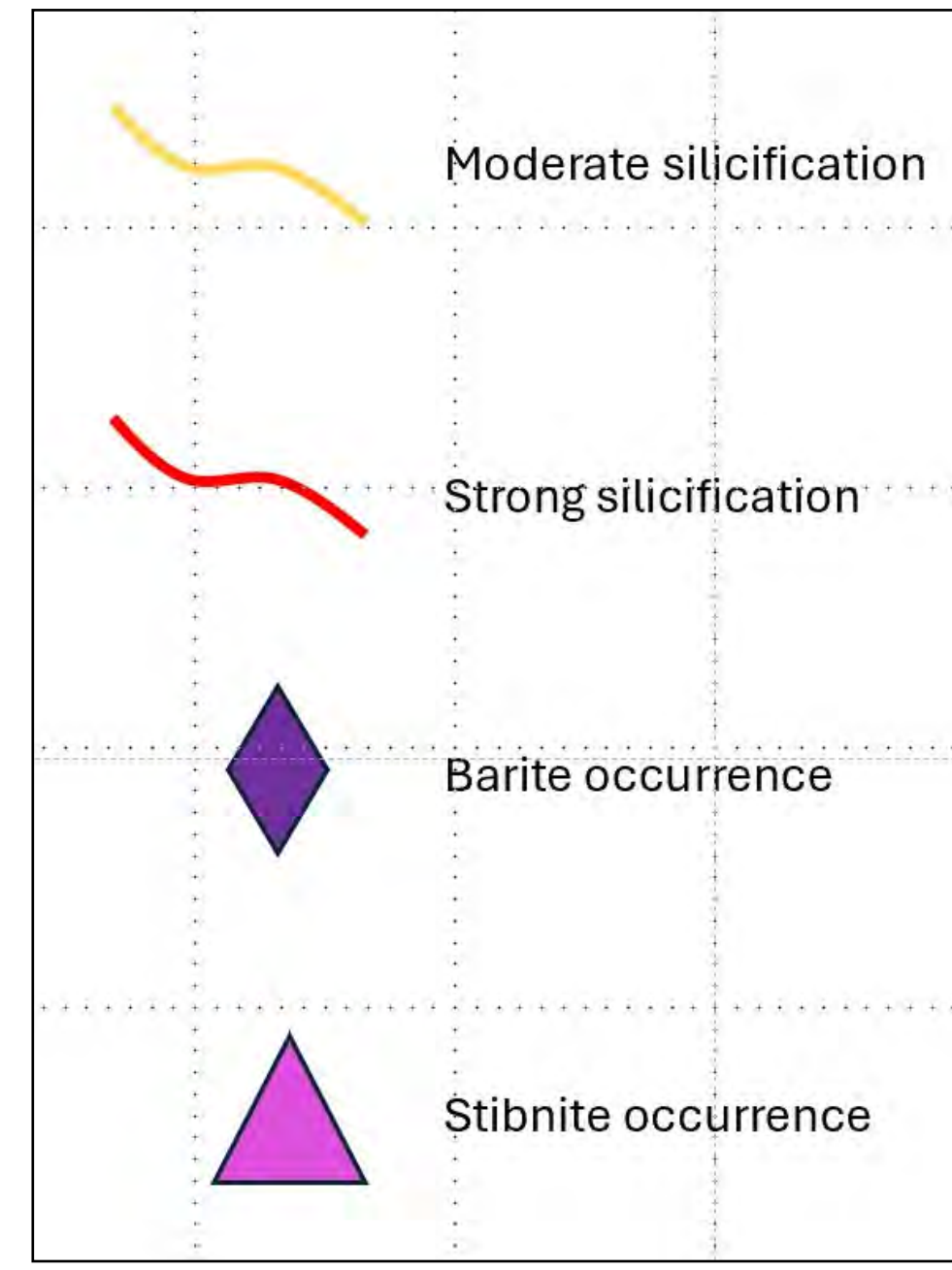
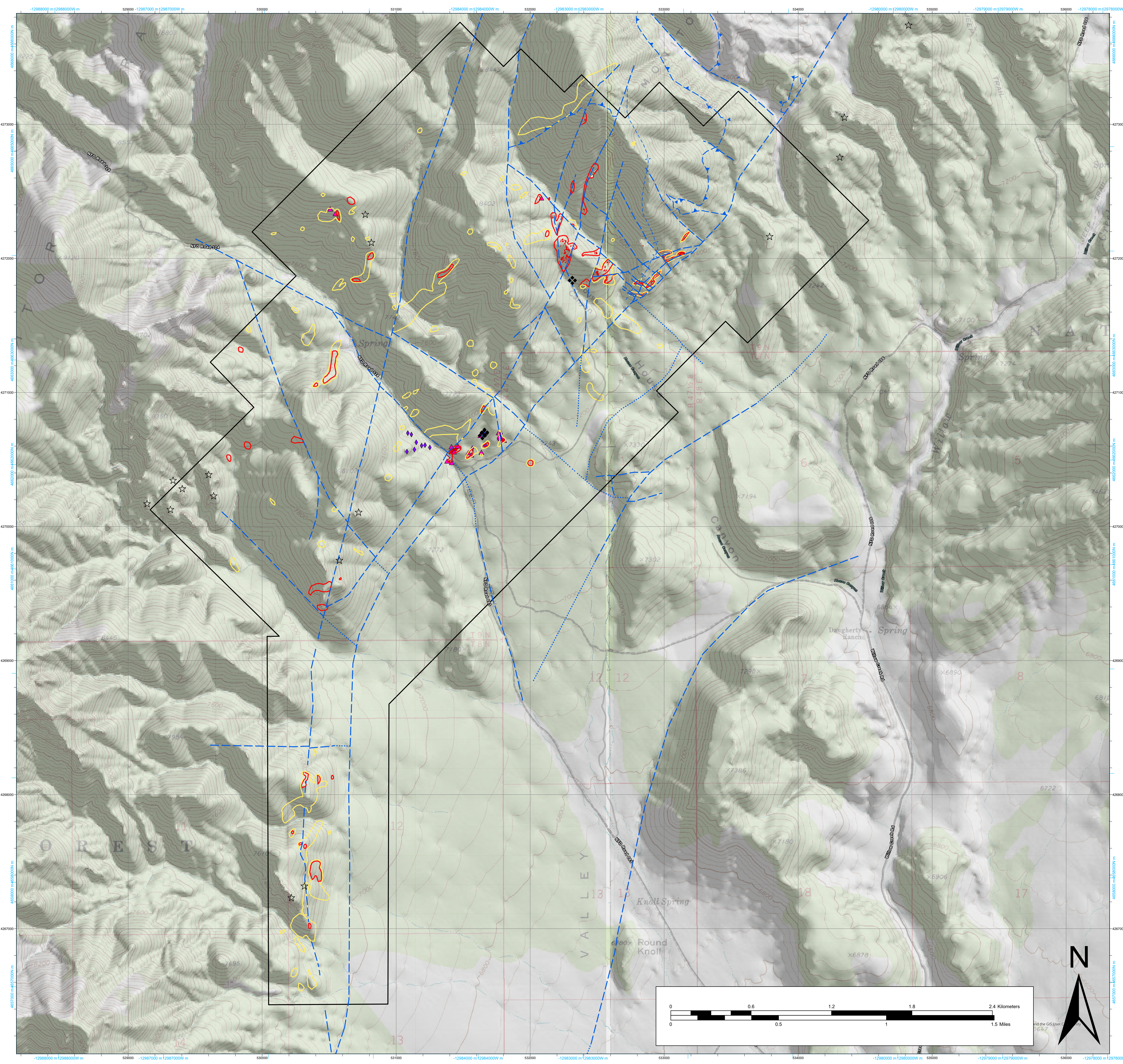
measurements

-  volcanic vent
-  anticline_approx
-  thrust_inferred
-  fault_inferred
-  bedding
-  volcanic foliation
-  tectonic foliation
-  veins
-  folds
-  faults
-  dikes

Alteration map of the King Solomon project, Monitor Range, northern Nye County, Nevada

Sericitization

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025



Alteration map of the King Solomon project, Monitor Range, northern Nye County, Nevada

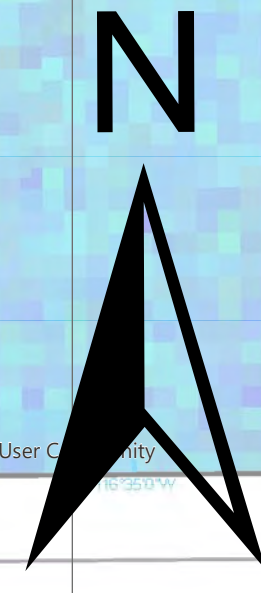
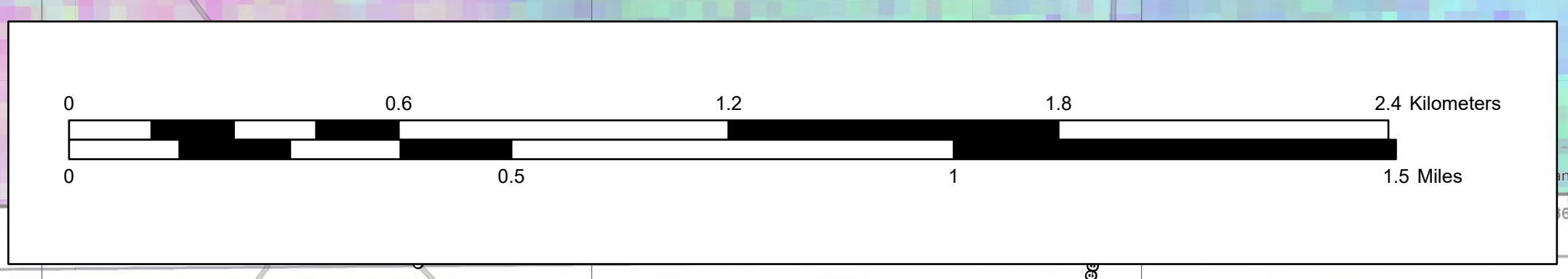
Silicification

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

Appendix F.
Geophysics

1. ASTER interpretation, 2006
2. CSAMT Resistivity 2006
3. Gravity Maps (compiled 2006, 2017, 2018)
4. IP (2006)
5. Magnetics (1995)

King Solomon Exploration Project,
Northern Nye County, Nevada
March 2026



- Moderate to strong sericitization
- Barite occurrence
- Stibnite occurrence

- map-scale alteration**
- jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
 - property outline
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E

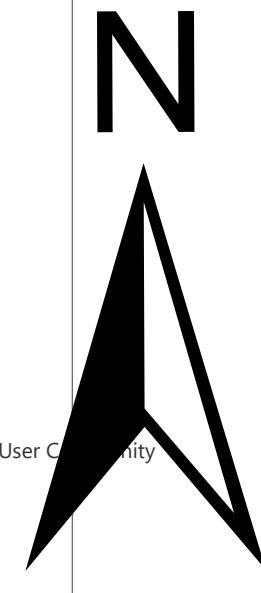
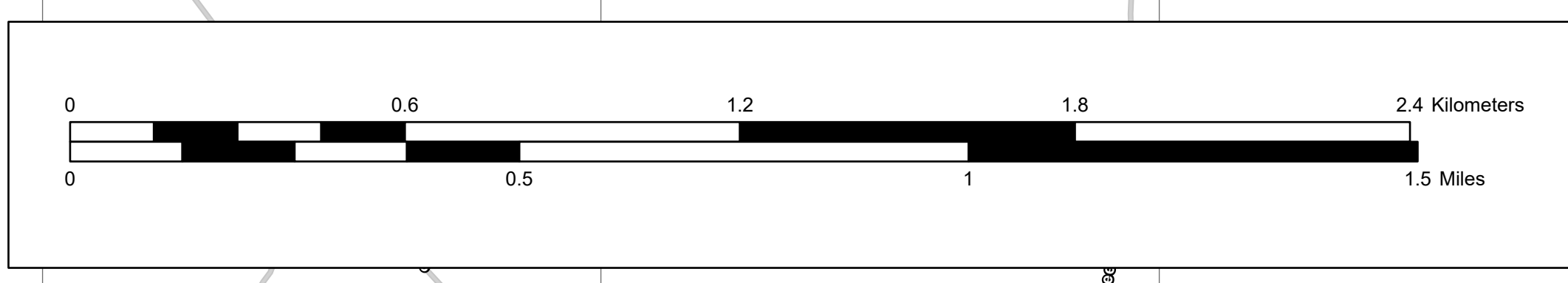
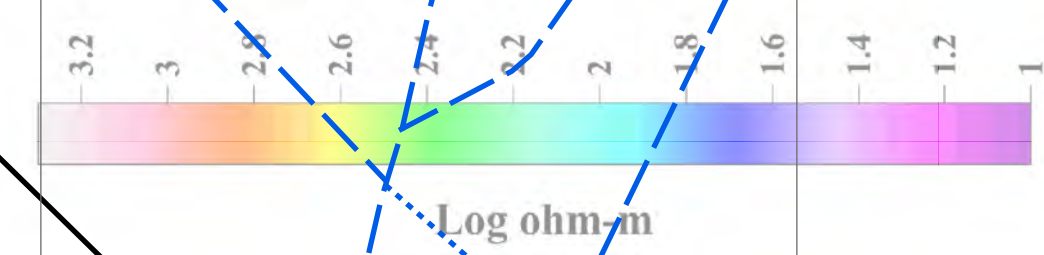
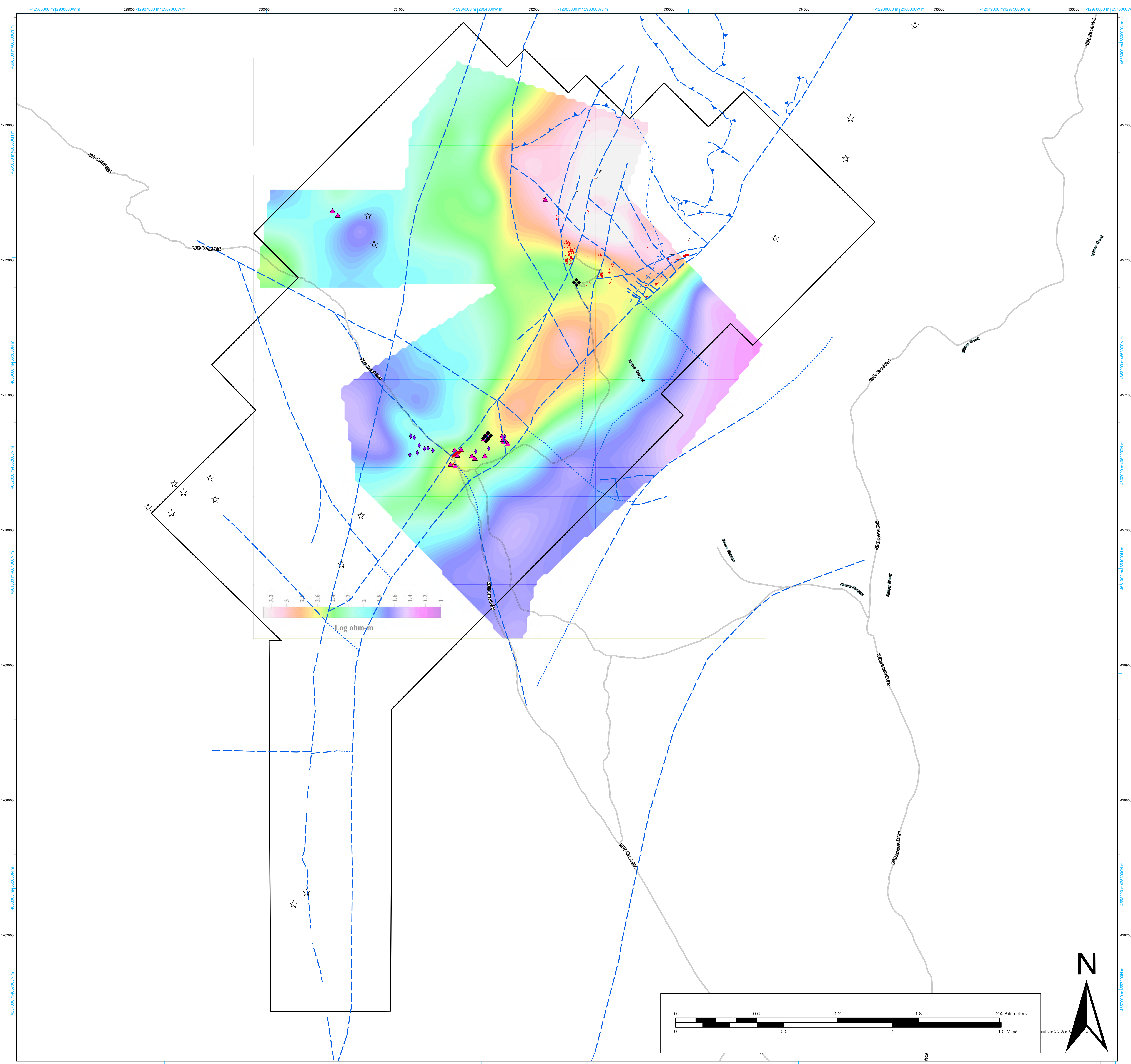
- measurements**
- volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - folds
 - faults
 - dikes

Alteration map of the King Solomon project, Monitor Range,
northern Nye County, Nevada

Sericitization

Aster SWIR Log Residual

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025



- Quaternary**
 - Qf_Qal_Qt Alluvial (Qal) and alluvial fan (Qaf) deposits; lesser talus apron deposits (Qt) on steeper slopes
 - Qoa Older alluvial and alluvial fan deposits present as perched terrace remnants at canyon mouths
 - Qls Landslide deposits
 - Tbt2 Upper beige biotite-phyric, fine pumice-bearing, vitrophyric rhyolite ash-flow tuff
 - Tbt Lower beige biotite-phyric, fine pumice-bearing vitrophyric ash-flow tuff
 - Trt Red tuff, welded pyroxene-bearing, coarse pumice-rich dacite ash-flow tuff
 - Ts2 White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks
 - Tfpr Stone Cabin rhyolite; yellow weathering, strongly flow foliated, sparsely porphyritic rhyolite lavas and domes
 - Ts White, planar-laminated to thin-bedded tuffaceous sedimentary rocks, generally altered and with Liesegang bands
 - Tpr Willow rhyolite. Spheroidally weathering, boldly outcropping, flow-foliated, maroon gray, plg>bio>qtz rhyolite lavas
 - Thr House rhyolite. Intrusive moderately porphyritic (plg>bio>qtz>san). May be intrusive equivalent to Tpr Willow lavas.
 - Tqprl Strongly flow foliated, white, altered, quartz-phyric rhyolite lavas associated with massive Tqpr dikes
 - Tqpr Massive, commonly wall-like exposures of spheroidally weathered, sparsely porphyritic qtz-rich rhyolite dikes
 - Tpbx Massive, speckled dark gray and white, polymictic lithic breccia with abundant rounded to subangular, coarse to fine, juvenile rhyolite and fine black sedimentary clasts. Groundmass consists of rock flour. Ubiquitous hydrothermal alteration. Associated with Tqpr rhyolite dikes
- Oligocene(?) King Solomon volcanic center**
 - Tcpt Older, maroon, hackly fractured, coarse to medium pumice-rich rhyolitic ash-flow tuff; multiple thin cooling units
- Oligocene(?) caldera-related ash-flow tuffs**
 - D-Osc Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. *Roberts Mountains or similar thrust*
 - Dd-stms Mostly tectonized, medium to light gray, hackly fractured, thin- to medium-bedded lime mudstone/wackestone, weathers to mottled gray and yellow. Possibly the Denay or Rabbit Hill formation.
 - Dd-ls Cliff-forming, massive, medium to pale gray carbonate packstone to grainstone; clasts rounded to 1cm, debris-flow carbonate; locally dolomitic. Possibly the lower Denay or Rabbit Hill formation.
 - SDrm Beige weathering, planar-laminated to thin-bedded, dark gray limestone and dolomitic limestone; has thin quartz-rich interlayers. Slope-former. Roberts Mountains Formation.
 - Ohc Massive, cliff-forming, yellowish-brown weathering, recrystallized, granular dolomite with considerable karst dissolution features. Hanson Creek Formation.
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 - Op Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

Structural features map of the King Solomon project, Monitor Range, northern Nye County, Nevada

CSAMT (resistivity, 2006)

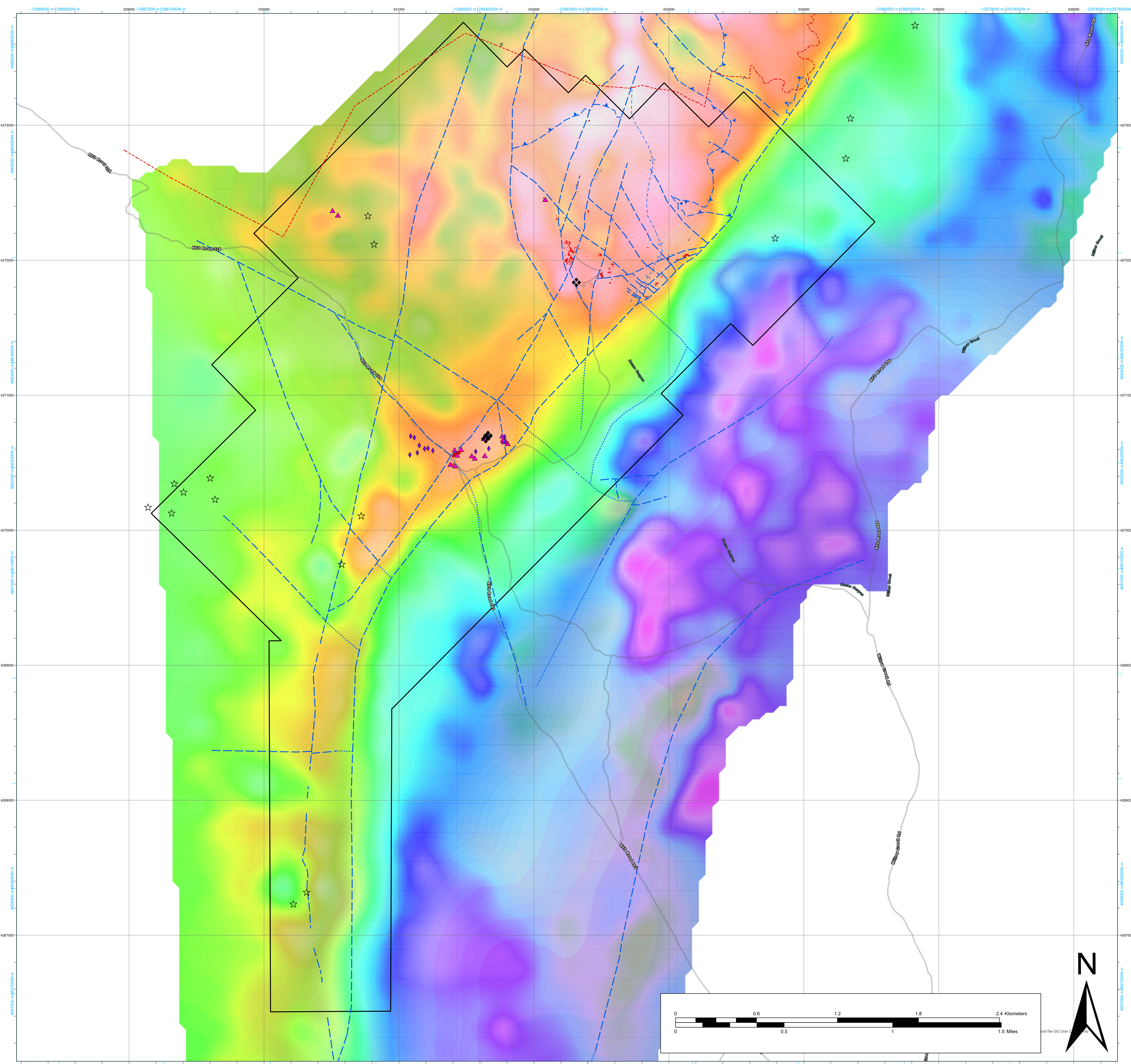
by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

map-scale alteration

- jasperoid breccia
- jasperoid
- wilderness boundary
- mine
- property outline
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E
- Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM
- Grid: UTM 1983, zone 11, 500 meters

measurements

- bedding
- volcanic foliation
- tectonic foliation
- veins
- folds
- faults
- dikes



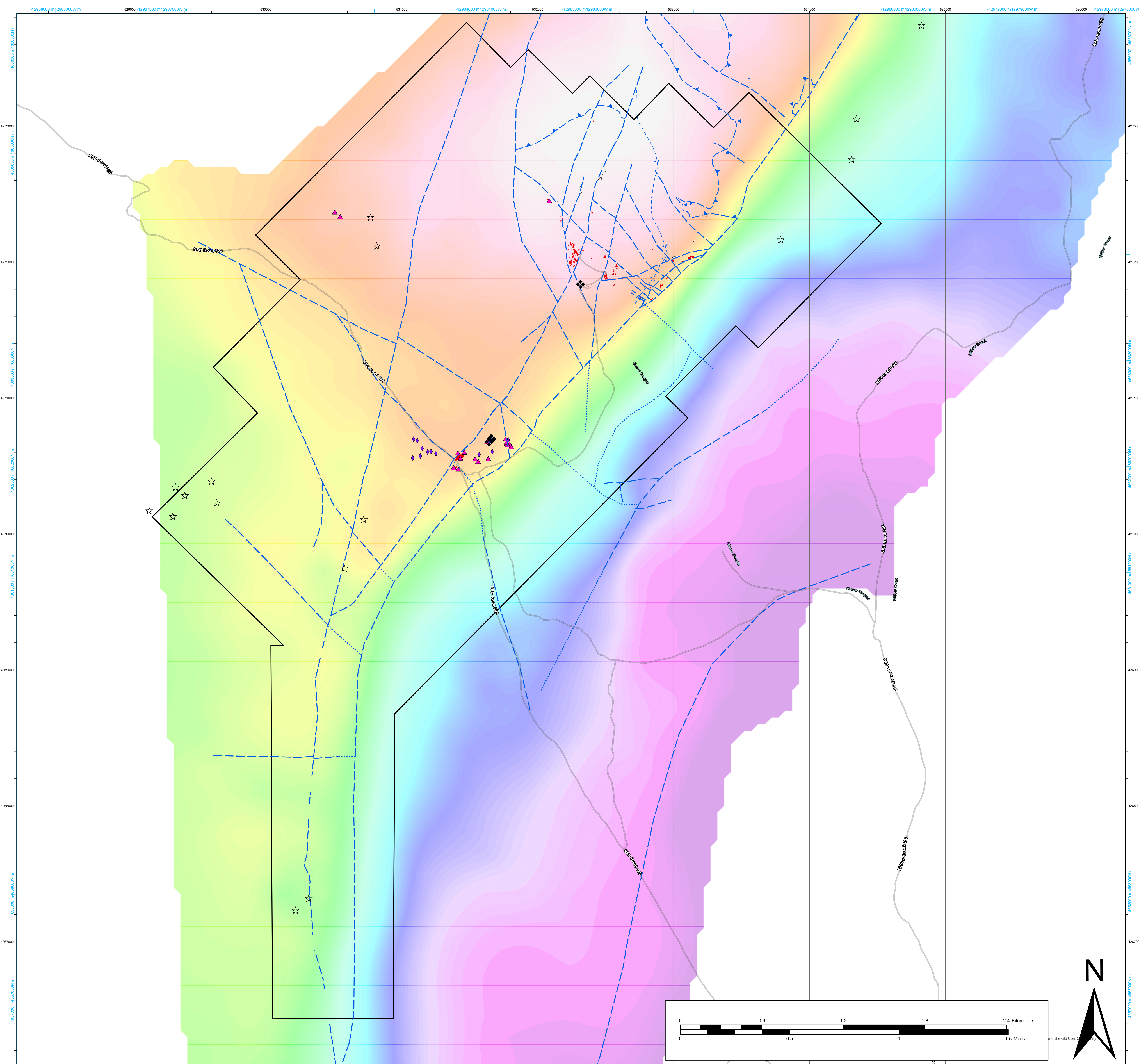
- Quaternary**
 - Qf_Qal_Qt Alluvial (Qal) and alluvial fan (Qaf) deposits; lesser talus apron deposits (Qt) on steeper slopes
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 - Qls Landslide deposits
 - Tbt2 Upper beige biotite-phyric, fine pumice-bearing, vitrophyric rhyolite ash-flow tuff
 - Tbt Lower beige biotite-phyric, fine pumice-bearing vitrophyric ash-flow tuff
 - Trt Red tuff, welded pyroxene-bearing, coarse pumice-rich dacite ash-flow tuff
 - Tts2 White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks
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 - Ohc Massive, cliff-forming, yellowish-brown weathering, recrystallized, granular dolomite with considerable karst dissolution features. Hanson Creek Formation.
 - Oe Massive, cliff-forming, spheroidally weathering, mottled pale gray with gray streaks, medium-grained quartzite, has distinct clear, rounded quartz grains. Eureka Quartzite.
 - Op Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

Structural features map of the King Solomon project, Monitor Range, northern Nye County, Nevada

Gravity - 1st Vertical Derivative (2017 compilation)

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

- map-scale alteration**
 - jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
 - property outline
- measurements**
 - volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - folds
 - faults
 - dikes
- *dark polygons within map units are outcrops
- UTM NAD83 (2011) datum
- Magnetic declination: 12°E
- Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM
- Grid: UTM 1983, zone 11, 500 meters



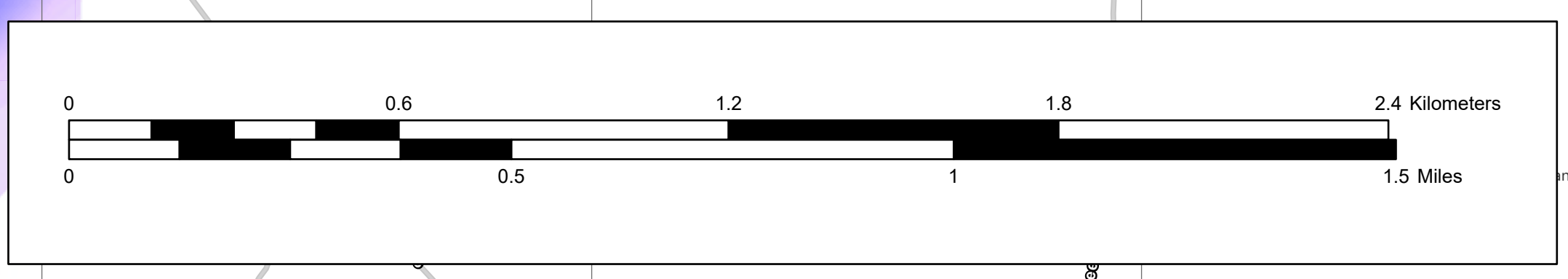
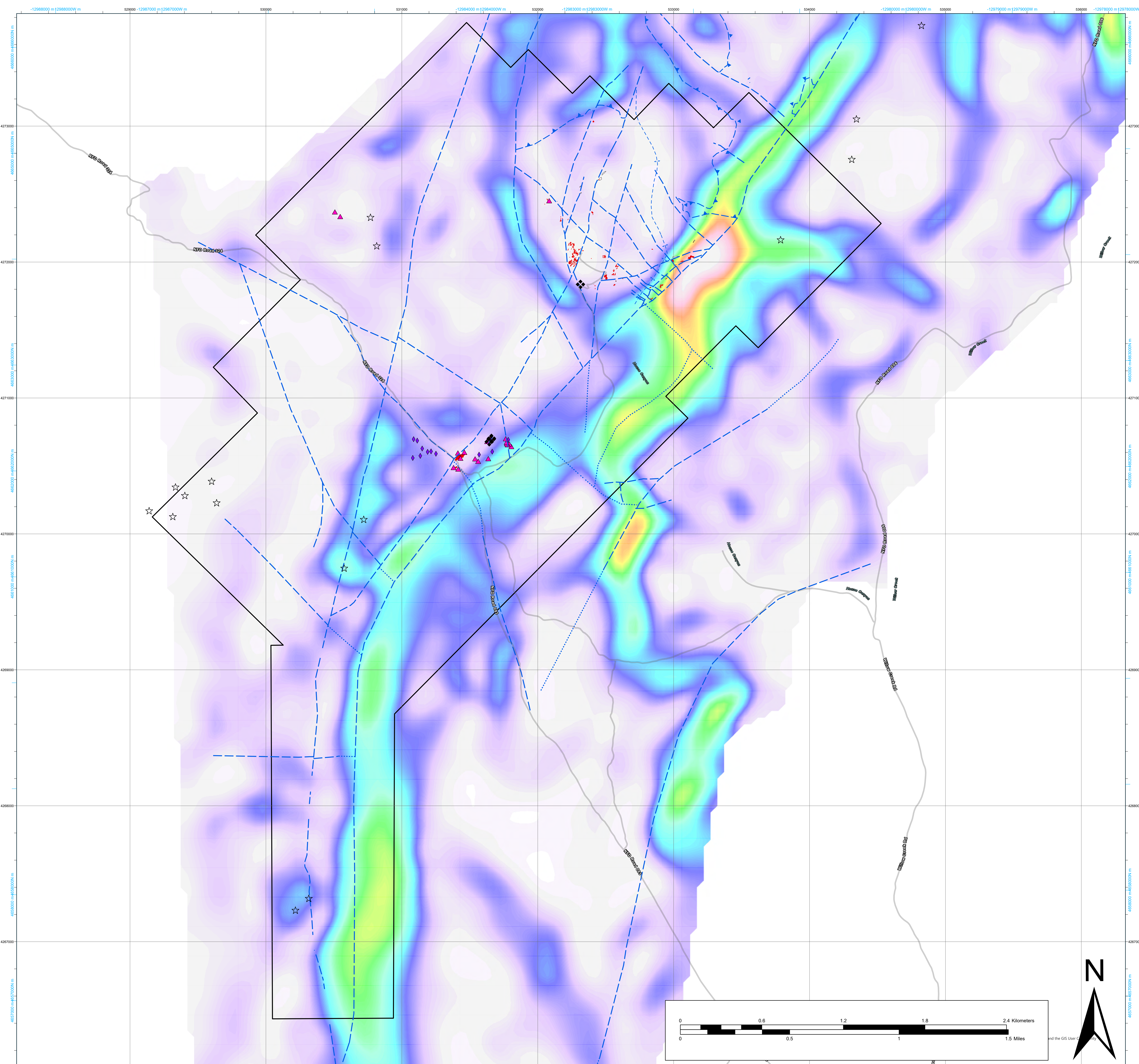
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 - Tbt Lower beige biotite-phyric, fine pumice-bearing vitrophyric ash-flow tuff
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 - Tts2 White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks
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 - Op Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

Structural features map of the King Solomon project, Monitor Range, northern Nye County, Nevada

Gravity – Complete Bouguer Anomaly (2017 compilation)

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

- map-scale alteration**
- jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
 - property outline
- measurements**
- volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - folds
 - faults
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- UTM NAD83 (2011) datum
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- Grid: UTM 1983, zone 11, 500 meters



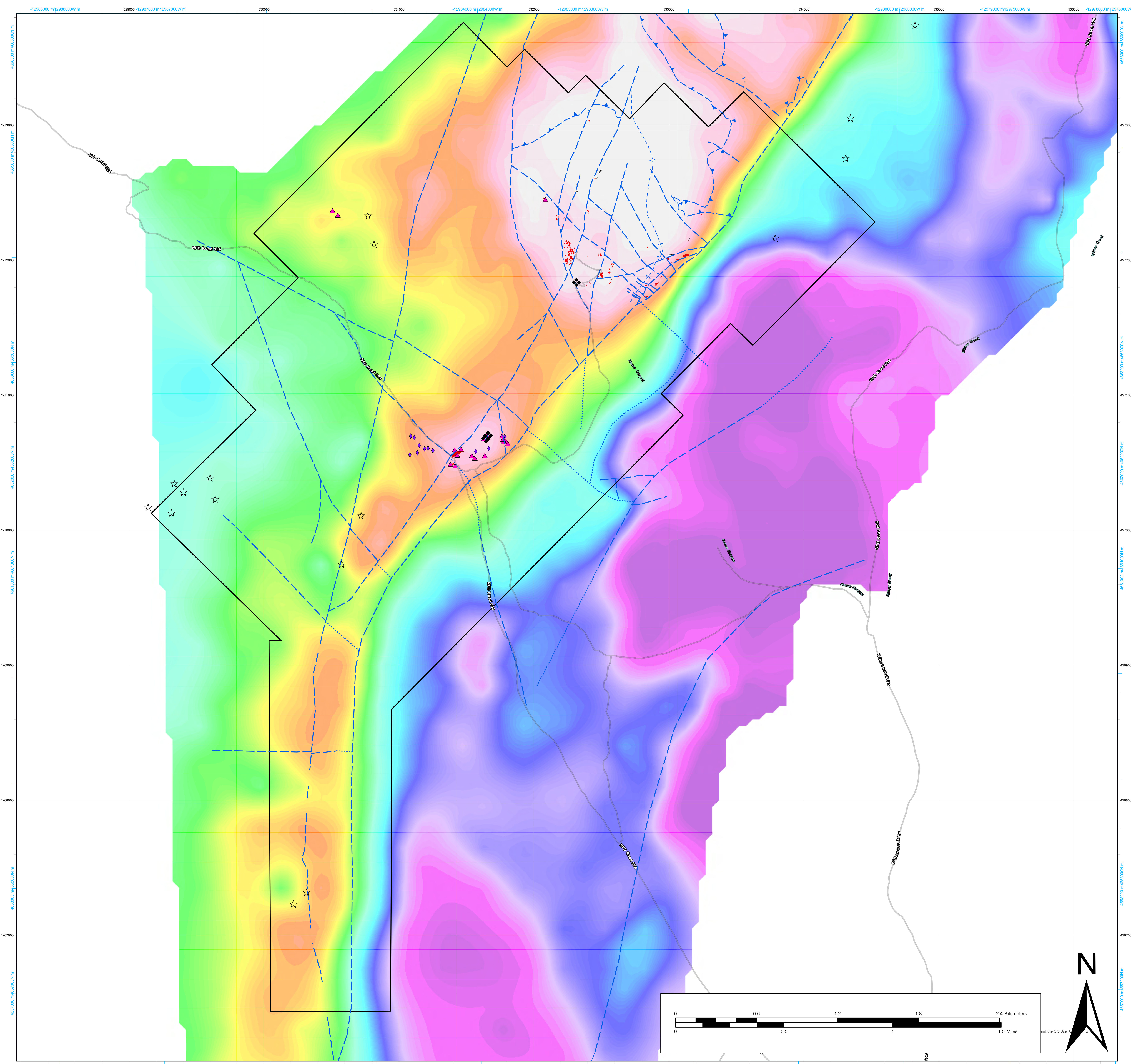
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Structural features map of the King Solomon project, Monitor Range, northern Nye County, Nevada

Gravity – Horizontal Gradient (2017 compilation)

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

- map-scale alteration**
- jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
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- measurements**
- volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
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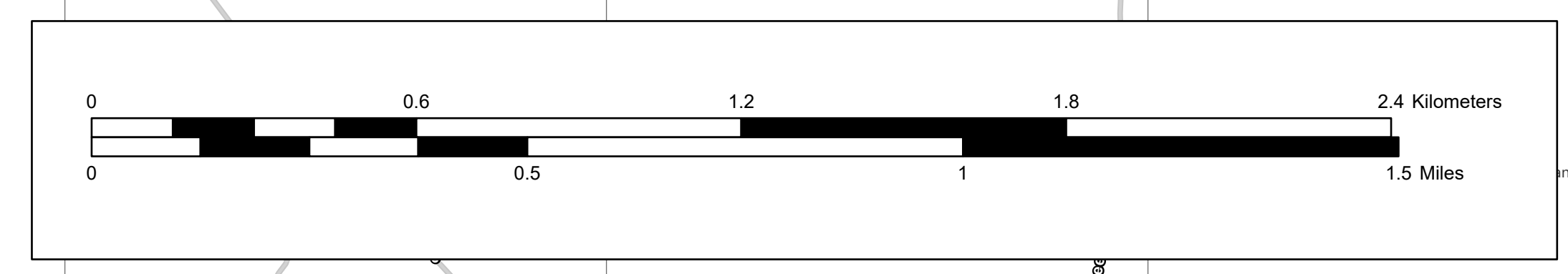
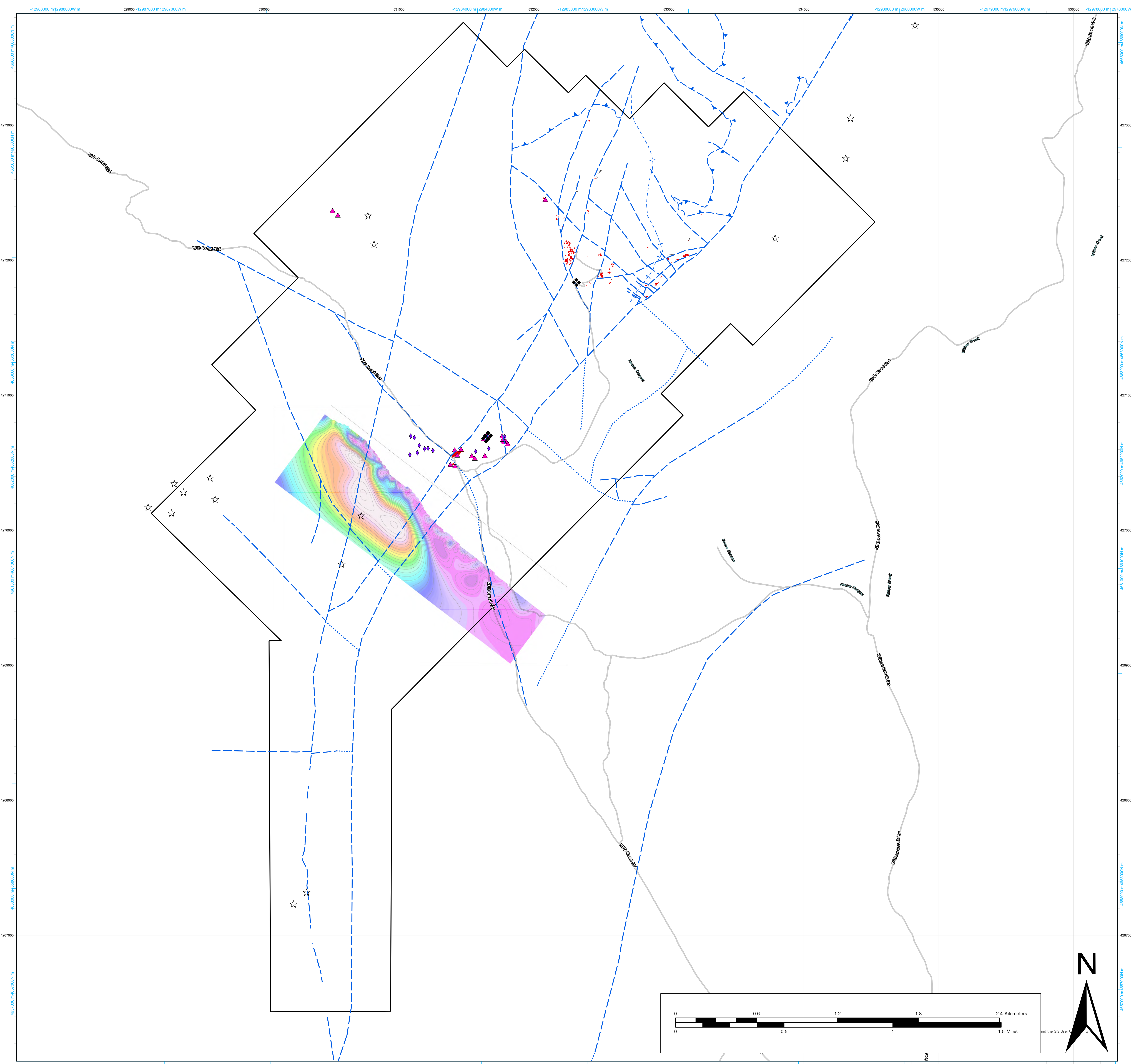
- Quaternary**
 - Qf_Qal_Qt Alluvial (Qal) and alluvial fan (Qaf) deposits; lesser talus apron deposits (Qt) on steeper slopes
 - Qoa Older alluvial and alluvial fan deposits present as perched terrace remnants at canyon mouths
 - Qls Landslide deposits
 - Tbt2 Upper beige biotite-phyric, fine pumice-bearing, vitrophyric rhyolite ash-flow tuff
 - Tbt Lower beige biotite-phyric, fine pumice-bearing vitrophyric ash-flow tuff
 - Trt Red tuff, welded pyroxene-bearing, coarse pumice-rich dacite ash-flow tuff
 - Tts2 White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks
 - Tfpr Stone Cabin rhyolite; yellow weathering, strongly flow foliated, sparsely porphyritic rhyolite lavas and domes
 - Ts White, planar-laminated to thin-bedded tuffaceous sedimentary rocks, generally altered and with Liesegang bands
 - Tpr Willow rhyolite. Spheroidally weathering, boldly outcropping, flow-foliated, maroon gray, plg>bio>qtz rhyolite lavas
 - Thr House rhyolite. Intrusive moderately porphyritic (plg>bio>qtz>san). May be intrusive equivalent to Tpr Willow lavas.
 - Tqpr1 Strongly flow foliated, white, altered, quartz-phyric rhyolite lavas associated with massive Tqpr dikes
 - Tqpr Massive, commonly wall-like exposures of spheroidally weathered, sparsely porphyritic qtz-rich rhyolite dikes
 - Tpbx Massive, speckled dark gray and white, polymictic lithic breccia with abundant rounded to subangular, coarse to fine, juvenile rhyolite and fine black sedimentary clasts. Groundmass consists of rock flour. Ubiquitous hydrothermal alteration. Associated with Tqpr rhyolite dikes
- Oligocene(?) King Solomon volcanic center**
 - Tcpt Older, maroon, hackly fractured, coarse to medium pumice-rich rhyolitic ash-flow tuff; multiple thin cooling units
- Oligocene(?) caldera-related ash-flow tuffs**
 - D-Osc Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. *Roberts Mountains or similar thrust*
 - Dd-stms Mostly tectonized, medium to light gray, hackly fractured, thin- to medium-bedded lime mudstone/wackestone, weathers to mottled gray and yellow. Possibly the Denay or Rabbit Hill formation.
 - Dd-ls Cliff-forming, massive, medium to pale gray carbonate packstone to grainstone; clasts rounded to 1cm, debris-flow carbonate; locally dolomitic. Possibly the lower Denay or Rabbit Hill formation.
 - SDrm Beige weathering, planar-laminated to thin-bedded, dark gray limestone and dolomitic limestone; has thin quartz-rich interlayers. Slope-former. Roberts Mountains Formation.
 - Ohc Massive, cliff-forming, yellowish-brown weathering, recrystallized, granular dolomite with considerable karst dissolution features. Hanson Creek Formation.
 - Oe Massive, cliff-forming, spheroidally weathering, mottled pale gray with gray streaks, medium-grained quartzite, has distinct clear, rounded quartz grains. Eureka Quartzite.
 - Op Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

**Structural features map of the King Solomon project, Monitor Range,
 northern Nye County, Nevada**

Residual Gravity (2017 compilation)

 by
 Michael W. Ressel
 for
 Phenom Resources Corp.
 October 2025

- map-scale alteration**
- jasperoid breccia
 - jasperoid
 - wilderness boundary
 - mine
 - property outline
- measurements**
- volcanic vent
 - anticline_approx
 - thrust_inferred
 - fault_inferred
 - bedding
 - volcanic foliation
 - tectonic foliation
 - veins
 - folds
 - faults
 - dikes
- *dark polygons within map units are outcrops
 UTM NAD83 (2011) datum
 Magnetic declination: 12°E
 Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM
 Grid: UTM 1983, zone 11, 500 meters



Structural features map of the King Solomon project, Monitor Range, northern Nye County, Nevada

Induced Polarization (chargeability, 2006)

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

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- Oligocene(?) caldera-related ash-flow tuffs**
 - D-Osc Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. *Roberts Mountains or similar thrust*
 - Dd-stms Mostly tectonized, medium to light gray, hackly fractured, thin- to medium-bedded lime mudstone/wackestone, weathers to mottled gray and yellow. Possibly the Denay or Rabbit Hill formation.
 - Dd-ls Cliff-forming, massive, medium to pale gray carbonate packstone to grainstone; clasts rounded to 1cm, debris-flow carbonate; locally dolomitic. Possibly the lower Denay or Rabbit Hill formation.
 - SDrm Beige weathering, planar-laminated to thin-bedded, dark gray limestone and dolomitic limestone; has thin quartz-rich interlayers. Slope-former. Roberts Mountains Formation.
 - Ohc Massive, cliff-forming, yellowish-brown weathering, recrystallized, granular dolomite with considerable karst dissolution features. Hanson Creek Formation.
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 - Op Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

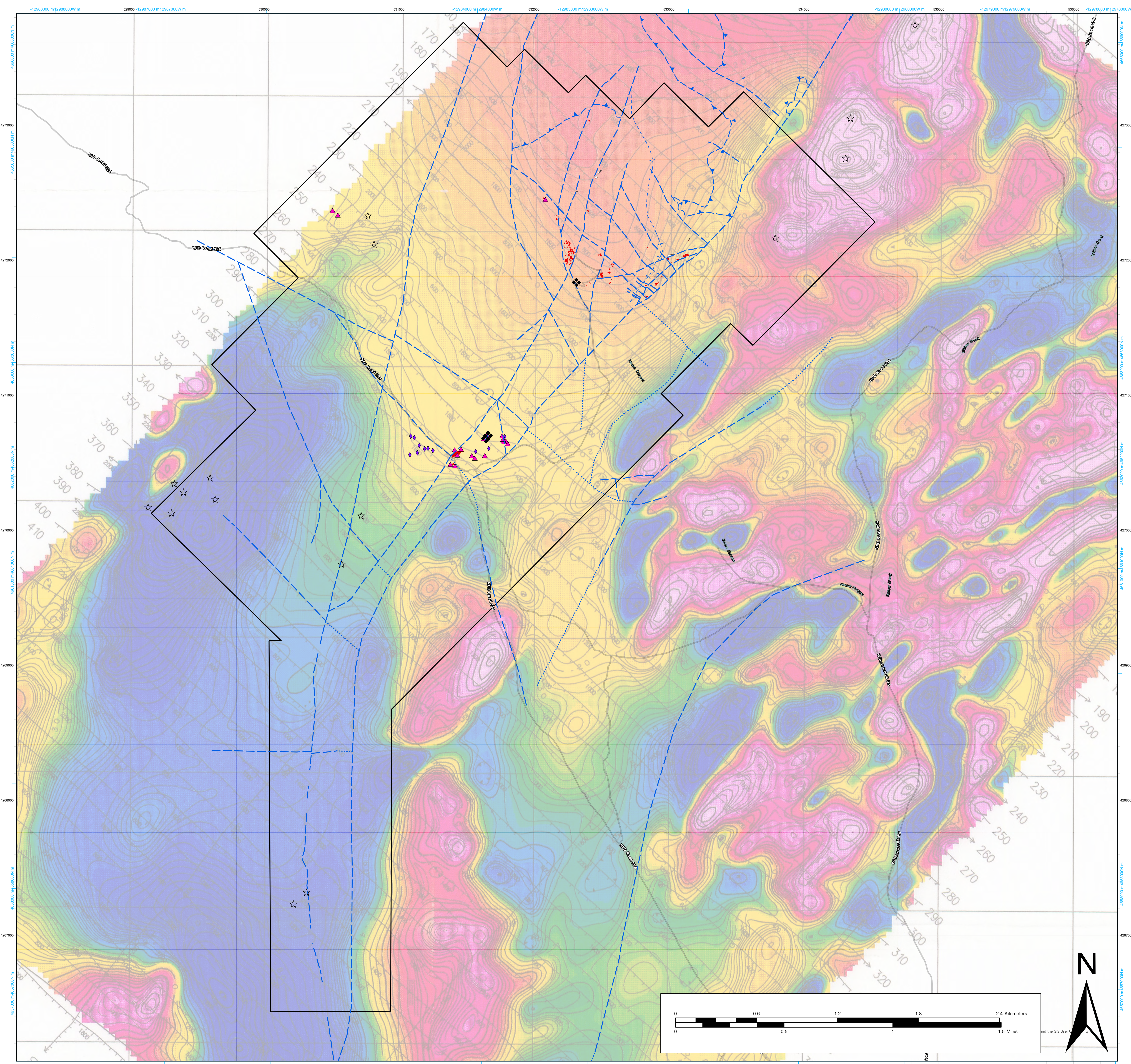
map-scale alteration

- jasperoid breccia
- jasperoid
- wilderness boundary
- mine
- property outline

*dark polygons within map units are outcrops
UTM NAD83 (2011) datum
Magnetic declination: 12°E
Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM
Grid: UTM 1983, zone 11, 500 meters

measurements

- bedding
- volcanic foliation
- tectonic foliation
- veins
- folds
- faults
- dikes



Quaternary	Qf_Qal_Qt	Alluvial (Qal) and alluvial fan (Qaf) deposits; lesser talus apron deposits (Qt) on steeper slopes
	Qoa	Older alluvial and alluvial fan deposits present as perched terrace remnants at canyon mouths
	Qls	Landslide deposits
	Tbt2	Upper beige biotite-phyric, fine pumice-bearing, vitrophyric rhyolite ash-flow tuff
	Tbt	Lower beige biotite-phyric, fine pumice-bearing vitrophyric ash-flow tuff
	Trt	Red tuff, welded pyroxene-bearing, coarse pumice-rich dacite ash-flow tuff
	Tts2	White, planar-laminated to thin-bedded, fine-grained, tuffaceous sedimentary rocks
	Tfpr	Stone Cabin rhyolite; yellow weathering, strongly flow foliated, sparsely porphyritic rhyolite lavas and domes
	Tts	White, planar-laminated to thin-bedded tuffaceous sedimentary rocks, generally altered and with Liesegang bands
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	Thr	House rhyolite. Intrusive moderately porphyritic (plg>bio>qtz>san). May be intrusive equivalent to Tpr Willow lavas.
	Tqprl	Strongly flow foliated, white, altered, quartz-phyric rhyolite lavas associated with massive Tqpr dikes
	Tqpr	Massive, commonly wall-like exposures of spheroidally weathered, sparsely porphyritic qtz-rich rhyolite dikes
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Oligocene(?) King Solomon volcanic center		
Oligocene(?) caldera- related ash- flow tuffs	Tcpt	Older, maroon, hackly fractured, coarse to medium pumice-rich rhyolitic ash-flow tuff; multiple thin cooling units
Paleozoic	D-Osc	Folded and sheared black siliceous mudstone and fine-grained red sandstone and siltstone with sparse gray micrite lenses. <i>Roberts Mountains or similar thrust</i>
	Dd-stms	Mostly tectonized, medium to light gray, hackly fractured, thin- to medium-bedded lime mudstone/wackestone, weathers to mottled gray and yellow. Possibly the Denay or Rabbit Hill formation.
	Dd-ls	Cliff-forming, massive, medium to pale gray carbonate packstone to grainstone; clasts rounded to 1cm, debris-flow carbonate; locally dolomitic. Possibly the lower Denay or Rabbit Hill formation.
	SDrm	Beige weathering, planar-laminated to thin-bedded, dark gray limestone and dolomitic limestone; has thin quartz-rich interlayers. Slope-former. Roberts Mountains Formation.
	Ohc	Massive, cliff-forming, yellowish-brown weathering, recrystallized, granular dolomite with considerable karst dissolution features. Hanson Creek Formation.
	Oe	Massive, cliff-forming, spheroidally weathering, mottled pale gray with gray streaks, medium-grained quartzite, has distinct clear, rounded quartz grains. Eureka Quartzite.
	Op	Platy, medium-gray, thin- to medium-bedded carbonate wackestone; slope-former. Pogonip Limestone.

Structural features map of the King Solomon project, Monitor Range,
northern Nye County, Nevada

Airborne Magnetics – Reduced to Pole (Kennecott, 1995)

by
Michael W. Ressel
for
Phenom Resources Corp.
October 2025

map-scale alteration

- jasperoid breccia
- jasperoid
- wilderness boundary
- mine
- property outline

measurements

- bedding
- volcanic foliation
- tectonic foliation
- veins
- folds
- faults
- dikes

☆ volcanic vent
 + anticline_approx
 - thrust_inferred
 - fault_inferred

*dark polygons within map units are outcrops

UTM NAD83 (2011) datum

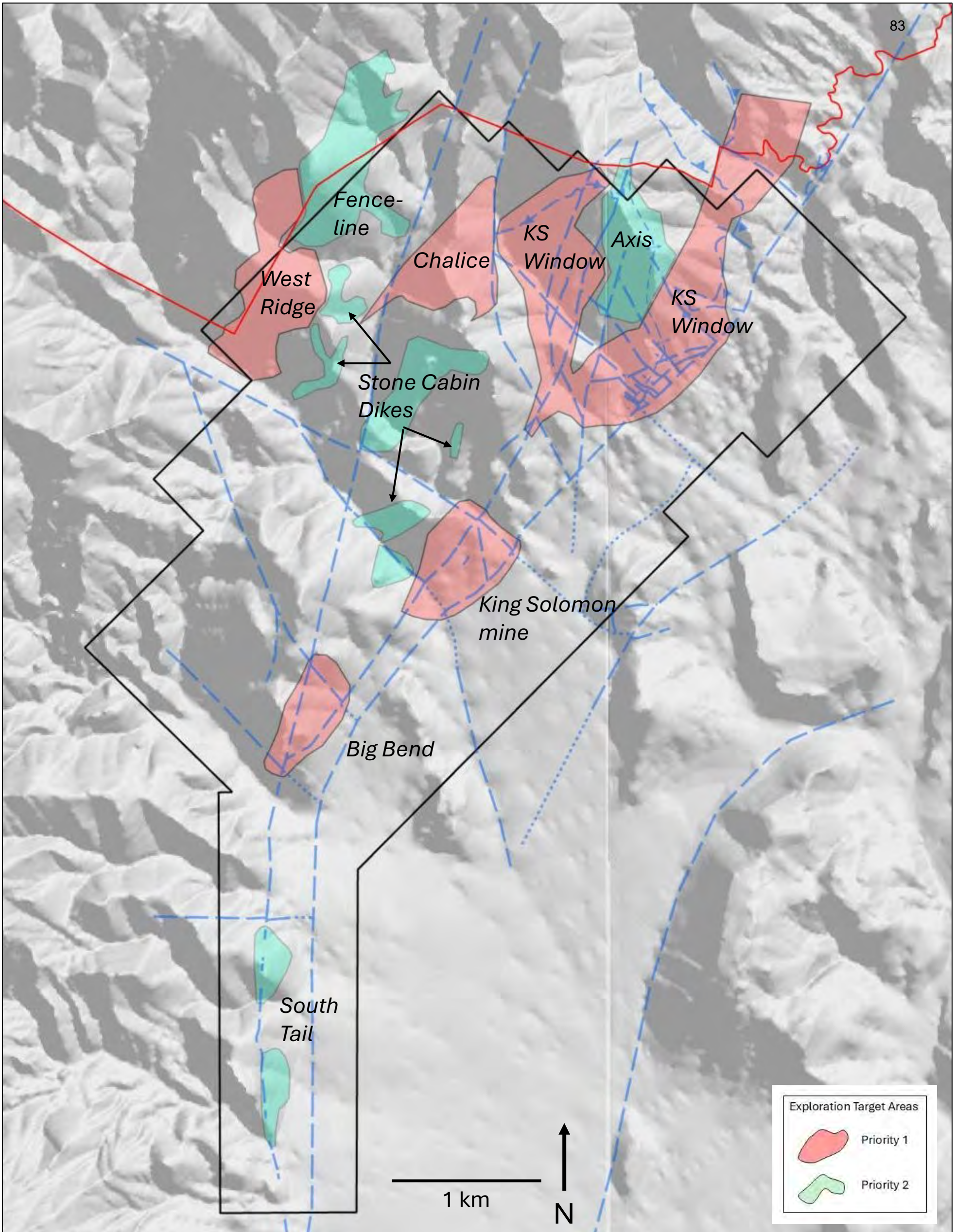
Magnetic declination: 12°E

Base maps: USGS Elkhorn Canyon and Eagle Pass 7.5-minute topographic and DEM

Grid: UTM 1983, zone 11, 500 meters

Appendix G.
Exploration Targets Map

King Solomon Exploration Project,
Northern Nye County, Nevada
March 2026



- King Solomon claim outline
- Table Mountain Wilderness boundary
- - - fault

Exploration Target Map
 King Solomon Project, Monitor Range,
 Northern Nye County, Nevada

 Phenom Resources Corp.
 April 2026

Exploration Target Areas	
	Priority 1
	Priority 2